

# Overview of data assimilation developments in LACE

Gergely Bölöni, Edit Adamcsek, Máté Mile, András Horányi, Mihály Szűcs, Alena Trojáková, Antonin Bucanek, Patrik Benacek, Xin Yan, Florian Meier, Tomislav Kovacic, Antonio Stanesic, Jure Cedilnik, Benedikt Strajnar, Mirela Niculae, Michal Nestiak, Maria Monteiro, Loïk Berre



# Outline

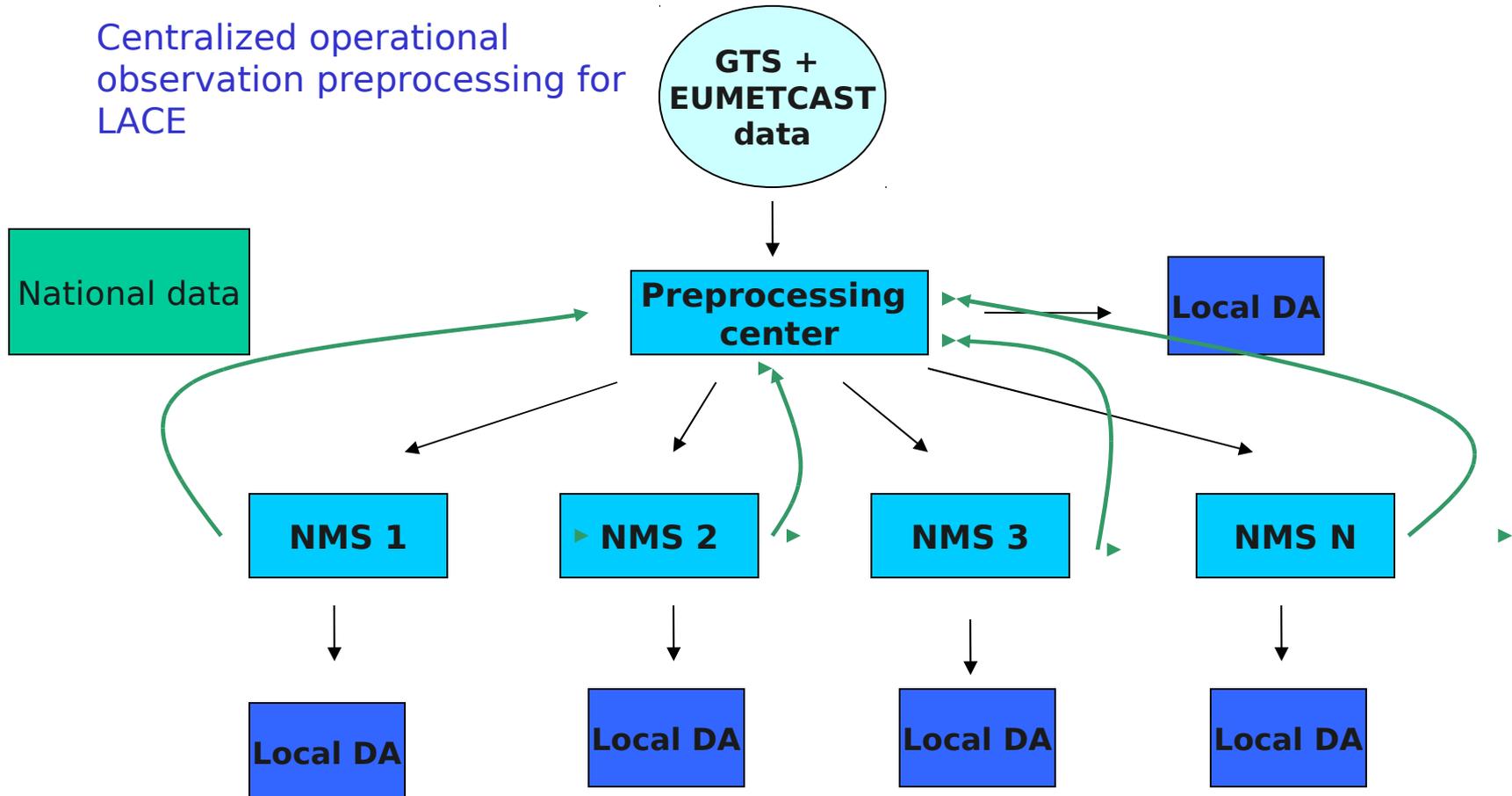
- Towards operational DA in LACE
- Diagnosis of Downscaled Global Ensemble DA and LAM Ensemble DA systems for sampling the B matrix
- Time variations of background error structures
- Observation, background and model error diagnosis in AROME
- Local surface perturbations in the LAMEPS system of Hungary

# Towards operational DA in LACE

- Regular DA runs in 5 countries: Austria, Croatia, Czech Rep., Hungary, Slovenia (Romania, Slovakia starts)
- Observations: OPLACE + local high-resolution SYNOP data
- Atm. 3DVAR + surf. OI (but wide range of cycling setups: IDFI, Blending of large scales, different LBCs)
- Improvements at 2m but not always above 850 hPa
- Room for improvement (LACE DA working days, 28-30 Sept., Ljubljana)
  - Order of the surf. (OI) and the atmospheric (VAR) analysis
  - Tune length scales in the surf. OI analysis
  - VARBC with daily cycling of the correction coefficients
  - Revise the use of AMSU channels
  - Jb/Jo tuning (more often, regularly)
  - Use mean orography instead of envelope

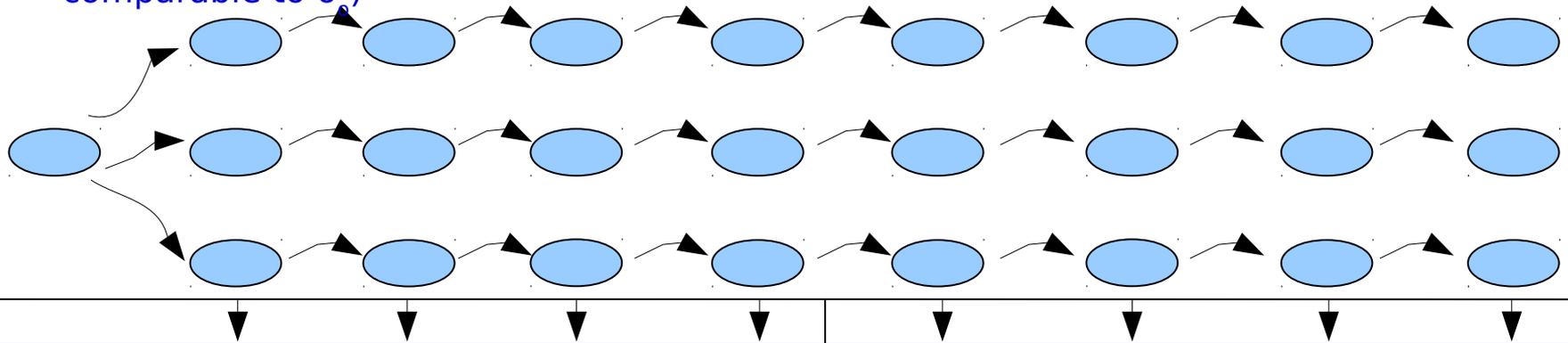
# Towards operational DA in LACE

Centralized operational observation preprocessing for LACE



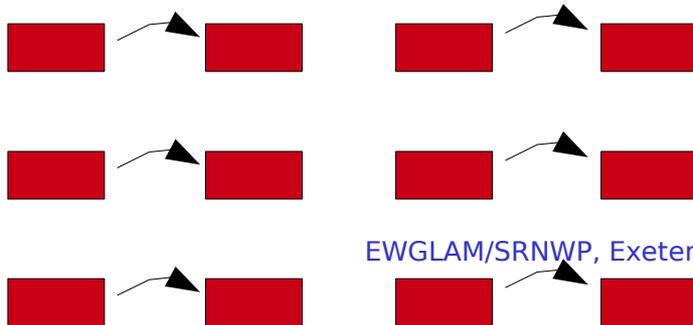
# Downscaled Global Ensemble DA versus LAM Ensemble DA

Global Ensemble Data Assimilation system (perturbed analyses through randomly perturbed observations with Gaussian distribution and rescaled amplitude comparable to  $\sigma$ )

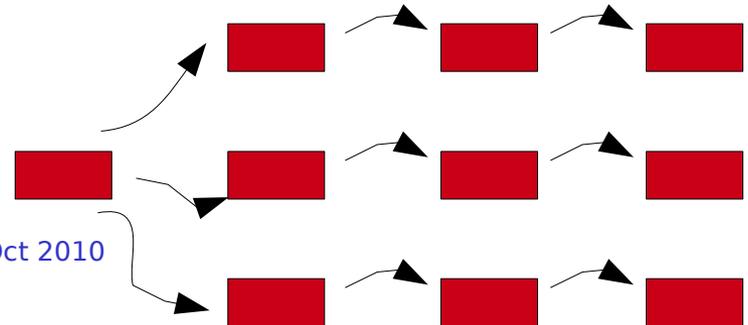


Perturbation through LBCs

**DscEnDa:** „Downscaling the large scale uncertainty” kept in the LBCs



**LamEnDa:** adding direct local uncertainty by perturbed LAM analysis



# Diagnosis of DscEnDa vs. LamEnDa for sampling the B matrix

What is the information we gain with LamEnDa vs. DscEnDa?

Diagnosis with PECA: Perturbation versus Error Correlation analysis  
(Wei and Toth, MWR, 2003)

$$PECA = (\epsilon_b, \epsilon_{ref})$$

$$\epsilon_b = X_b^{pert} - X_b^{ct}$$

$$\epsilon_{ref} = X_b^{ct} - X_a^{verif}$$

In what extent our ensemble differences explain the B error variance?

