



# Evaluating the representation of clouds and precipitation in non-hydrostatic atmospheric models for the Belgian region

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## Integrations with ARPS

Integrations with COSMO2.8

Comparison with MODIS

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I. Meirold-Mautner, S. Crewell<sup>2</sup>, L. Delobbe<sup>3</sup>, G. Haase<sup>4</sup>*

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Introduction

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## QUEST (Crewell et al.)

Quantitative evaluation of regional precipitation forecasts using multi-dimensional remote sensing observations  
See also poster Thorsten Reinhardt et al

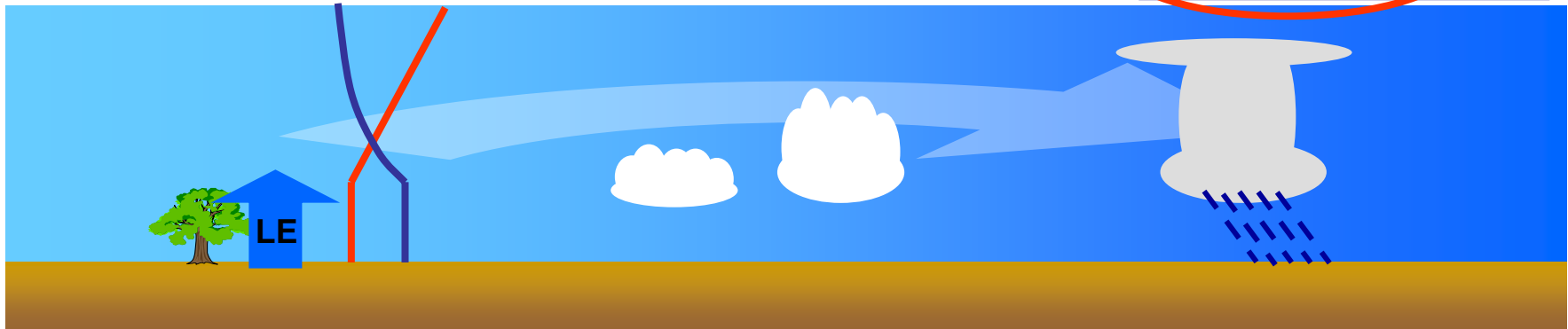
### satellite



MSG ~ 5km; 15min

- Cloud Mask
  - Cloud top pressure
- MODIS ~ 1km; 1day

- Cloud Mask
- Optical thickness

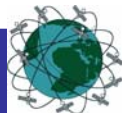


### IPT / Micro-wave



- 1D vertical;  
Lindenberg (and Cabauw)
- temperature profile
  - humidity profile
  - LWC

### GPS



- 147 stations;  
Germany;  
30min
- IWV

### Ceilometer



- 17 stations;  
Germany; 1min;  
ranges up to 4km
- Cloud base height
  - Cloud cover (<4km)

### Radar



- DX radar composite;  
1km; 5min
- Rain rate
- Polarimetric radar (DLR)
- Weather radars  
Wideumont

# Introduction

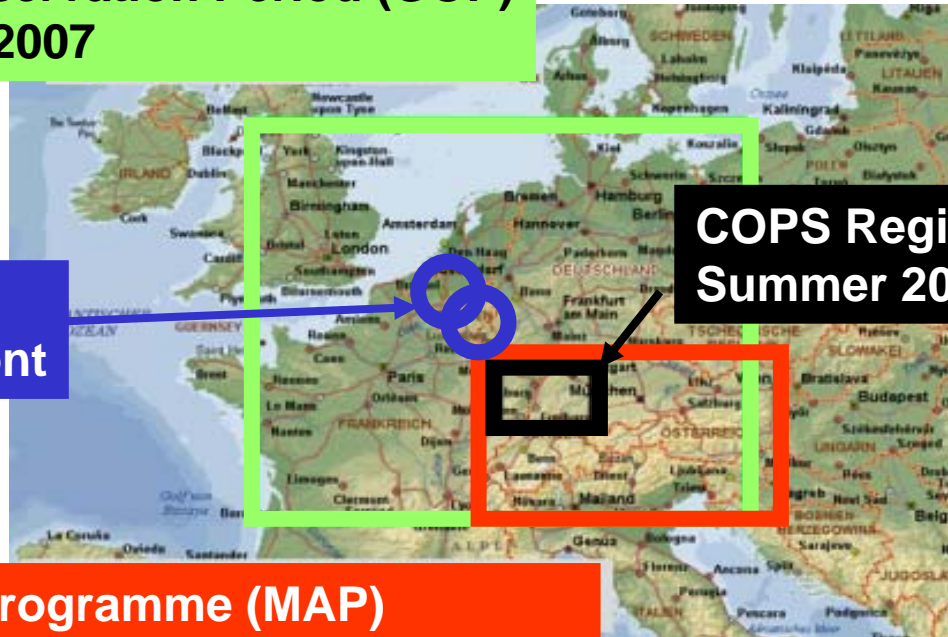
## Our region of interest

**General Observation Period (GOP)  
Entire year 2007**

**Belgian Radars  
Zaventem-Wideumont**

**COPS Region  
Summer 2007**

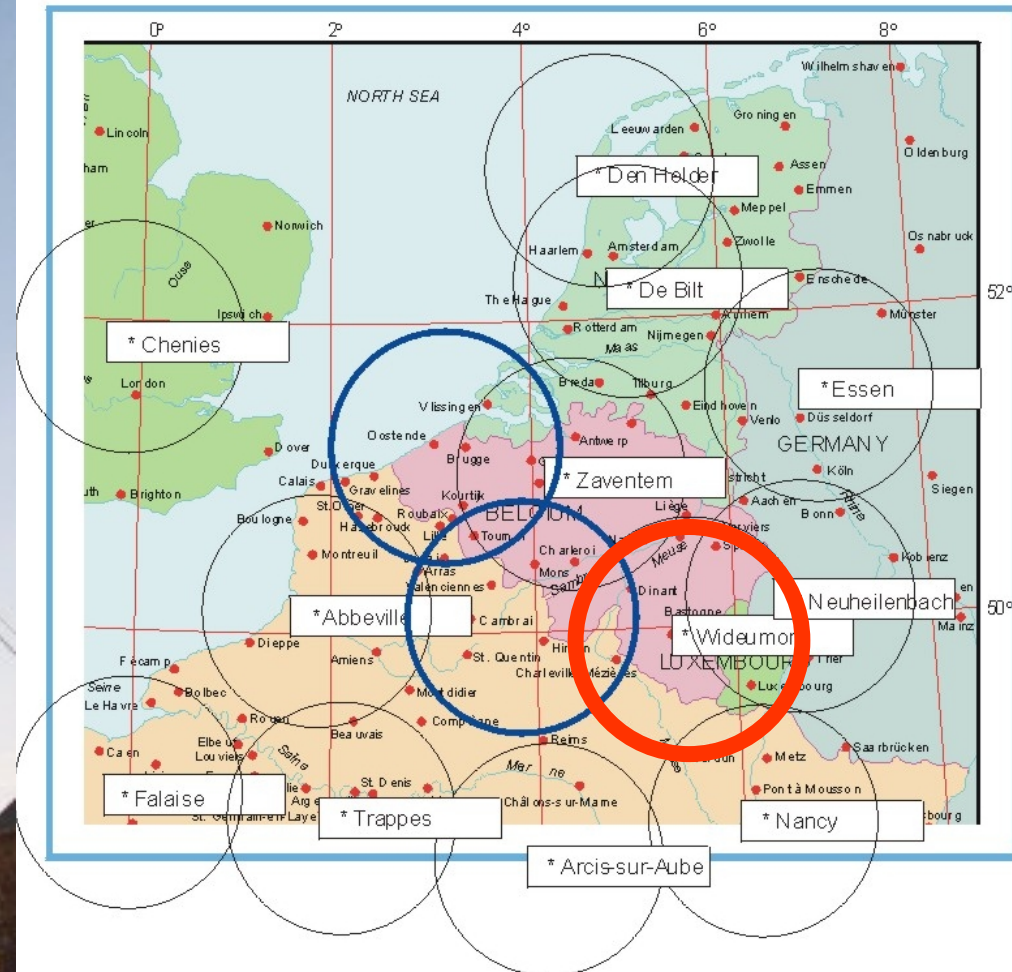
**Mesoscale Alpine Programme (MAP)  
Forecast Demonstration Project (FDP) region**



