



Idealized tests of ALADIN-NH dynamical kernel at very high resolution (comparison of H and NH versions)

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Presentation outline

Part I – basic facts about ALADIN-NH

- brief overview of current ALADIN-NH configuration
- important steps making this configuration possible
- future challenges

Part II – idealized 2D simulations

- orographic flow \Rightarrow robustness of ALADIN-NH dynamical kernel
- dry convection \Rightarrow limitations of hydrostatic approximation

Part I – basic facts about ALADIN-NH

Current ALADIN-NH configuration (1)

- fully compressible system
- mass based hybrid eta coordinate
- two additional prognostic variables derived from NH pressure departure and vertical divergence:

$$\hat{q} = \ln\left(\frac{p}{\pi}\right) \qquad d_4 = d_3 + X$$

- hydrostatic pressure, $d_3 = \frac{\partial w}{\partial z}, \ X = -\frac{\partial \mathbf{v}}{\partial z} \cdot \nabla_{\eta} z$

• representation of fields:

 π

horizontal	spectral	(unstaggered grid)
vertical	finite difference	(Lorenz type staggering)

• 2TL timestepping with SL advection

Current ALADIN-NH configuration (2)

- ICI scheme (Iterated Centered Implicit) using SI model as preconditioner
- resting isothermal SI background leading to solver with constant coefficients
- recomputation of SL trajectories in each iteration
- possibility of truly 2TL scheme (i.e. non-extrapolating)

Steps towards current configuration (1)

Starting line for listed developments was first ALADIN-NH version, finished in 1994 (3TL SI scheme with iteration of *X*-term, Eulerian advection). Attempts to implement SL advection faced instability, which was particularly severe for 2TL scheme. Problem was solved in several steps during period 1999–2005:

- choice of NH prognostic variables \hat{q}, d_3 based on linear stability analysis \Rightarrow removal of some instabilities related to imperfect SI background
- relaxation of angular momentum conservation \Rightarrow removed instability rising from inconsistent formulation of linear and non-linear models
- introduction of prognostic variable $d_4 = d_3 + X \Rightarrow$ reduced orographic instability in non-linear model
- ICI scheme \Rightarrow further stabilisation of non-linear model, making it sufficiently robust for operational use

Steps towards current configuration (2)

- relaxation of SI operator by introducing concept of acoustic background temperature ⇒ removed thermal instability for SL2TL SI scheme
- advection of $w \Rightarrow$ removal of SL chimney, clean bubble simulations
- diagnostic BBC \Rightarrow removal of SL chimney, achieved more consistently with ALADIN-NH code design
- SLHD scheme (Semi-Lagrangian Horizontal Diffusion) \Rightarrow removal of HD chimney by replacing spectral diffusion applied on $d_{3,4}$ with gridpoint one

All these steps led to final goal – sufficiently robust and accurate SL2TL ICI scheme with only one iteration needed.

ALADIN slang – SL chimney (1)

non-linear potential flow, vertical velocity \boldsymbol{w}



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ALADIN slang – SL chimney (2)

vertical velocity difference (left minus right)



ALADIN slang – HD chimney (1)

non-linear potential flow, vertical velocity w



ALADIN slang – HD chimney (2)

vertical velocity difference (left minus right)



Future challenges

- more transparent lateral boundary treatment what can be achieved in spectral model? (not strictly NH problem)
- bottom boundary treatment consistent with diabatic tendencies after SL and HD chimneys, will we face also Φ -sics one?
- non-reflective upper boundary condition RUBC or something else?
- performance above steep orography any bad surprises hidden? (common topic with HIRLAM)
- advection of w yes or no? (currently implemented only for SL2TL non-extrapolating scheme)
- vertical finite element discretization will it make ALADIN-NH dynamical kernel usable for ECMWF? (common topic with HIRLAM)

Part II – idealized 2D simulations

Case 1 – orographic flow

Model robustness will be demonstrated using non-linear potential flow. Reasons for this choice were following:

- purely NH regime (only trapped waves, excluded from H solution)
- strong orographic forcing, non-linear response
- analytical solution can be constructed iteratively
- simple enough solution, easier detection of model deficiencies
- little sensitivity to treatment of upper and lateral boundaries