PECA defined in the ALADIN spectral space (after B cov computation by Berre, 2000):

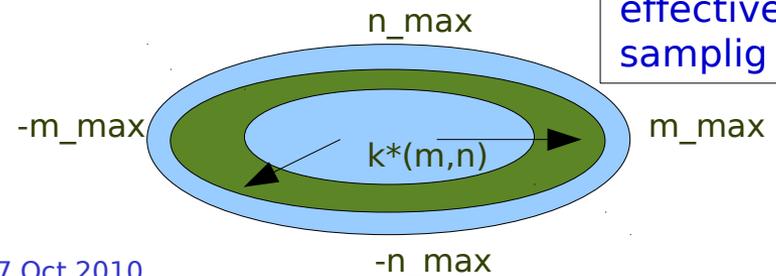
$$PECA(k^*, z) = \frac{1}{2\pi} \int_0^{2\pi} \frac{\epsilon_b^{m,n,z}}{\epsilon_{ref}^{m,n,z}} d\theta$$

$$\theta = \tan^{-1} \frac{n/nyl}{m/nxl}$$

On which scales our ensemble is effective for sampling B?

where

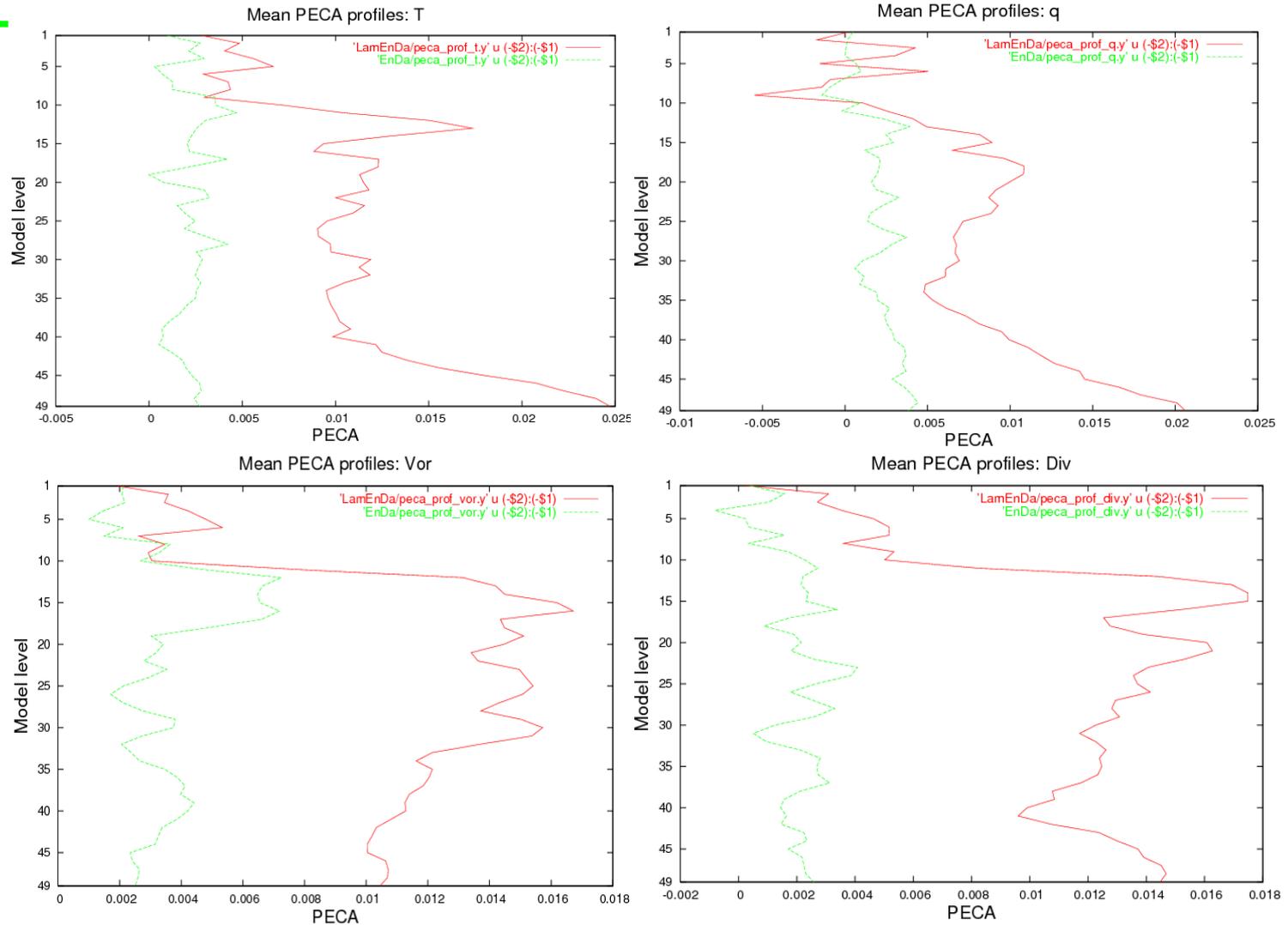
$$Ns \sqrt{\left(\frac{m}{nxl}\right)^2 + \left(\frac{n}{nyl}\right)^2} = k^*$$



# Diagnosis of DscEnDa vs. LamEnDa for sampling the B matrix

Larger PECA for LamEnDa:

LamEnDa sampling explains more the error background error variances than DscEnDa

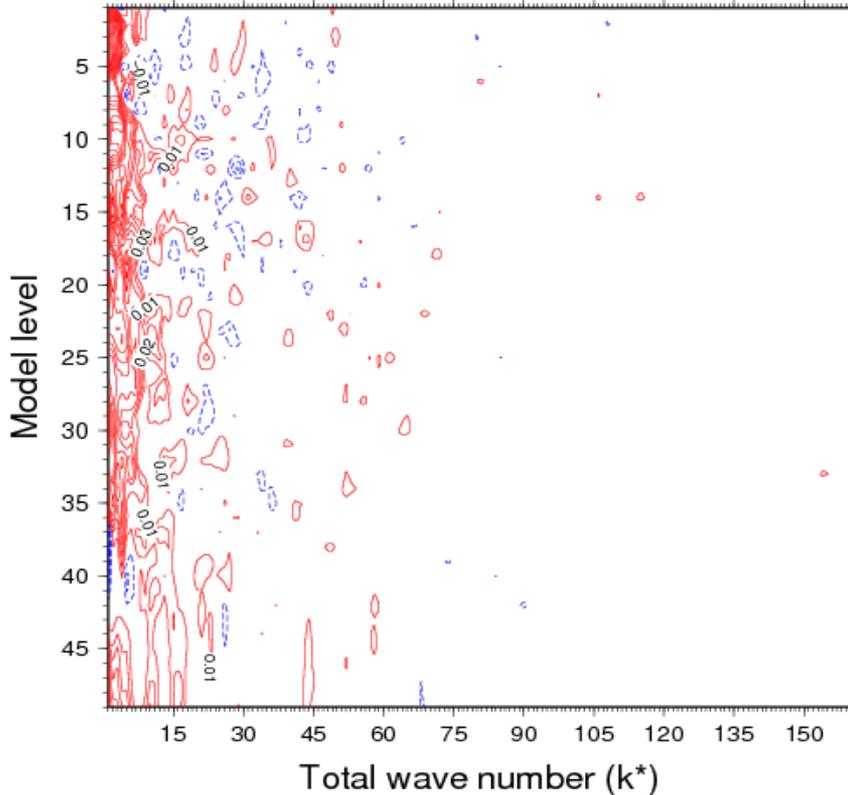


# Diagnosis of DscEnDa vs. LamEnDa for sampling the B matrix

Larger PECA for the small scales for LamEnDa: small scale errors are better sampled by LamEnDa than by DscEnDa