# Methods: Radar reflectivity

## C-band weather Radar Wideumont (RMI)

- Scans at 10 elevation angles each 15 minutes (0.5 – 17.5 °)
- Horizontal resolution is 500 m in range and 1 degree in azimuth



## Methods: Radar reflectivity

Based on Smith et al., 1975

Marshall-Palmer-type size distribution:  $N_k(D) = N_{0k} \exp(-\lambda D)$  (1)

$$Z_k = \int_0^{\infty} N_k(D) D^6 dD \quad (2)$$

$$\text{Radar reflectivity factor for rain} = \text{cnt}_1 \cdot \frac{(\rho q_r)^{7/4}}{N_r^{3/4} \rho_r^{7/4}} \quad (3)$$

$$\text{Radar reflectivity factor for dry snow} = \text{cnt}_2^* \cdot \frac{(\rho q_{s-})^{7/4}}{N_s^{3/4} \rho_i^{7/4}} \quad (4)$$

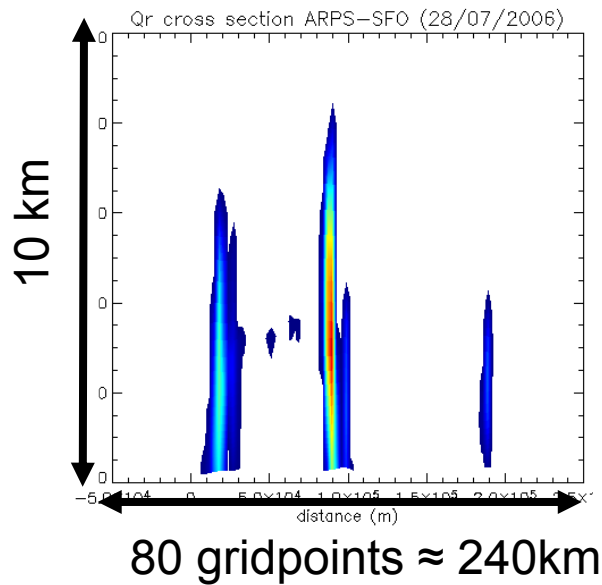
$$\text{Radar reflectivity factor for wet snow} = \text{cnt}_1 \cdot \frac{(\rho q_{s+})^{7/4}}{N_s^{3/4} \rho_s^{7/4}} \quad (5)$$

$$\text{Radar reflectivity factor for hail} = \text{cnt}_3 \cdot \left( \frac{(\rho q_h)^{7/4}}{N_h^{3/4} \rho_h^{7/4}} \right)^{0.95} \quad (6)$$

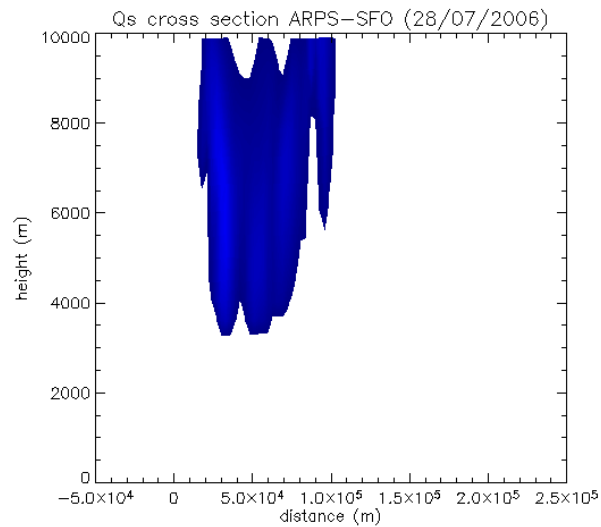
\* Taking into account the ratio of dielectric factors for water and ice

## An illustration: Precipitating hydrometeors in ARPS

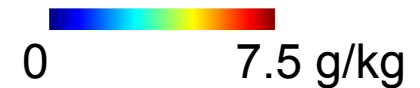
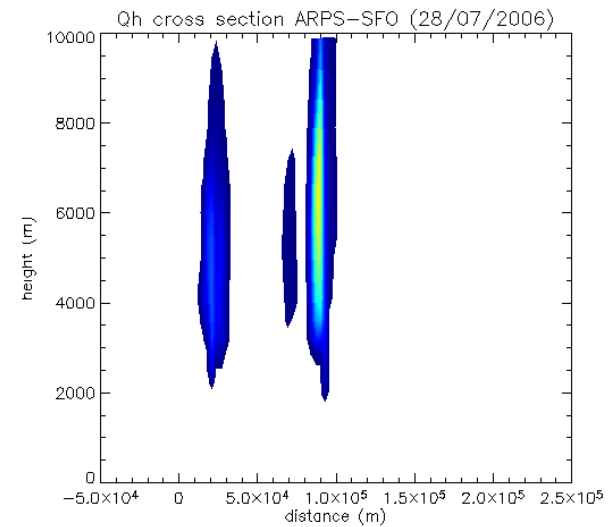
rain



snow

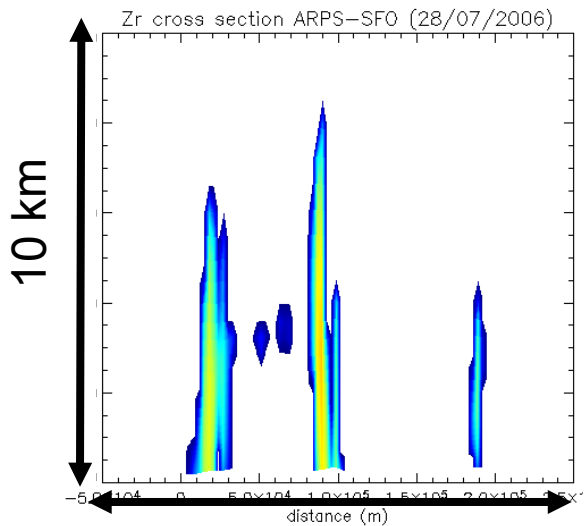


hail



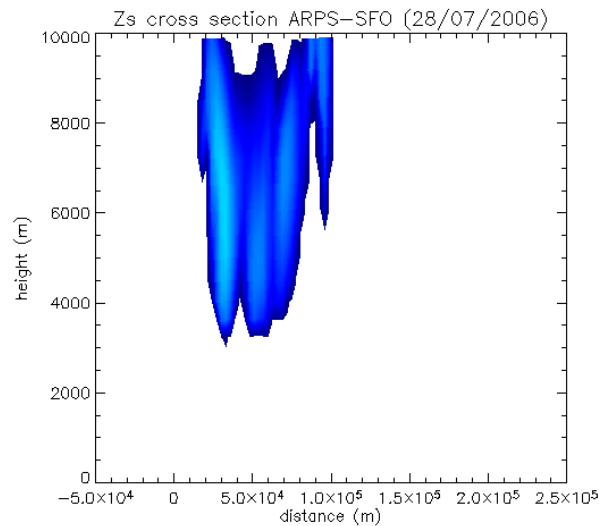
## An illustration: Radar reflectivities in ARPS