Experimental setup

• background state: $\bar{\pi}_S = 101325 \text{ Pa}$ $\bar{u} = 15 \text{ ms}^{-1}$ (constant wind)

• bell shaped mountain:
$$a = 100 \text{ m} \text{ (halfwidth)}$$

 $h = 100 \text{ m} \text{ (height)}$

• stationary final state obtained by short integration (t = 200 s) with time constant LBC

Basic ALADIN-NH configuration

• non-extrapolating SL2TL ICI scheme with 1 iteration

• SI background:
$$T^* = 300 \text{ K}$$

 $\pi^*_a = 100 \text{ K}$
 $\pi^*_S = 90\,000 \text{ Pa}$

- timestep $\Delta t = 1.0 \, \mathrm{s}$
- prognostic variable d_4
- diagnostic BBC
- diffusion of SLHD type
- lateral coupling with background state, sponge

Aspects to be tested

- stability and accuracy, behaviour for long timesteps
- impact of prognostic variable d_4
- impact of acoustic background temperature T_a^*

Stability and accuracy versus timestep (1)

vertical velocity w



Stability and accuracy versus timestep (2)

vertical velocity w



Impact of prognostic variable d_4



variable d_3 requires 3 iterations to become stable

Impact of acoustic background temperature T_a^*



without T_a^* , SL2TL SI scheme blows after 9 timesteps

with T_a^* , SL2TL SI scheme is stable

Summary for the case 1

- tested ALADIN-NH configuration proved to be sufficiently accurate and stable
- there is little sensitivity to timestep length, response to orographic forcing is acceptable up to CFL \approx 5
- for ICI scheme with 1 iteration, prognostic variable d_4 is necessary to remove orographic instability
- acoustic background temperature T_a^* removes thermal instability for SL2TL SI scheme, its importance diminishes for ICI scheme
- no detrimental effects of non-extrapolating scheme were detected

Case 2 – dry convection

Why convective case?

- convection is driven by local imbalance between buoyancy and gravity, which is non-hydrostatic effect
- comparison of H and NH simulations at different resolutions enables to limit validity of hydrostatic approximation
- explicit convection is very sensitive numerical test

Complications:

- analytical solution is not known
- \bullet advection of w had to be used to get clean solution

Common experimental setup

- resting initial state (u, w = 0) with horizontally constant surface pressure $\pi_S = 101325 \text{ Pa}$
- neutral background stratification with $\bar{\theta} = 300 \,\mathrm{K}$
- cold bubble perturbation in initial θ field ($\theta = \overline{\theta} + \theta'$)
- flat orography
- horizontally periodic domain, no coupling

Simulations at 10 m horizontal resolution

- \bullet domain $1.2\times1.2\,\text{km},$ additional 30 layers at the top with isothermal stratification
- resolution: $\begin{array}{l} \Delta x = 10 \, \mathrm{m}, \ \mathrm{quadratic} \ \mathrm{truncation} \\ \Delta z = 10 \, \mathrm{m} \end{array}$
- \bullet circular bubble with r = 150 m and $\theta'_{\rm max}$ = $-0.5\,{\rm K}$ placed 600 m above ground
- ALADIN configuration as in case 1, except:
 - timestep $\Delta t = 0.4 \,\mathrm{s}$
 - diagnostic BBC replaced by advection of \boldsymbol{w}
 - hydrostatic simulations with extrapolating SL2TL SI scheme

hydrostatic

non-hydrostatic



 $\Delta x = 10 \,\mathrm{m}, t = 0 \,\mathrm{s}$

hydrostatic

non-hydrostatic



 $\Delta x = 10 \,{\rm m}, t = 20 \,{\rm s}$

hydrostatic

non-hydrostatic



 $\Delta x = 10 \,{\rm m}, t = 50 \,{\rm s}$

hydrostatic

non-hydrostatic



 $\Delta x = 10 \, \text{m}, t = 100 \, \text{s}$

hydrostatic

non-hydrostatic



 $\Delta x = 10 \, \text{m}, t = 200 \, \text{s}$

hydrostatic

non-hydrostatic



 $\Delta x = 10 \,\mathrm{m}, t = 400 \,\mathrm{s}$

hydrostatic

non-hydrostatic



 $\Delta x = 10 \, \text{m}, t = 600 \, \text{s}$

Are the numerical results trustable?