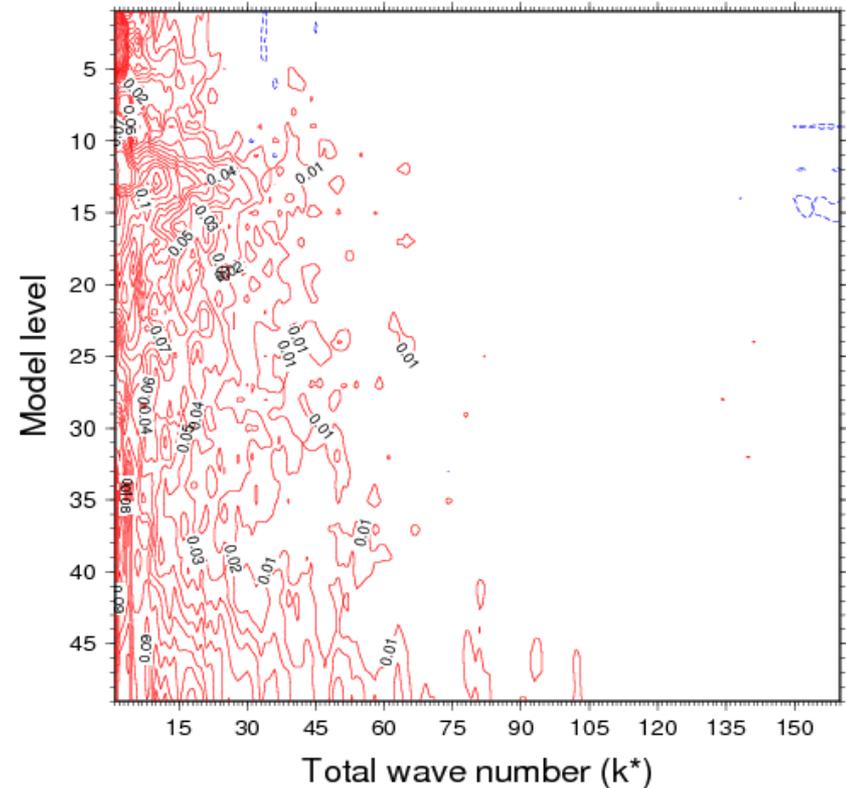
file: peca\_cort.xy

scaling : -1, max: 0.164590254, min: -0.0546603538, contours: 0.0



file: peca\_cort.xy

scaling : -1, max: 0.17724213, min: -0.015601974, contours: 0.01



# Time variations of background error structures

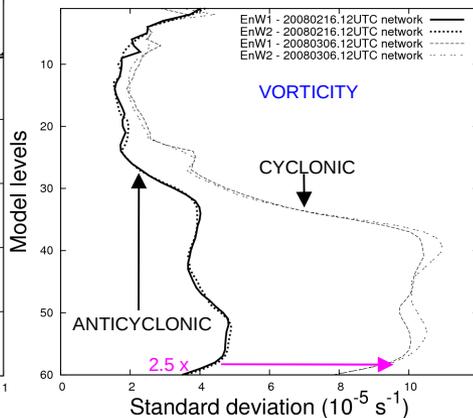
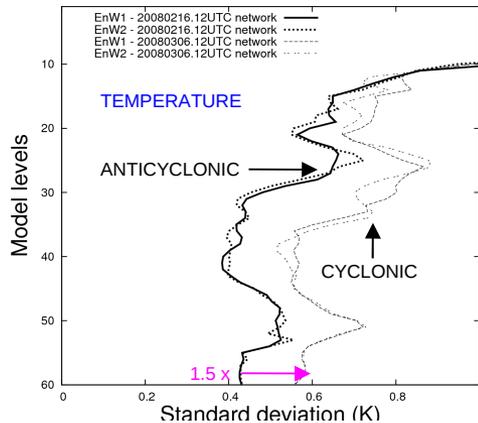
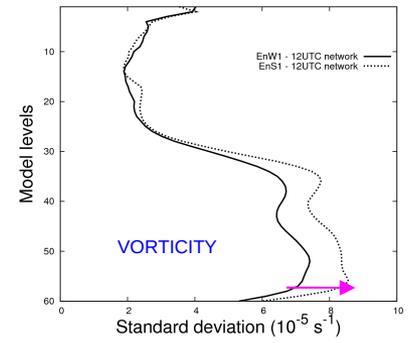
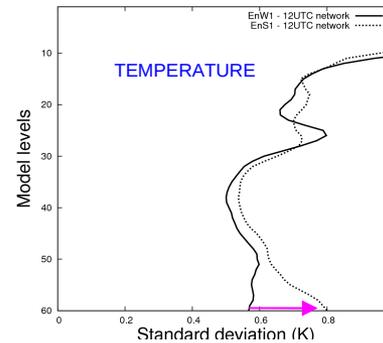
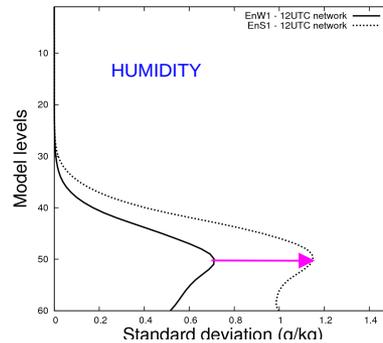
(Monteiro, M. and Berre, L., JGR Atmospheres 2010)

Background errors estimated with 6 perturbed members of 3D-var ALADIN/France ensemble DA system (RenDA)

**Season:** winter → summer

**Variations:** Variances ↑ Length-scales ↓

**Weather association:** convective activity in summer over ALADIN/France domain



**Day-to-day:** anticyclonic → cyclonic

**Variations:** Variances ↑ Length-scales ↓

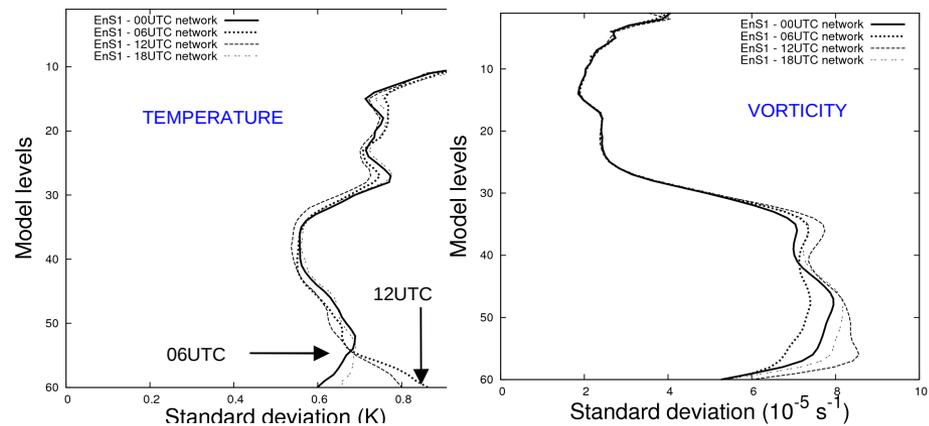
**Weather association:** instabilities in cyclonic situations

# Time variations of background error structures

**Hour :**

**Weather association :**

related to diurnal cycle mainly during summer in accordance with expected effects of afternoon convective activity



## basic conclusions:

- weather consistent variations of background error statistical structures
- robustness of the sampling method from “small”-size ensemble checked by using 2 independent ensembles with similar results
- sampling from “small”-size ensembles makes possible to compute flow dependent background error structure functions

# Observation, background and model error diagnosis in AROME

- model error  $\mathbf{Q}$  can be seen as part of  $\mathbf{B}$  ( $\mathbf{Q}=\alpha\mathbf{B}$ )
- indirect estimate: compare  $\mathbf{B}$  (in a “perfect model” framework, e.g. computed by ensemble method) to  $\mathbf{B}$  from a posteriori diagnostics
- Method I: Desroziers et. al. (2005):

$$E[(\mathbf{y} - \mathbf{x}_a)(\mathbf{y} - H(\mathbf{x}_b))^T] = \mathbf{R}$$

$$E[(\mathbf{x}_a - H(\mathbf{x}_b))(\mathbf{x}_a - H(\mathbf{x}_b))^T] = \mathbf{HBH}^T$$

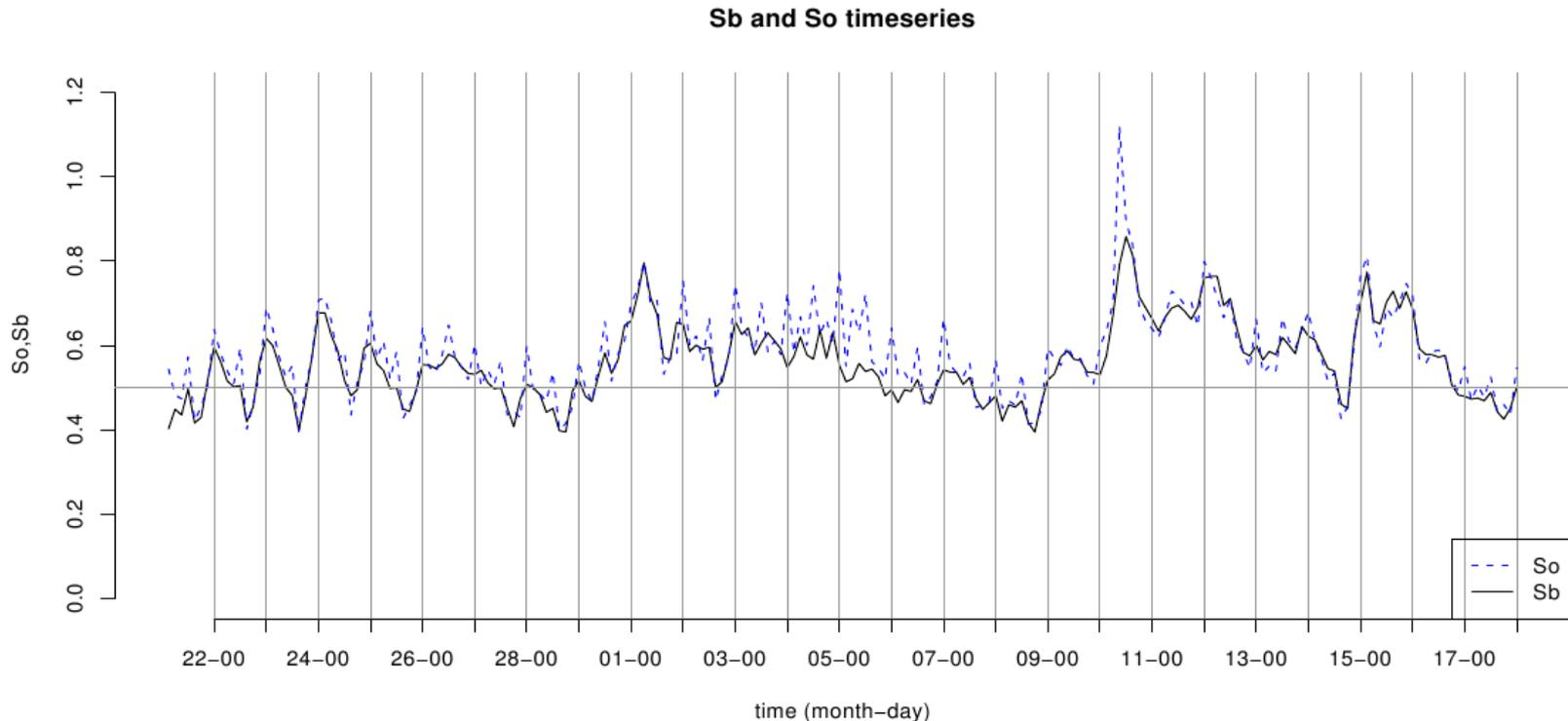
- Method II: Desroziers and Ivanov (2001), Berre et. al. (2006):
  - $\text{tr}(\mathbf{HK})$  estimate using ensemble of assimilations