rain

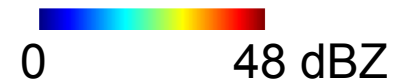
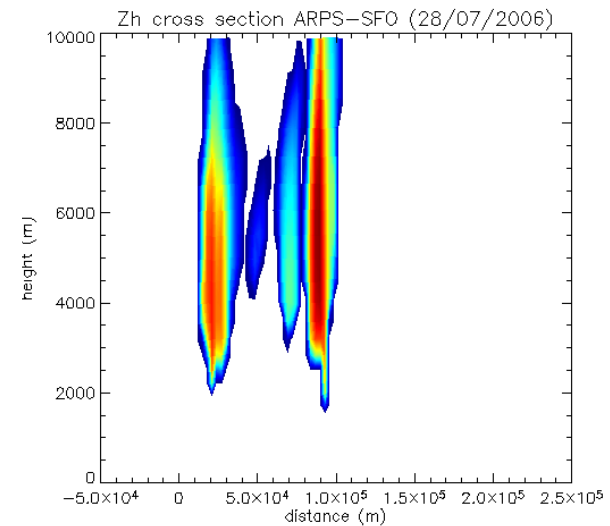


80 gridpoints  $\approx$  240km

snow



hail



$$Z = 10 \log_{10} \left( \frac{1 \text{ mm}^6 \text{ m}^{-3}}{\text{m}^3} \right)$$

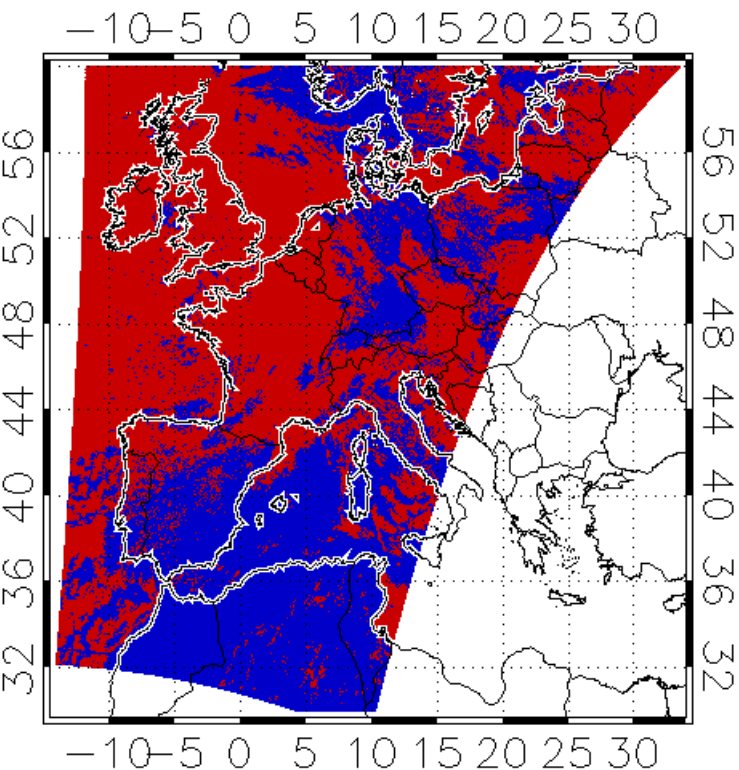
$$Z_e = Z_{er} + Z_{es} + Z_{eh}$$



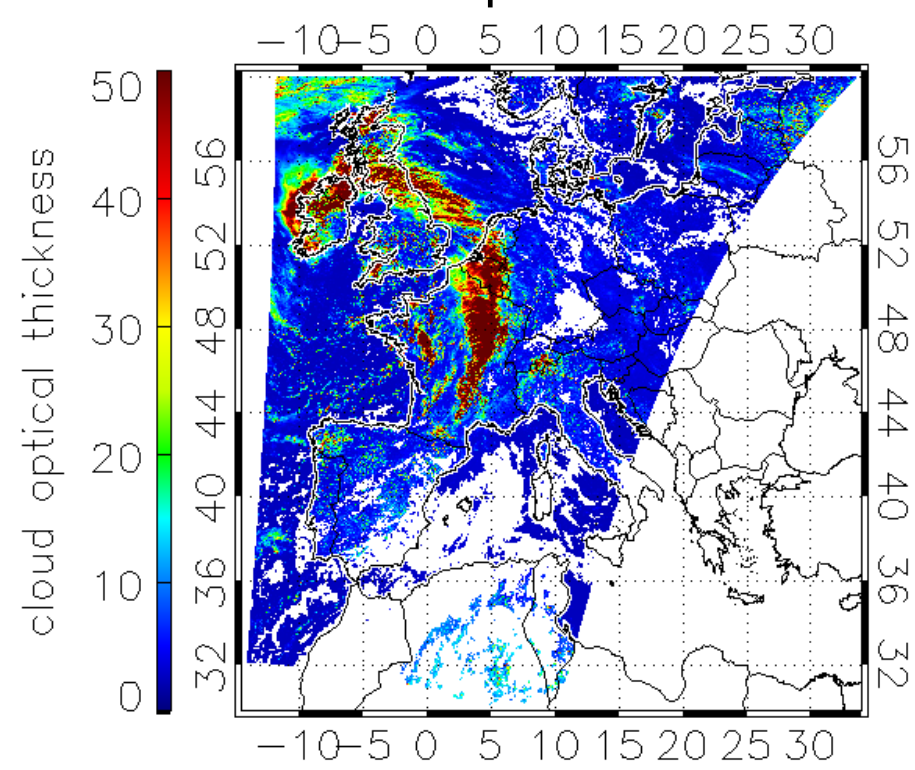
Developing a method to evaluate cloud optical thickness using satellite remote sensing (Van Lipzig et al., 2006; Schröder et al, 2006)

12 Aug 2004 MODIS overpass 10:55UTC

Cloud cover



Cloud optical thickness



# Methods: cloud optical thickness

Clouds in atmospheric models affect:

- 1: precipitation:  $q^c$ ,  $q^i$
- 2: source/sink of heat related to phase changes:  $q^c$ ,  $q^i$
- 3: radiation: optical properties

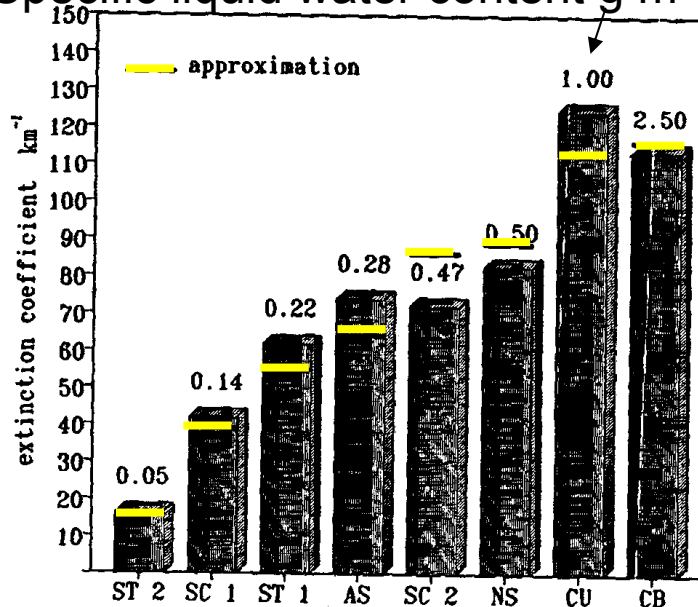
Radiation code LM (Ritter and Geleyn, 1992)