- convection and turbulence are non-linear phenomena consisting of multiple interacting scales
- in reality, subgrid scales can have significant influence on resolved ones, which leads to the necessity of their parameterization
- simple test to judge if this is the case is to increase horizontal resolution and redo the simulation

perturbation of potential temperature $\theta - \overline{\theta}$

 $\Delta x = 10 \text{ m}$

 $\Delta x = 5 \,\mathrm{m}$



non-hydrostatic run, $t = 600 \, \text{s}$

Simulations at 1 km horizontal resolution

- domain 120×1.2 km, additional 30 layers at the top with isothermal stratification
- resolution: $\begin{array}{lll} \Delta x = & 1 \, \mathrm{km}, \ \mathrm{quadratic} \ \mathrm{truncation} \\ \Delta z = & 10 \, \mathrm{m} \end{array}$
- initial bubble stretched to ellipse with a = 15 km and b = 150 m
- 10 times stronger initial perturbation ($\theta'_{max} = -5 \text{ K}$)
- timestep $\Delta t = 40 \,\mathrm{s}$
- other settings unchanged

hydrostatic

non-hydrostatic



 $\Delta x = 1 \,\mathrm{km}, t = 0 \,\mathrm{min}$

hydrostatic

non-hydrostatic



 $\Delta x = 1 \,\mathrm{km}, t = 10 \,\mathrm{min}$

hydrostatic

non-hydrostatic



 $\Delta x = 1 \,\mathrm{km}, t = 20 \,\mathrm{min}$

hydrostatic

non-hydrostatic



 $\Delta x = 1 \text{ km}, t = 30 \text{ min}$

hydrostatic

non-hydrostatic



 $\Delta x = 1 \,\mathrm{km}, t = 1 \,\mathrm{h}$



hydrostatic

non-hydrostatic



 $\Delta x = 1 \,\mathrm{km}, t = 2 \,\mathrm{h}$

How trustable are numerical results now?

- with bigger Δx influence of subgrid scales might become stronger
- it is necessary to repeat the test with increased horizontal resolution

 $\Delta x = 1 \,\mathrm{km}$

 $\Delta x = 0.5 \,\mathrm{km}$



non-hydrostatic run, $t = 0 \min$

 $\Delta x = 1 \,\mathrm{km}$

 $\Delta x = 0.5 \,\mathrm{km}$



non-hydrostatic run, t = 10 min

 $\Delta x = 1 \,\mathrm{km}$





non-hydrostatic run, t = 20 min

 $\Delta x = 1 \,\mathrm{km}$





non-hydrostatic run, t = 30 min

 $\Delta x = 1 \,\mathrm{km}$





non-hydrostatic run, t = 1 h

 $\Delta x = 1 \,\mathrm{km}$





non-hydrostatic run, t = 2 h

Summary for the case 2

- ALADIN-NH kernel (with advection of w) passed also cold bubble test
- at kilometric horizontal resolution, subgrid convective scales can still play an important role
- slight differences between H and NH simulations start to be visible
- it can be expected that these differences will become more pronounced for deep convective systems with diabatic forcing
- going to 10 m horizontal resolution, differences between H and NH simulations increase ⇒ NH model becomes compulsory
- at the same time influence of subgrid scales diminishes, they can be satisfactorily represented by viscous dissipation

Conclusions

- ALADIN-NH dynamical kernel is now in stable state, ready for operational use
- there are still some details to be solved (advection of w, diabatic tendencies in BBC, non-reflective upper boundary treatment, ...)
- shape of the kernel can change with implementation of vertical finite elements, but its general concept will most likely remain untouched
- idealized bubble simulations show that non-hydrostatic effects start to be visible at kilometric horizontal resolution and they completely take over for 10 m horizontal resolution