$$S_o = \frac{2J_o(\mathbf{x}_a)]}{\text{Tr}(I_{p*p} - \mathbf{HK})}$$

$$S_b = \frac{2J_b(\mathbf{x}_a)]}{\text{Tr}(\mathbf{KH})}$$

# Observation, background and model error diagnosis in AROME

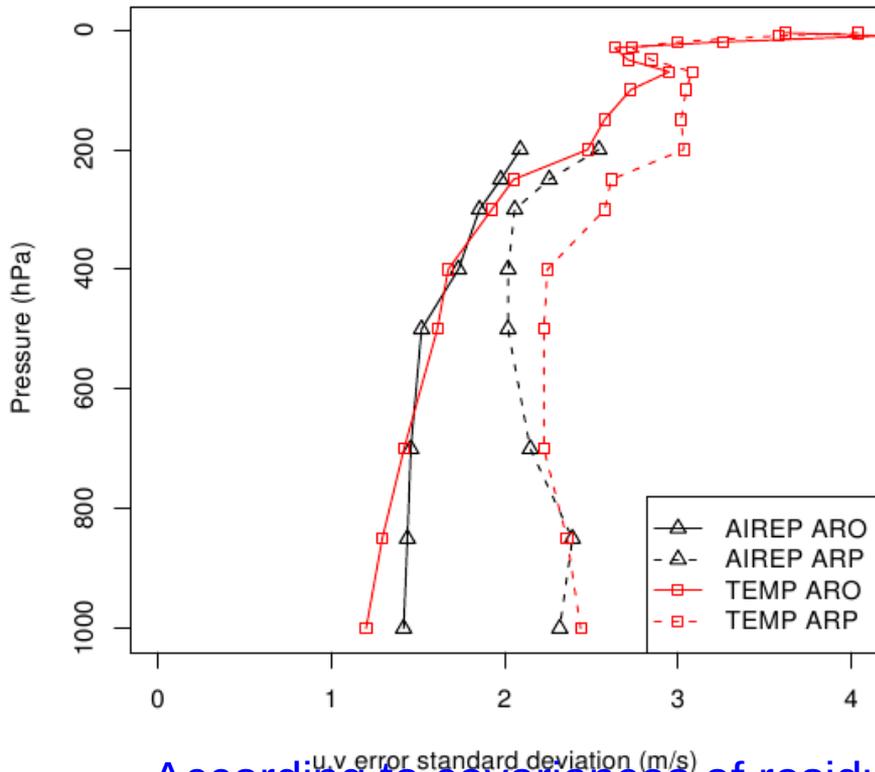
- Used methods give different results for  $\sigma_b$  and similar for  $\sigma_o$
- For method I, diagnosed  $\sigma_o$  is highly correlated with  $\sigma_b$  – this can indicate limitations of the method in high resolution
- Conclusion:  $\alpha = 1.54$  (method I) ,  $\alpha = 0.64$  (method II)



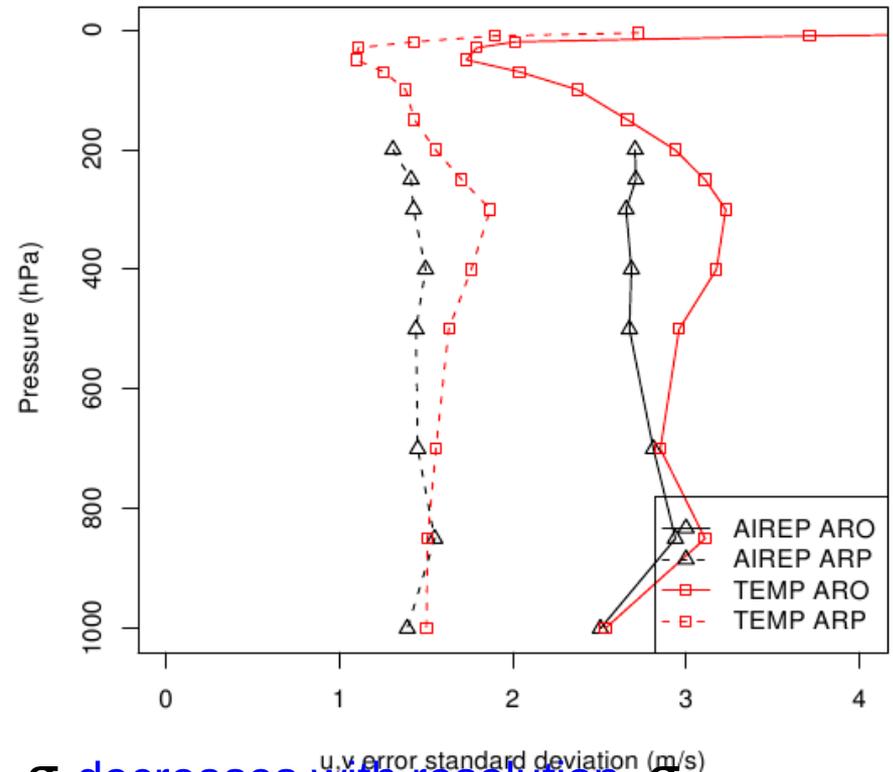
# Observation, background and model error diagnosis in AROME

## Resolution impact on $\sigma_o$ and $\sigma_b$

Profile of diagnosed observational error STD



Profile of diagnosed background error STD



- According to covariances of residuals,  $\sigma_o$  decreases with resolution,  $\sigma_b$  increases
- One should not use the same  $\sigma_o$ , e.g. in ARPEGE and AROME

# Local surface perturbation in ALADIN-HUNEPS

Upgrade of PEARP:

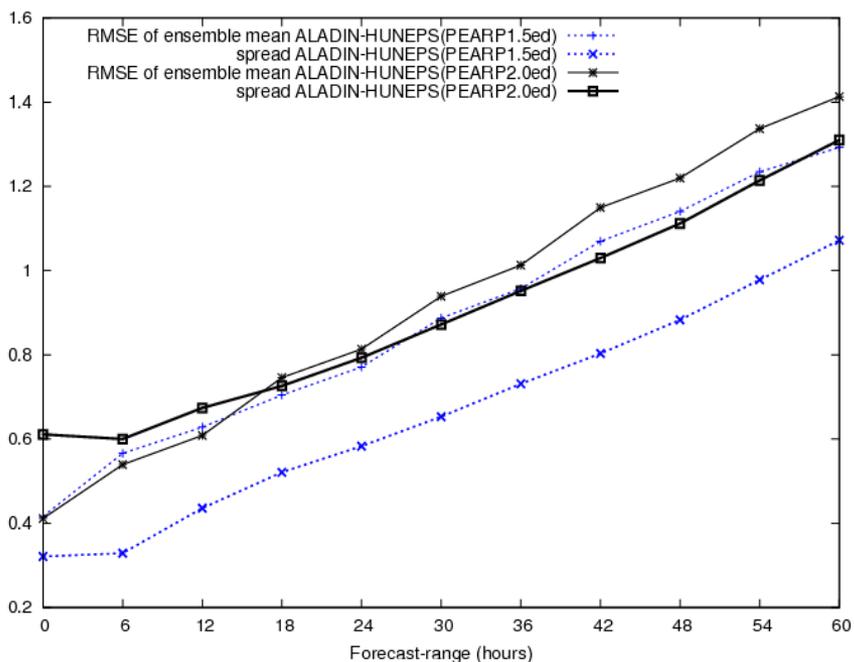
PEARP1.5: Targeted Singular Vectors

PEARP2: Targeted SV + perturbations from the ARPEGE Ensemble

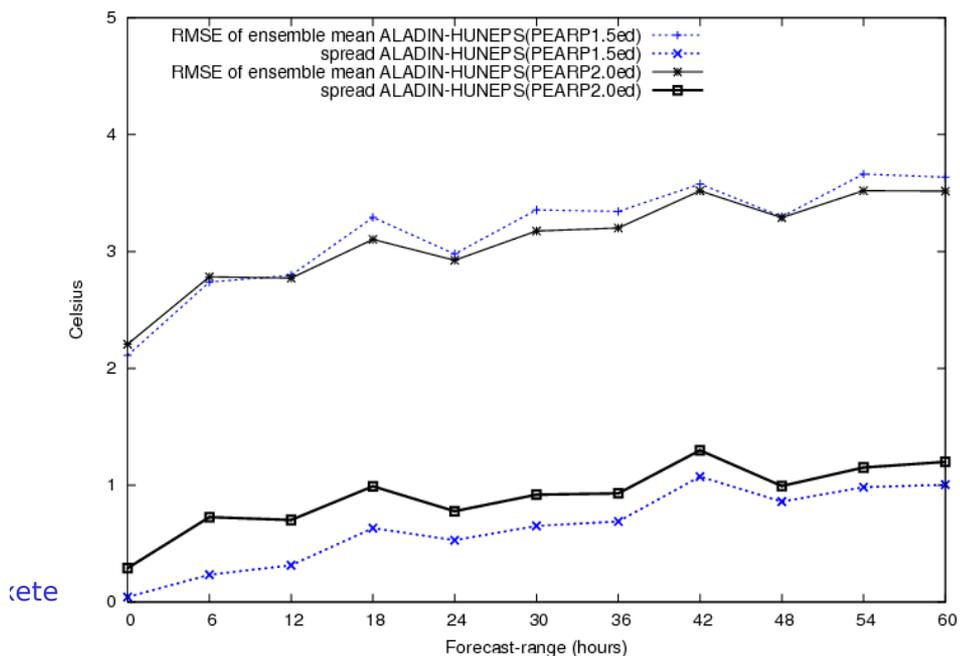
DA system (EnVar)