Radiative fluxes and heating rates

Cloud optical properties

Bulk model quantities

Specific liquid water content  $g\ m^{-3}$



Coefficients for 8 different cloud types

$$r_e = c_7 + c_8 \rho_{LW}$$
$$\delta = \left\{ c_1 + \frac{c_2}{r_e} \right\} \rho_{LW} \Delta z$$

# Methods: cloud optical thickness

Clouds in atmospheric models affect:

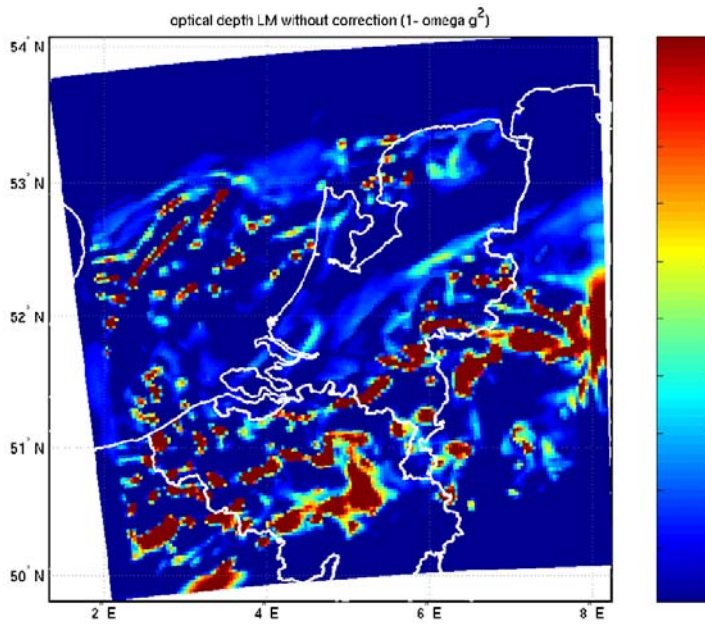
- 1: precipitation:  $q^c$ ,  $q^i$
- 2: source/sink of heat related to phase changes:  $q^c$ ,  $q^i$
- 3: radiation: optical properties

Radiation code other models (e.g. ECMWF)

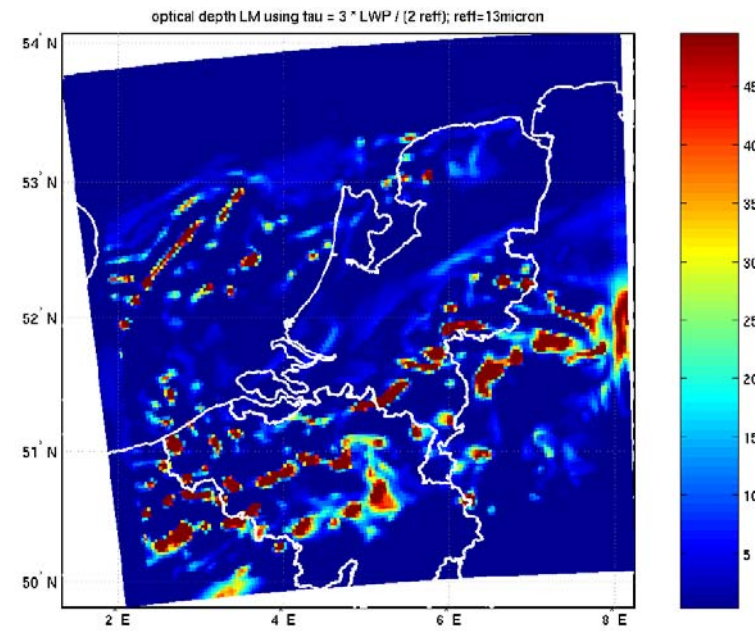
$$\delta = \frac{3\rho_{LW}\Delta z}{2r_e} \quad \text{With } r_e \text{ constant}$$

Formulation from COSMO-code

COSMO2.8  
Cloud  
optical  
thickness  
12 Aug 2004  
MODIS  
overpass  
10:55UTC

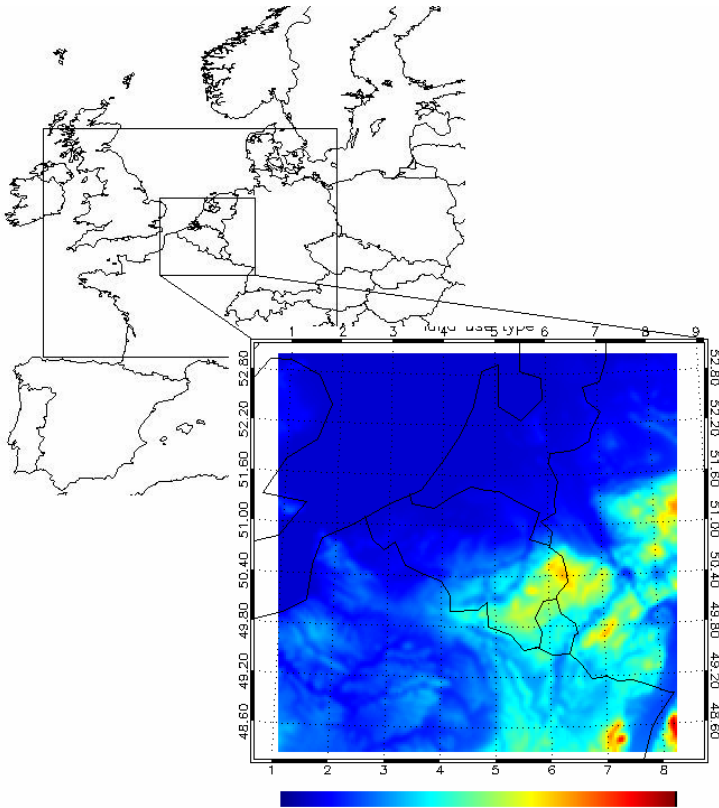


Approximation



**Advanced Regional Prediction System (ARPS):** Non-hydrostatic mesoscale model (Xue et al. 2000, 2001), developed at CAPS

**COSMO2.8:** Non-hydrostatic mesoscale model (Doms et al., 1999)



**ARPS** (integrations by Kwinten Van Weverberg):  
Double one-way nested grid with successive  $\Delta x$  of 9 km and 3 km. Smallest model domain centered over Belgium. Boundary and initial conditions derived from ECMWF operational analysis. 50 vertical levels

No convection parameterization in smallest domain, Kain-Fritsch convection parameterization larger domain

Lin-Tao microphysics (including rain, snow, hail, cloud ice, cloud water)

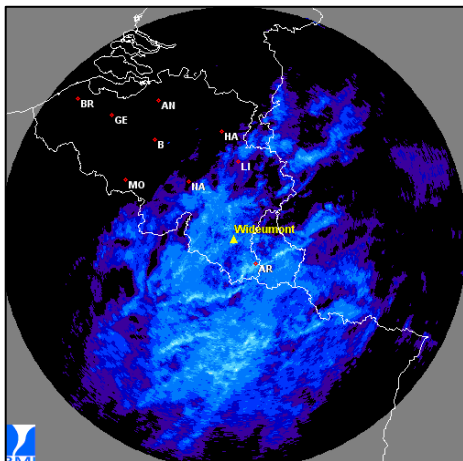
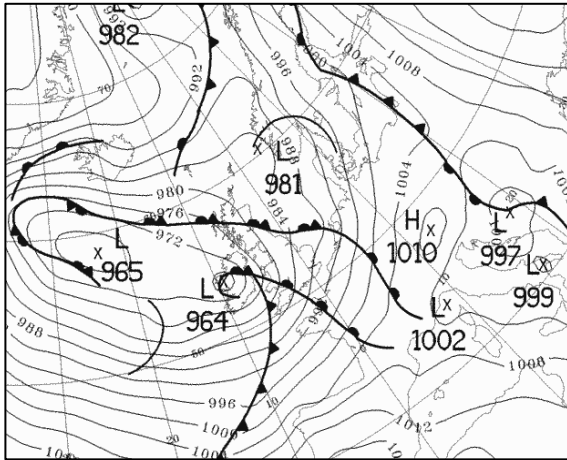
**COSMO2.8** (integrations by Ingo Meirold-Mautner):  
Version 3.21 integrations centered above Belgium,  $\Delta x = 2.8$  km, 160x160 grid points, 50 levels, prognostic variables for cloud, ice, rain, snow, graupel, driven by COSMO-LME analyses



# Description of the cases

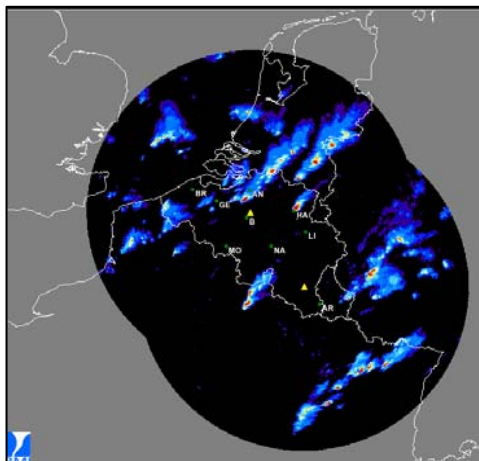
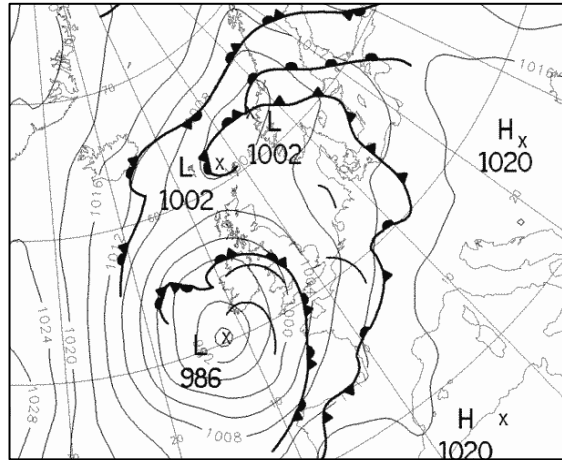
## Stratiform precipitation case

23/11/2006



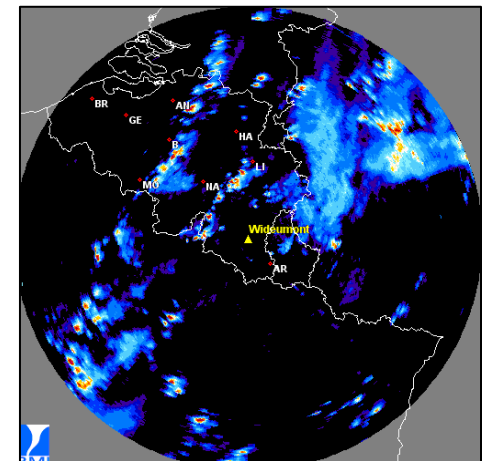
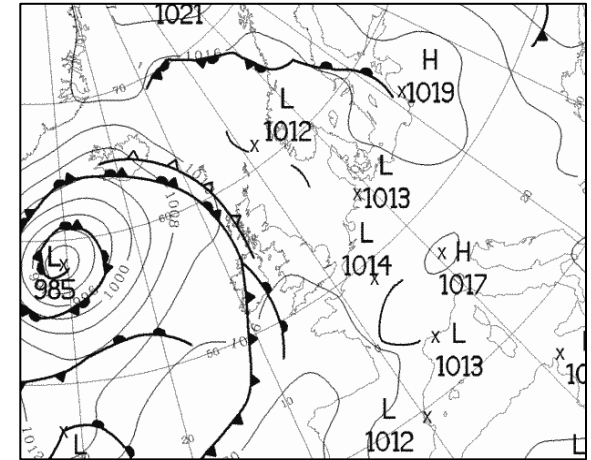
## Dynamically driven convective case

01/10/2006



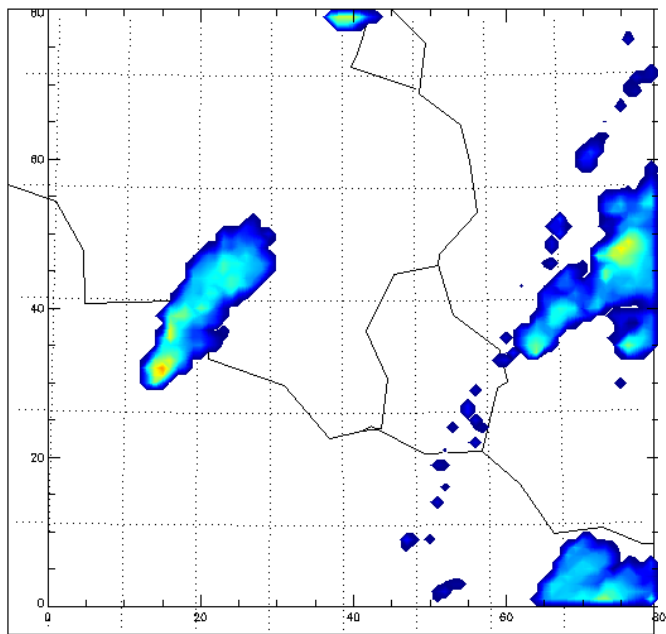
## Thermodynamically driven convective case