Improved spread-skill relationship  
but rather on high levels

T at 500 hPa

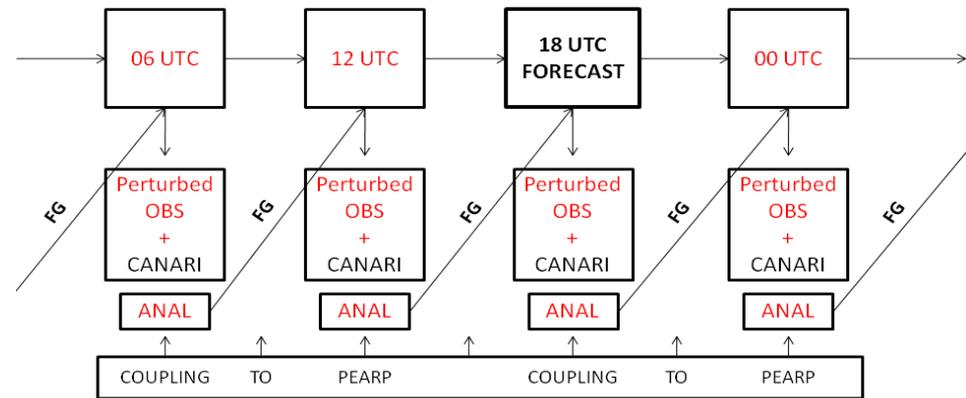


T 2m

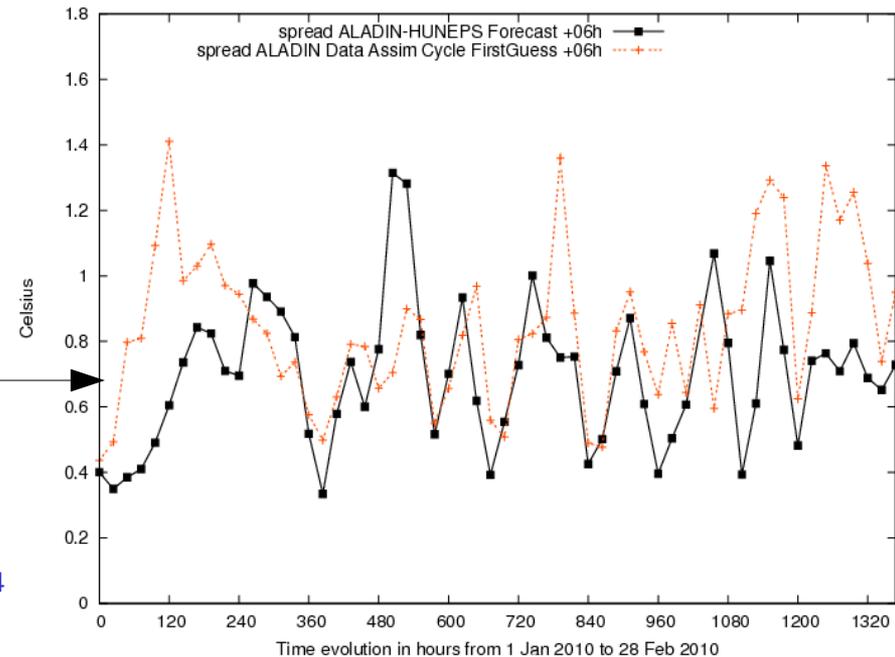


# Local surface perturbation in ALADIN-HUNEPS

Proposal: additional surface (soil) perturbations locally by a „surface ensemble DA” (Horányi et al., 2010, submitted to Tellus)



Somewhat increased spread at 2m



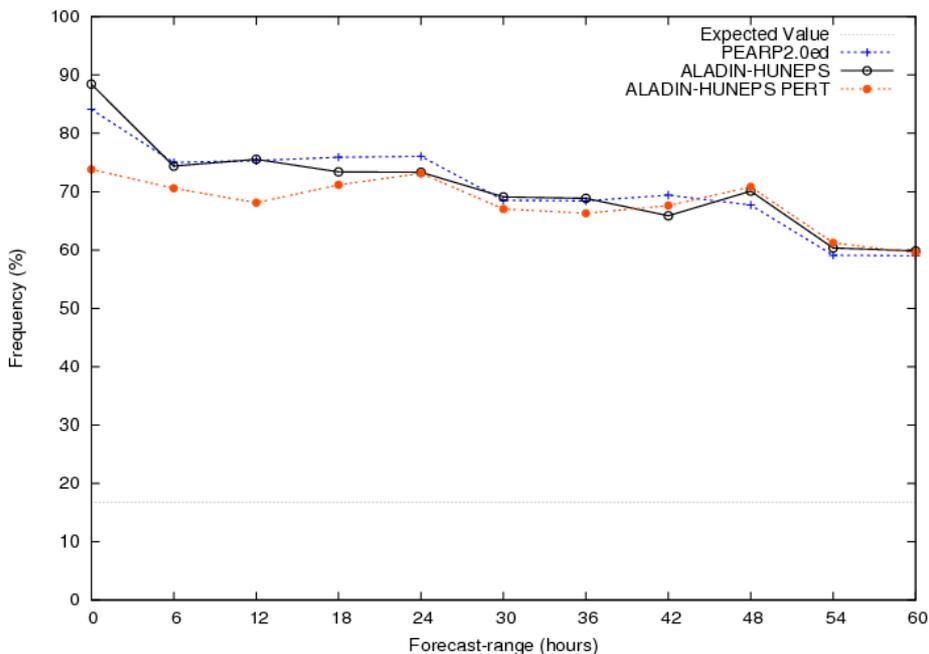
# Local surface perturbation in ALADIN-HUNEPS

PEARP2.0 downscaling + surface ensemble DA:

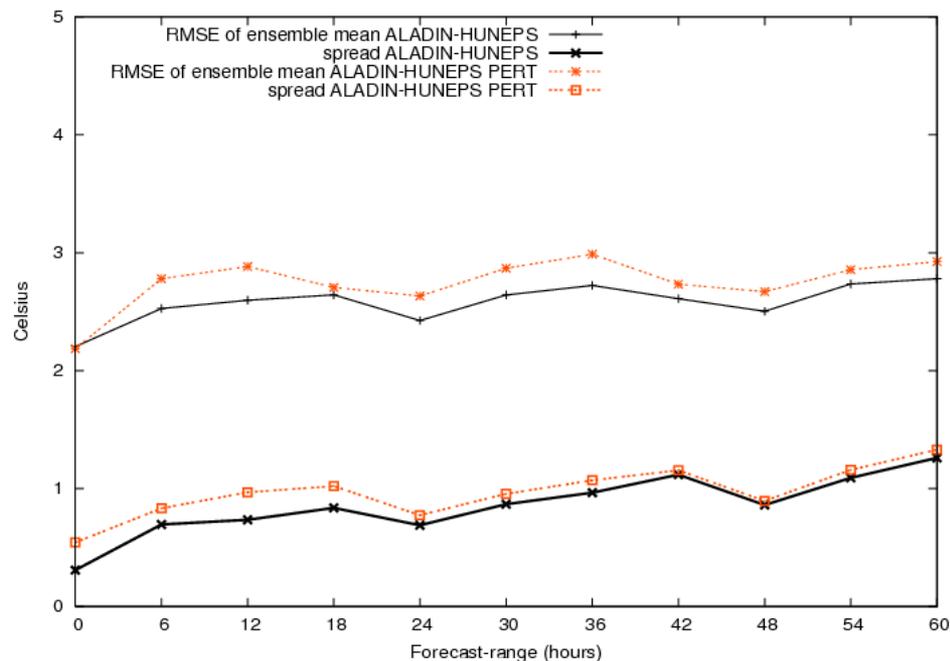
Percentage of outliers improved:

Spread-skill: spread increased but RMSE as well

T 2m



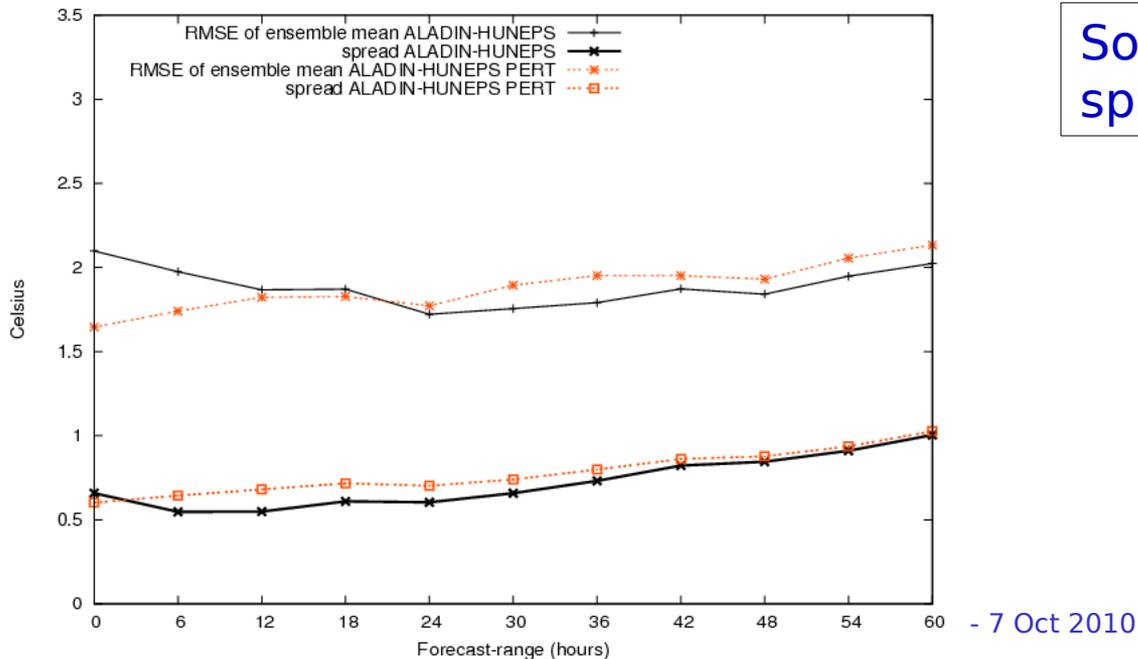
T 2m



# Local surface perturbation in ALADIN-HUNEPS

PEARP2.0 downscaling + surface ensemble DA:

1000 hPa

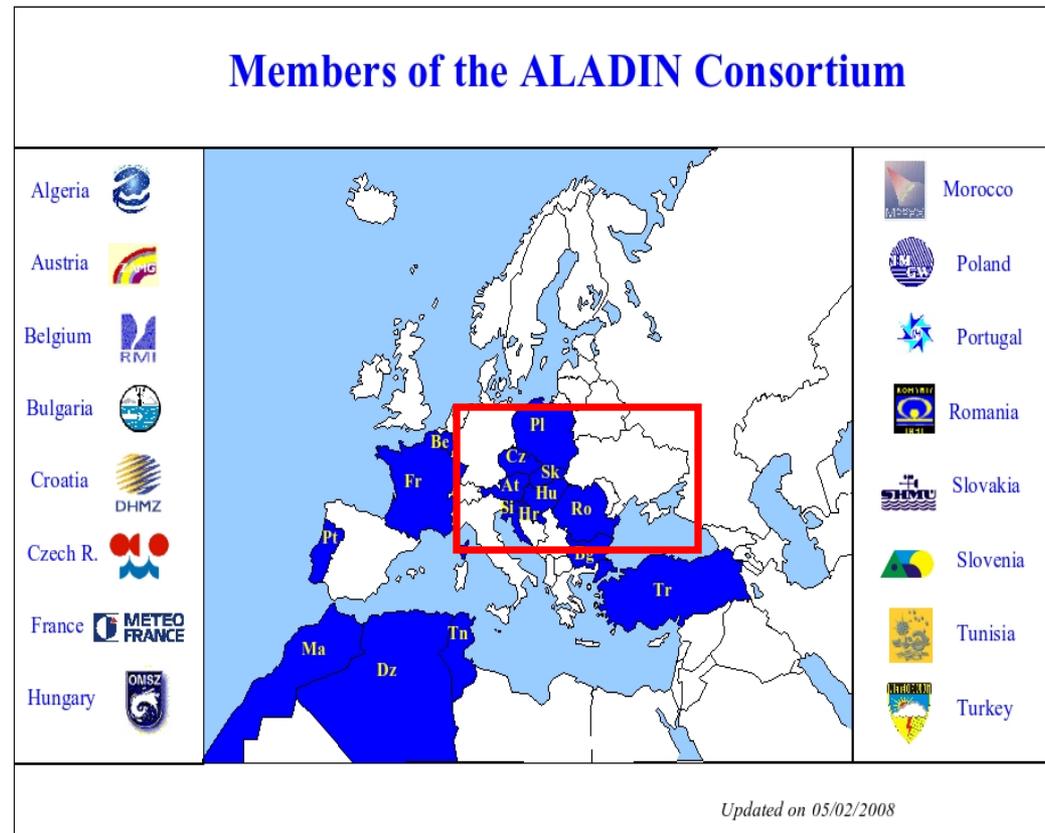


Some improvement in the spread-skill at 1000 hPa

Thank you for your attention

# LACE

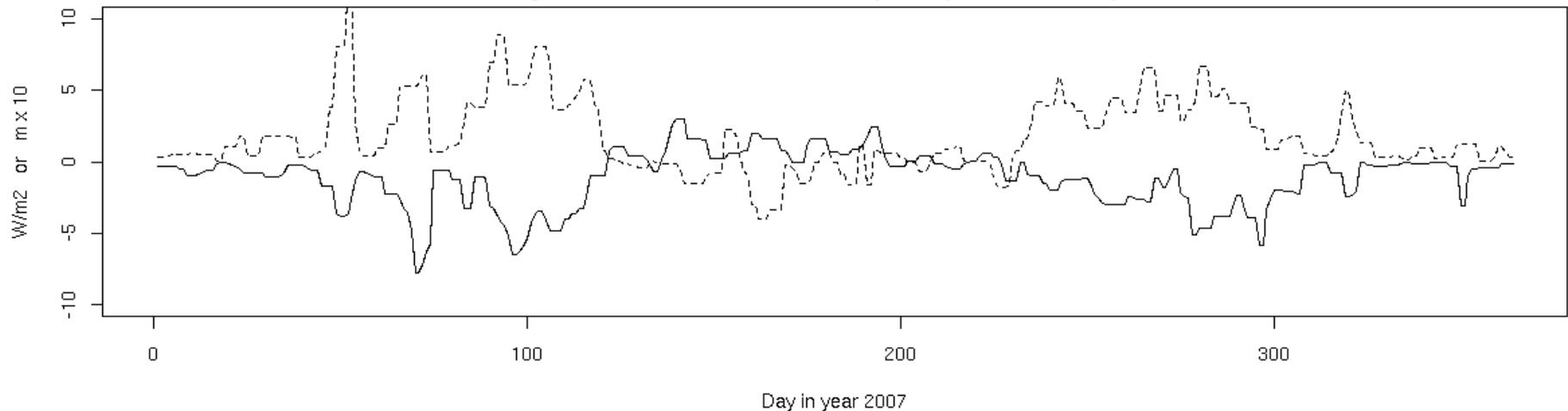
- Austria
- Croatia
- Czech Republic
- Hungary
- Romania
- Slovakia
- Slovenia



# Initialisation with LandSAF albedo

- Technical developments: use FA instead of previous binary formats (oper use is easier technically)
- 1 Year test: albedo assimilation acts as systematic bias correction of 2m fields
- LSAF albedo assim has an important effect on the PBL height

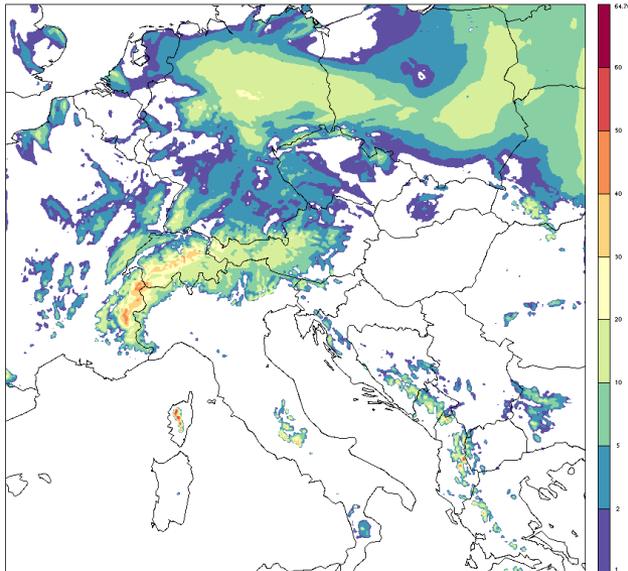
Hesse  
Daily average of sensible heat flux difference (exp-ref) (solid)  
and PBL height difference at 12 UTC, as analysed by the model (exp-ref) (dashed)



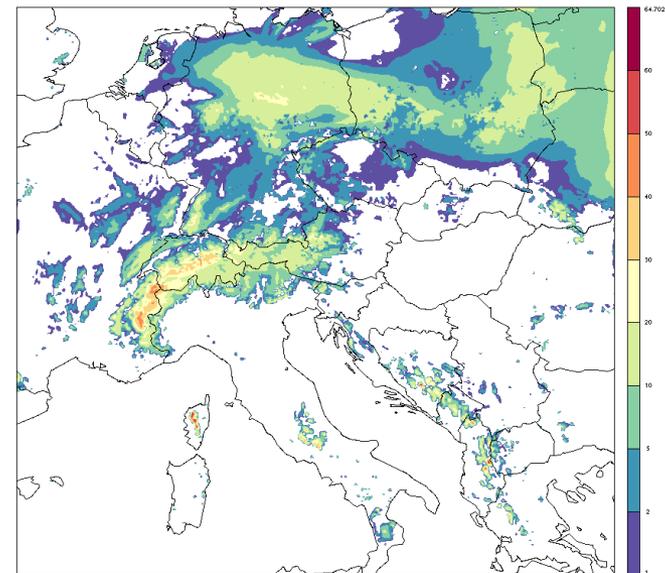
# Initialisation with LandSAF albedo

- Combine CANARI snow assimilation and LSAF product
- LSAF/MF-albedo-snow cover products gives an improved guess for CANARI snow analysis
- Use of snow height instead of snow water equivalent in CANARI
- Try to define evolving B errors for snow analysis (previous analysis error + increase where precipitation was falling)

Snow height [cm] (SWE/dens.), 4.4km SLOV domain  
canari snow height analysis, valid for 2010010206



Snow height [cm] (SWE/dens.), 4.4km SLOV domain  
canari snow analysis based on LSAF modified FG, valid for 2010010206



# Overview of data assimilation developments in LACE

Gergely Bölöni, Edit Adamcsek, Máté Mile, András Horányi, Mihály Szűcs, Alena Trojáková, Antonin Bucanek, Patrik Benacek, Xin Yan, Florian Meier, Tomislav Kovacic, Antonio Stanesic, Jure Cedilnik, Benedikt Strajnar, Mirela Niculae, Michal Nestiak, Maria Monteiro, Loïk Berre



EWGLAM/SRNWP, Exeter, 4 - 7 Oct 2010



# Outline

- Towards operational DA in LACE
- Diagnosis of Downscaled Global Ensemble DA and LAM Ensemble DA systems for sampling the B matrix
- Time variations of background error structures
- Observation, background and model error diagnosis in AROME
- Local surface perturbations in the LAMEPS system of Hungary

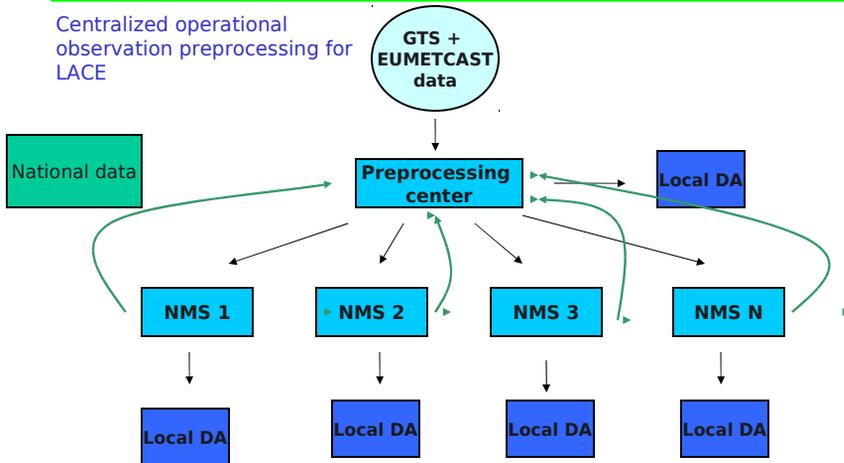
## Towards operational DA in LACE

- Regular DA runs in 5 countries: Austria, Croatia, Czech Rep., Hungary, Slovenia (Romania, Slovakia starts)
- Observations: OPLACE + local high-resolution SYNOP data
- Atm. 3DVAR + surf. OI (but wide range of cycling setups: IDFI, Blending of large scales, different LBCs)
- Improvements at 2m but not always above 850 hPa
- Room for improvement (LACE DA working days, 28-30 Sept., Ljubljana)
  - Order of the surf. (OI) and the atmospheric (VAR) analysis
  - Tune length scales in the surf. OI analysis
  - VARBC with daily cycling of the correction coefficients
  - Revise the use of AMSU channels
  - Jb/Jo tuning (more often, regularly)
  - Use mean orography instead of envelope