28/07/2006

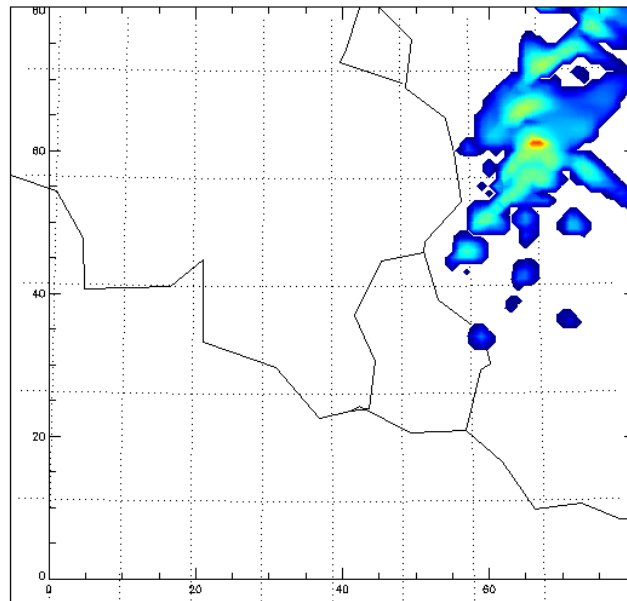


# Results: October case

## Radar 18:00



## ARPS 18:00



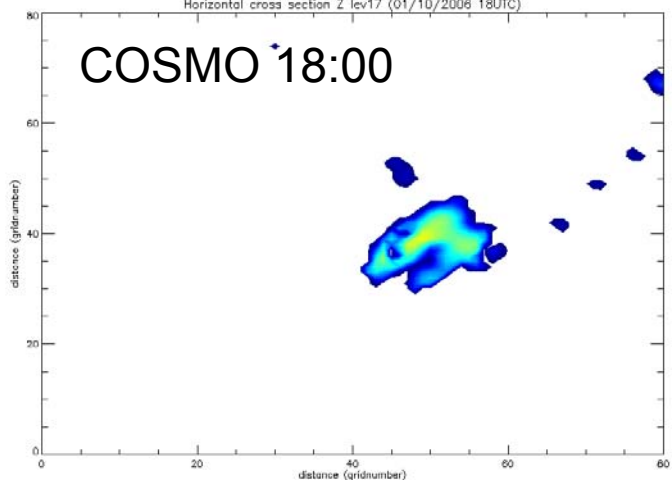
80 gridpoints  $\approx$  240km

80 gridpoints  $\approx$  240km

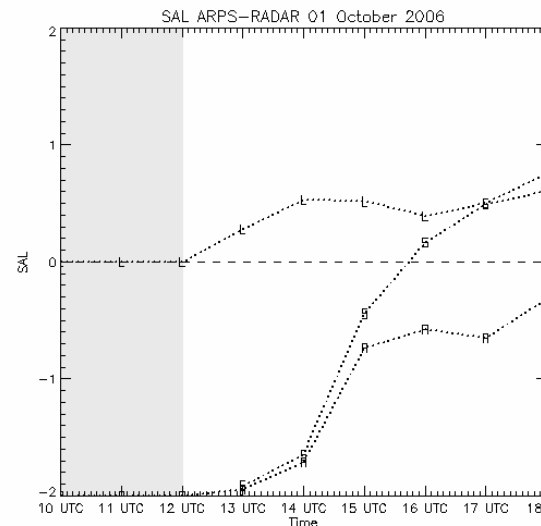
80 gridpoints  $\approx$  240km

### Horizontal cross section Z lev17 (01/10/2006 18UTC)

## COSMO 18:00

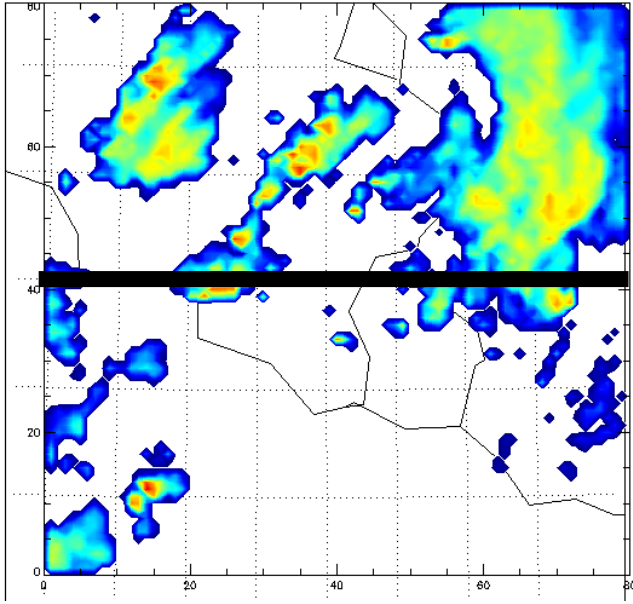


Reflectivities



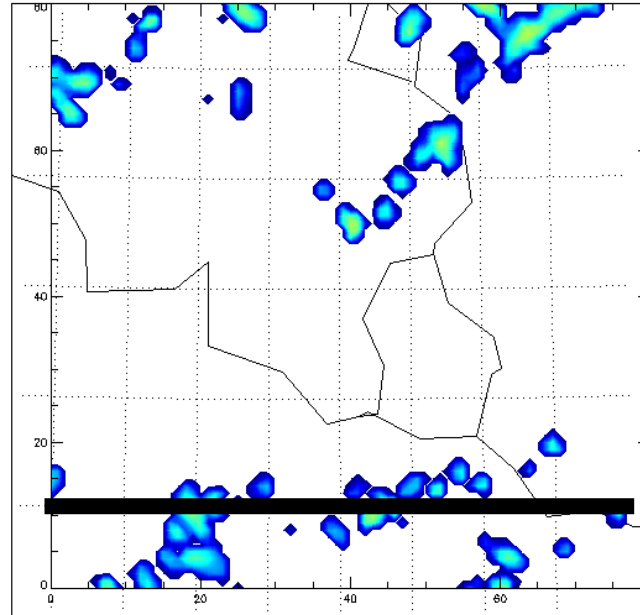
# Results: July case

## Radar 15:00



Horizontal cross section Z lev17 (28/07/2006 15UTC)

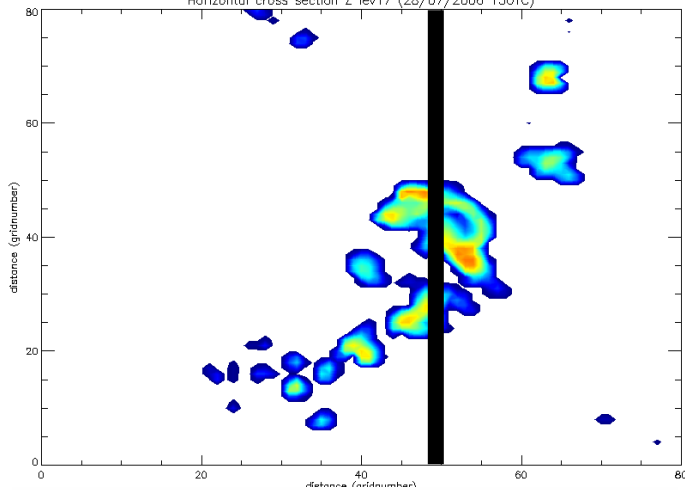
## ARPS 15:00



80 gridpoints  $\approx$  240km

80 gridpoints  $\approx$  240km

80 gridpoints  $\approx$  240km

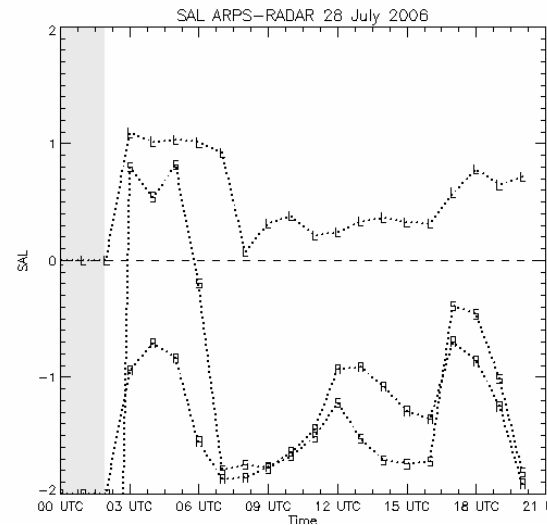


## COSMO 15:00

### Reflectivities

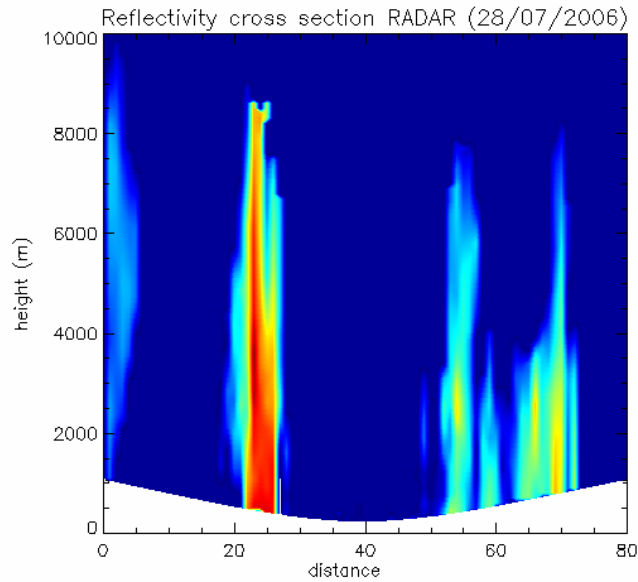


0 dBZ 56

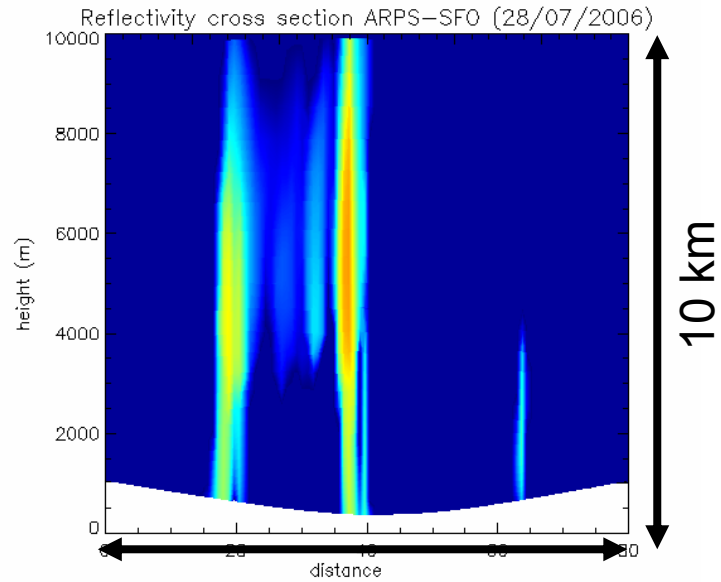


# Results: July case

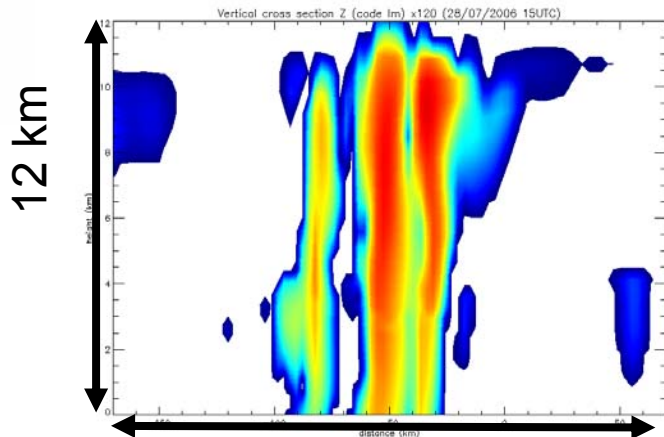
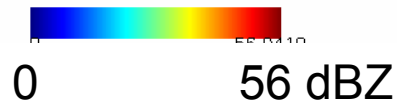
## Radar 15:00