EWGLAM/SRNWP, Exeter, 4 - 7 Oct 2010

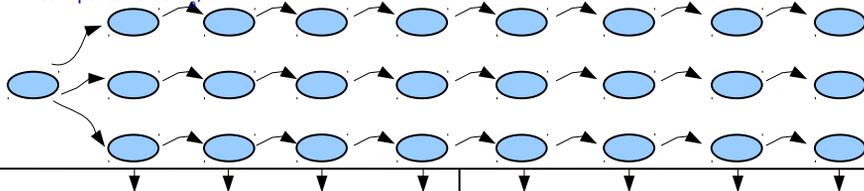
# Towards operational DA in LACE

Centralized operational observation preprocessing for LACE



# Downscaled Global Ensemble DA versus LAM Ensemble DA

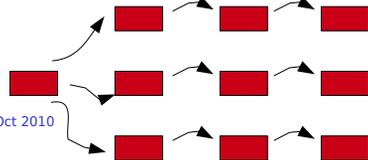
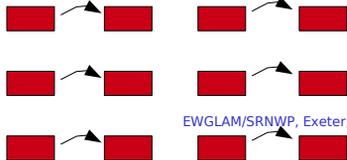
Global Ensemble Data Assimilation system (perturbed analyses through randomly perturbed observations with Gaussian distribution and rescaled amplitude comparable to  $\sigma$ )



Perturbation through LBCs

**DscEnDa:** „Downscaling the large scale uncertainty” kept in the LBCs

**LamEnDa:** adding direct local uncertainty by perturbed LAM analysis



EWGLAM/SRNWP, Exeter, 4 - 7 Oct 2010

# Diagnosis of DscEnDa vs. LamEnDa for sampling the B matrix

What is the information we gain with LamEnDa vs. DscEnDa?

Diagnosis with PECA: Perturbation versus Error Correlation analysis  
(Wei and Toth, MWR, 2003)

$$PECA = (\epsilon_b, \epsilon_{ref})$$

$$\begin{aligned} \epsilon_b &= X_b^{pert} - X_b^{ct} \\ \epsilon_{ref} &= X_b^{ct} - X_a^{verif} \end{aligned}$$

In what extent our ensemble differences explain the B error variance?

PECA defined in the ALADIN spectral space (after B cov computation by

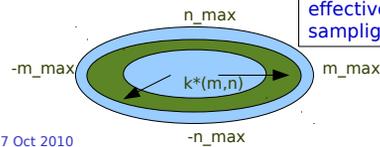
Berre, 2000):

$$PECA(k^*, z) = \frac{1}{2\pi} \int_0^{2\pi} \epsilon_b^{m,n,z} \epsilon_{ref}^{m,n,z} d\theta \quad \theta = \tan^{-1} \frac{n/nyl}{m/nxl}$$

On which scales our ensemble is effective for samplig B?

where

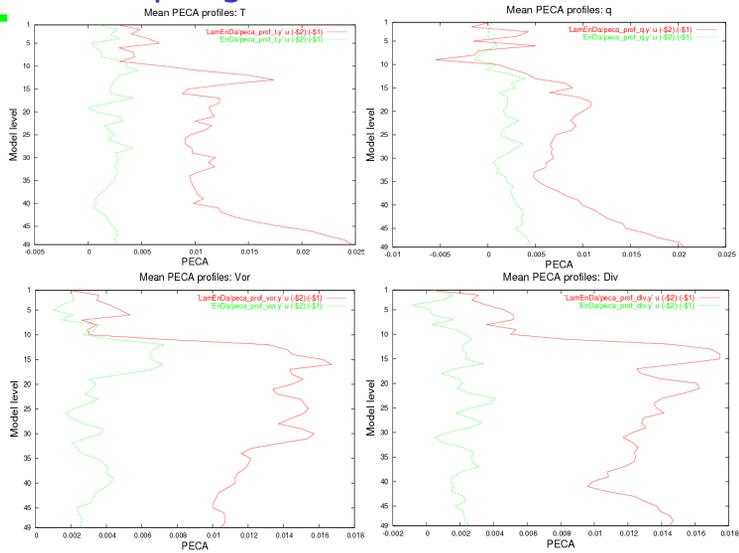
$$Ns \sqrt{\left(\frac{m}{nxl}\right)^2 + \left(\frac{n}{nyl}\right)^2} = k^*$$



# Diagnosis of DscEnDa vs. LamEnDa for sampling the B matrix

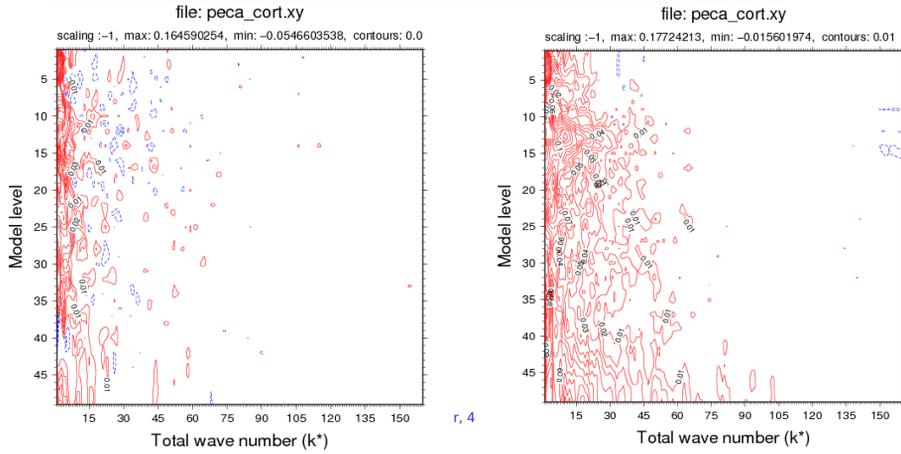
Larger PECA for LamEnDa:

LamEnDa sampling explains more the error background error variances than DscEnDa



# Diagnosis of DscEnDa vs. LamEnDa for sampling the B matrix

Larger PECA for the small scales for LamEnDa: small scale errors are better sampled by LamEnDa than by DscEnDa



# Time variations of background error structures

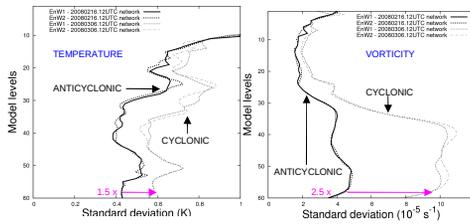
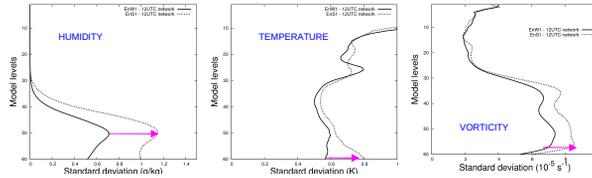
(Monteiro, M. and Berre, L., JGR Atmospheres 2010)

Background errors estimated with 6 perturbed members of 3D-var ALADIN/France ensemble DA system (RenDA)

**Season:** winter → summer

**Variations:** ↑ Length-scales ↓

**Weather association:**  
convective activity in summer  
over ALADIN/France domain



© 2010

**Day-to-day:** anticyclonic → cyclonic

**Variations:** ↑ Length-scales ↓

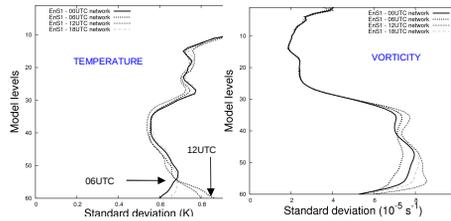
**Weather association:**  
instabilities in cyclonic situations

# Time variations of background error structures

**Hour :**

**Weather association :**

related to diurnal cycle mainly during summer in accordance with expected effects of afternoon convective activity



**basic conclusions:**

- weather consistent variations of background error statistical structures
- robustness of the sampling method from “small”-size ensemble checked by using 2 independent ensembles with similar results
- sampling from “small”-size ensembles makes possible to compute flow dependent background error structure functions

## Observation, background and model error diagnosis in AROME

- model error  $\mathbf{Q}$  can be seen as part of  $\mathbf{B}$  ( $\mathbf{Q}=\alpha\mathbf{B}$ )
- indirect estimate: compare  $\mathbf{B}$  (in a “perfect model” framework, e.g. computed by ensemble method) to  $\mathbf{B}$  from a posteriori diagnostics
- Method I: Desroziers et. al. (2005):

$$E[(\mathbf{y} - \mathbf{x}_a)(\mathbf{y} - H(\mathbf{x}_b))^T] = \mathbf{R}$$

$$E[(\mathbf{x}_a - H(\mathbf{x}_b))(\mathbf{x}_a - H(\mathbf{x}_b))^T] = \mathbf{HBH}^T$$

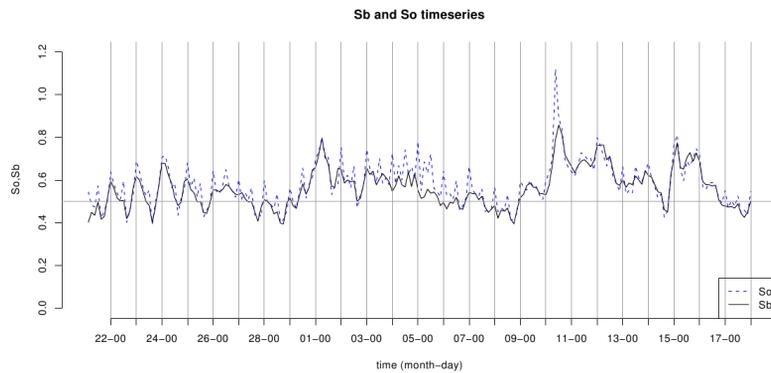
- Method II: Desroziers and Ivanov (2001), Berre et. al. (2006):
  - $\text{tr}(\mathbf{HK})$  estimate using ensemble of assimilations

$$S_o = \frac{2J_o(\mathbf{x}_a)}{\text{Tr}(I_{p \times p} - \mathbf{HK})}$$

$$S_b = \frac{2J_b(\mathbf{x}_a)}{\text{Tr}(\mathbf{KH})}$$

## Observation, background and model error diagnosis in AROME

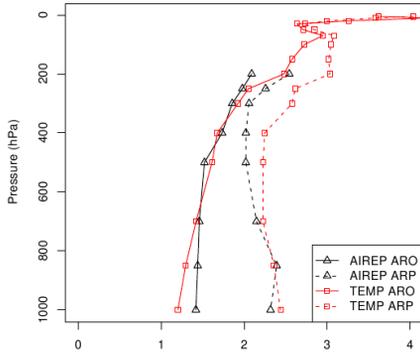
- Used methods give different results for  $\sigma_b$  and similar for  $\sigma_o$
- For method I, diagnosed  $\sigma_o$  is highly correlated with  $\sigma_b$  – this can indicate limitations of the method in high resolution
- Conclusion:  $\alpha = 1.54$  (method I) ,  $\alpha = 0.64$  (method II)



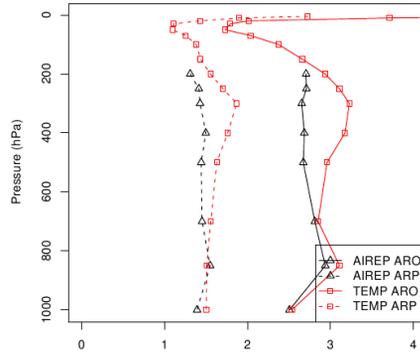
# Observation, background and model error diagnosis in AROME

## Resolution impact on $\sigma_o$ and $\sigma_b$

Profile of diagnosed observational error STD



Profile of diagnosed background error STD



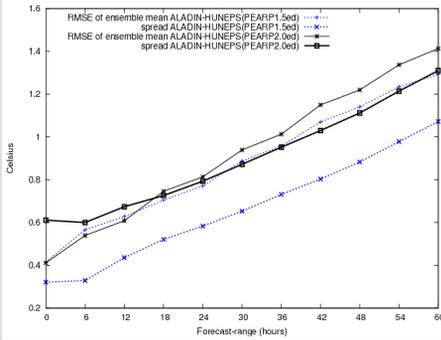
- According to covariances of residuals,  $\sigma_o$  decreases with resolution,  $\sigma_b$  increases
- One should not use the same  $\sigma_o$ , e.g. in ARPEGE and AROME

# Local surface perturbation in ALADIN-HUNEPS

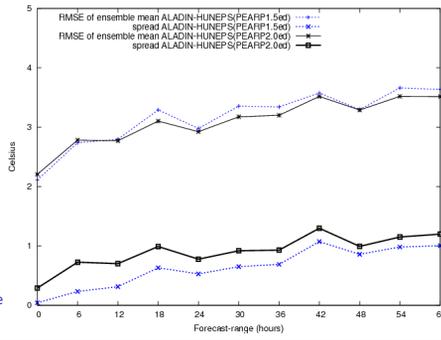
Upgrade of PEARP:  
 PEARP1.5: Targeted Singular Vectors  
 PEARP2: Targeted SV + perturbations from the ARPEGE Ensemble DA system (EnVar)

Improved spread-skill relationship but rather on high levels

T at 500 hPa

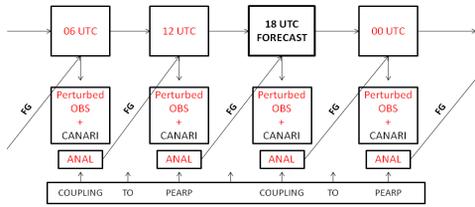


T 2m



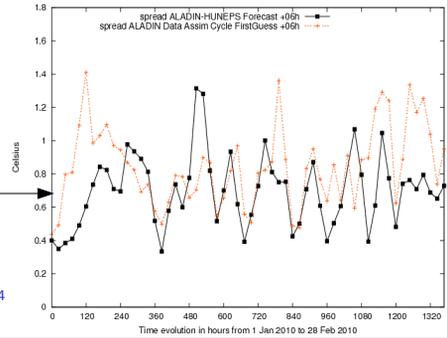
# Local surface perturbation in ALADIN-HUNEPS

Proposal: additional surface (soil) perturbations locally by a „surface ensemble DA” (Horányi et al., 2010, submitted to Tellus)



Somewhat increased spread at 2m

EWGLAM/SRNWP, Exeter, 4

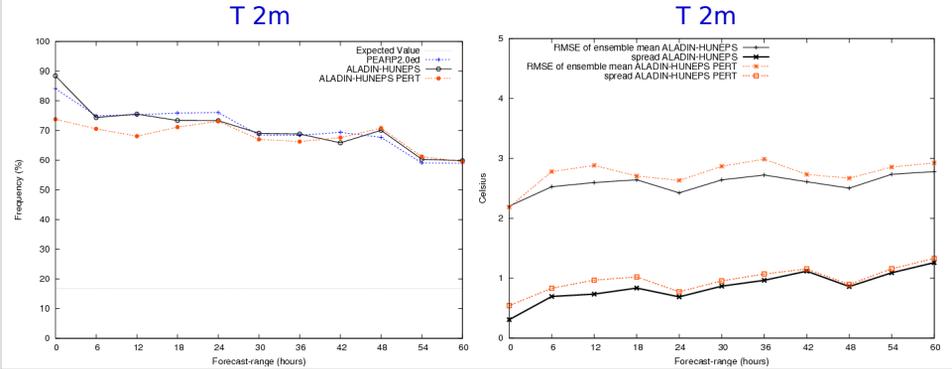


# Local surface perturbation in ALADIN-HUNEPS

PEARP2.0 downscaling + surface ensemble DA:

Percentage of outliers improved:

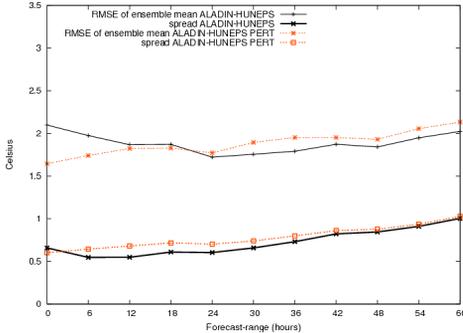
Spread-skill: spread increased but RMSE as well



# Local surface perturbation in ALADIN-HUNEPS

PEARP2.0 downscaling + surface ensemble DA:

1000 hPa



Some improvement in the spread-skill at 1000 hPa

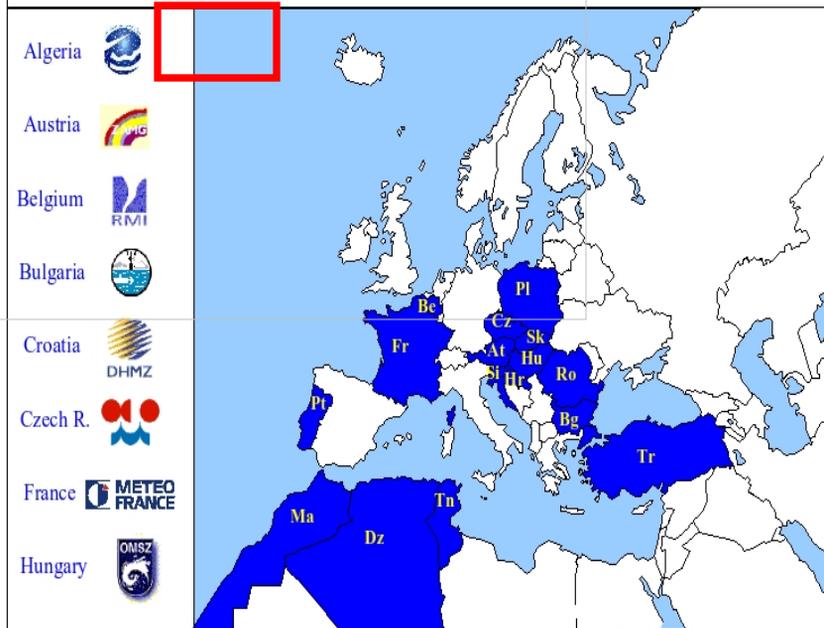
- 7 Oct 2010

Thank you for your attention

# LACE

- Austria
- Croatia
- Czech Republic
- Hungary
- Romania
- Slovakia
- Slovenia

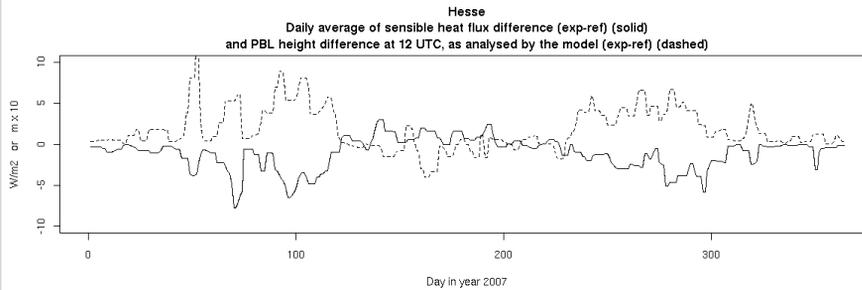
## Members of the ALADIN Consortium



Updated on 05/0

## Initialisation with LandSAF albedo

- Technical developments: use FA instead of previous binary formats (oper use is easier technically)
- 1 Year test: albedo assimilation acts as systematic bias correction of 2m fields
- LSAF albedo assim has an important effect on the PBL height



EWGLAM/SRNWP, Exeter, 4 - 7 Oct 2010

## Initialisation with LandSAF albedo

- Combine CANARI snow assimilation and LSAF product
- LSAF/MF-albedo-snow cover products gives an improved guess for CANARI snow analysis
- Use of snow height instead of snow water equivalent in CANARI
- Try to define evolving B errors for snow analysis (previous analysis error + increase where precipitation was falling)