## ARPS 15:00



80 gridpoints  $\approx$  240km

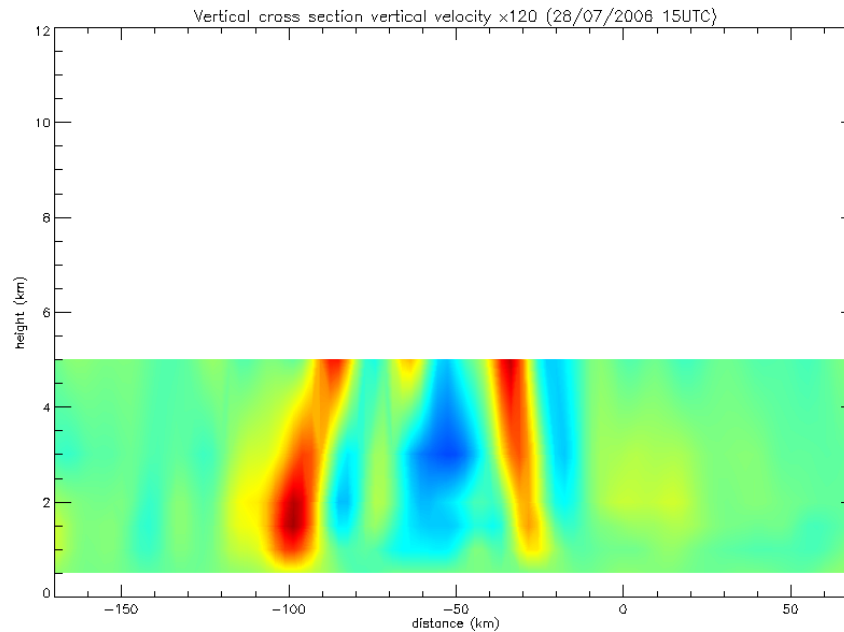


80 gridpoints  $\approx$  240km

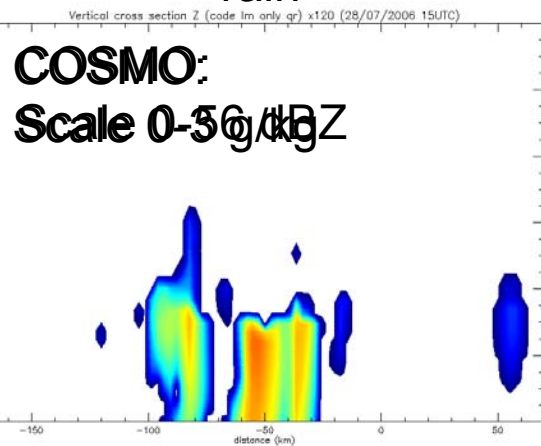
## COSMO 15:00



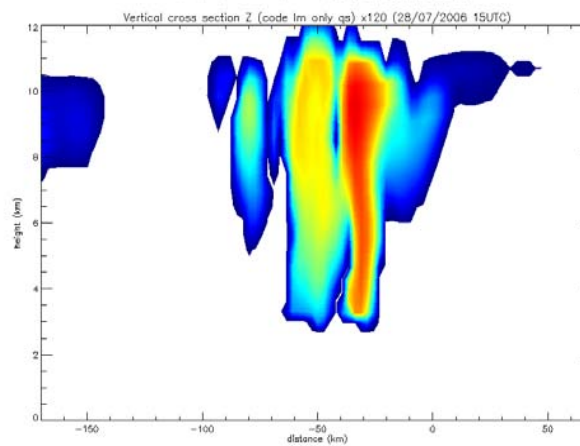
# Results: July case



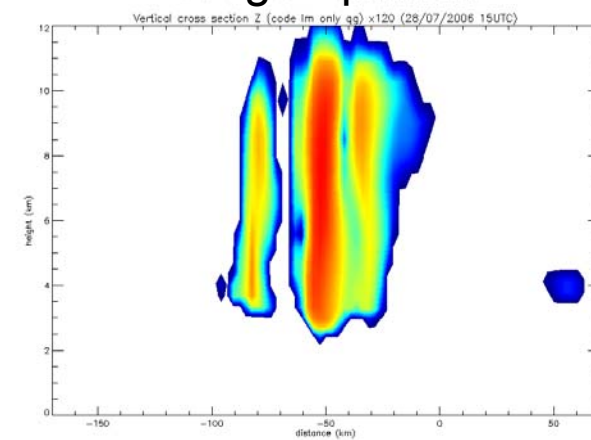
rain



snow



graupel

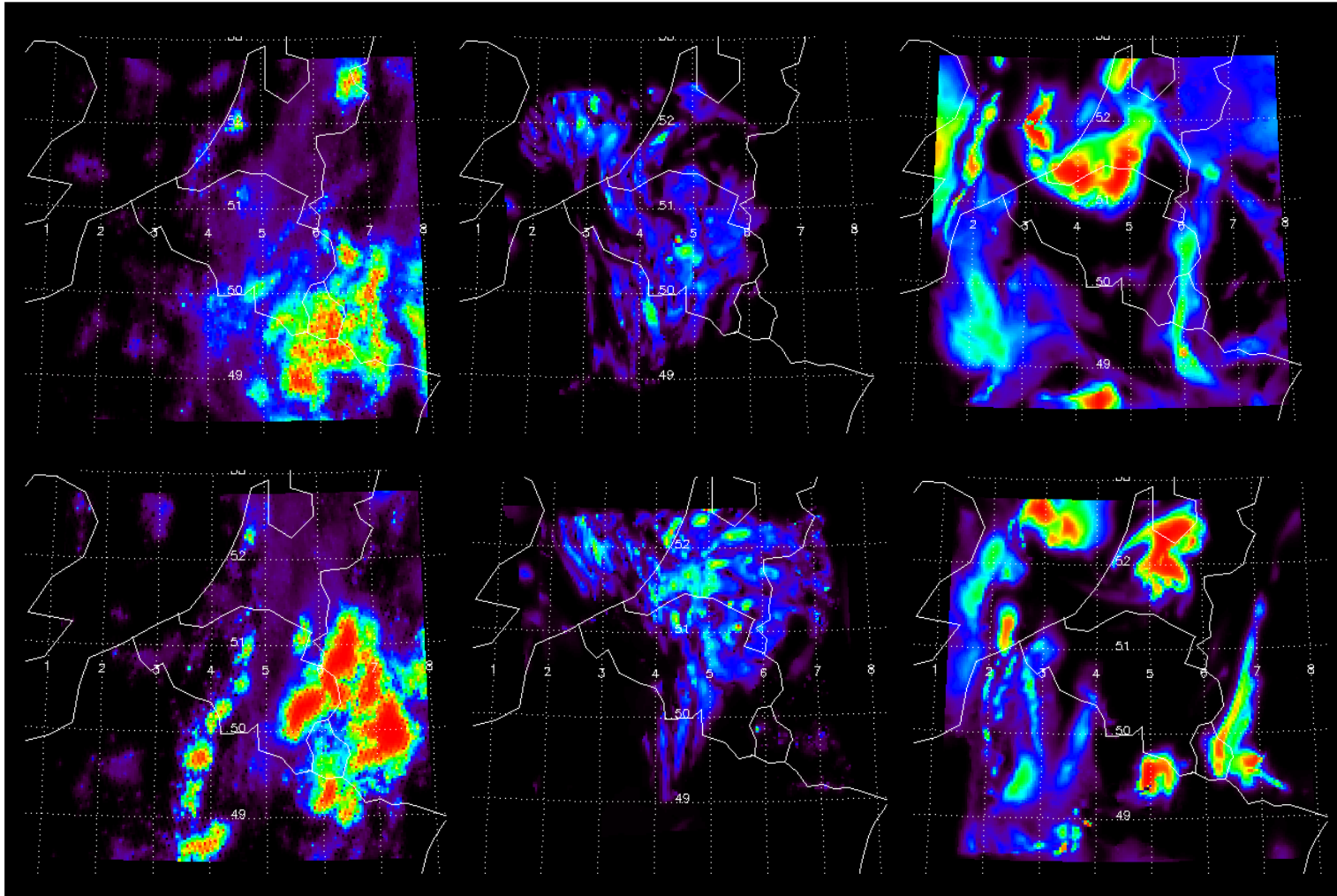


# Results: July case

MODIS 10:30

COSMO 9:00

ARPS 9:00



MODIS 12:15

COSMO 12:00

ARPS 12:00



0

25

50

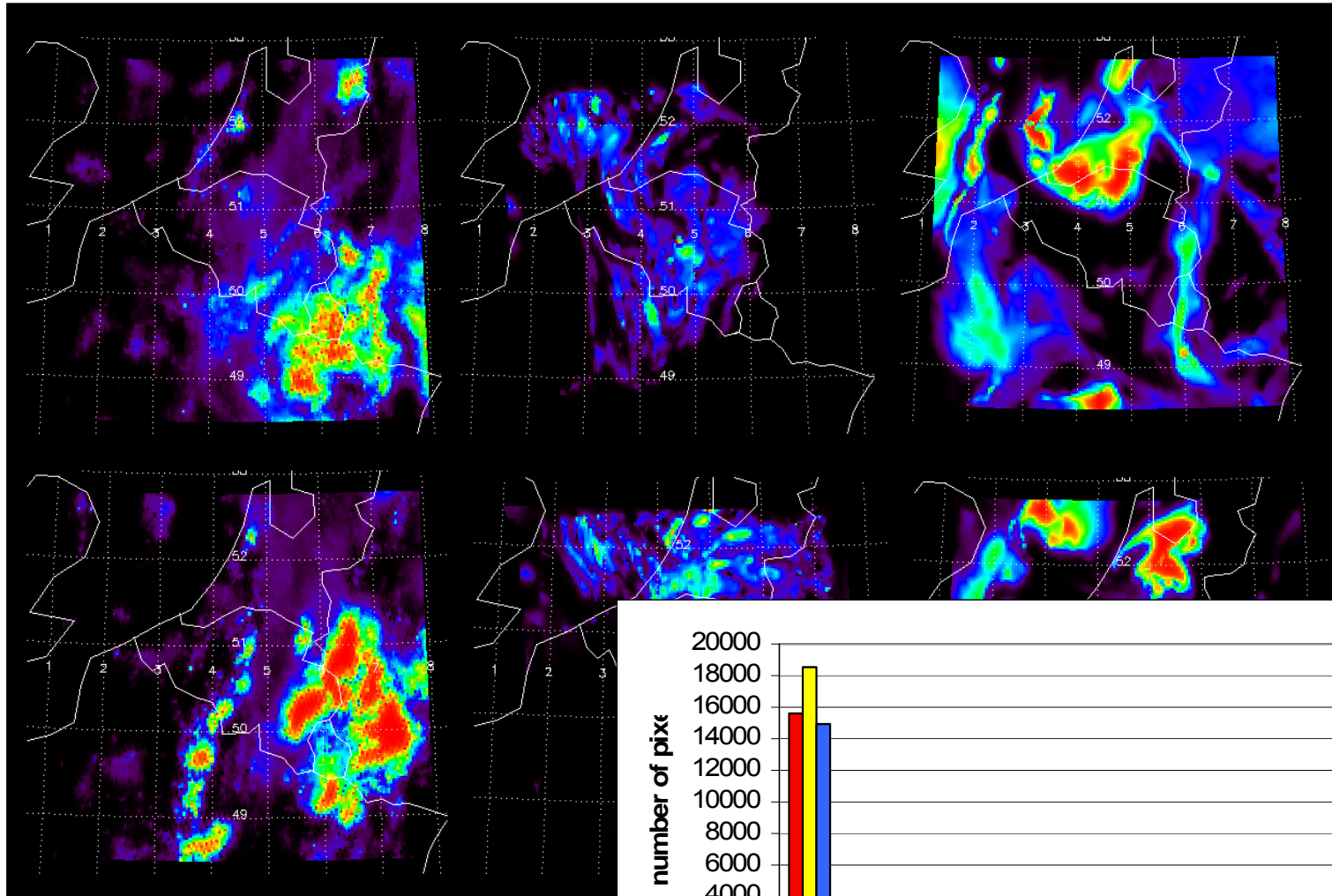
Cloud optical thickness July case

# Results: July case

MODIS 10:30

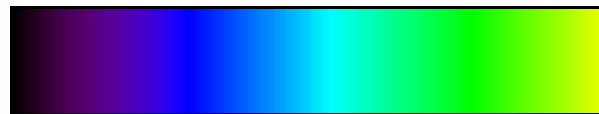
COSMO 9:00

ARPS 9:00



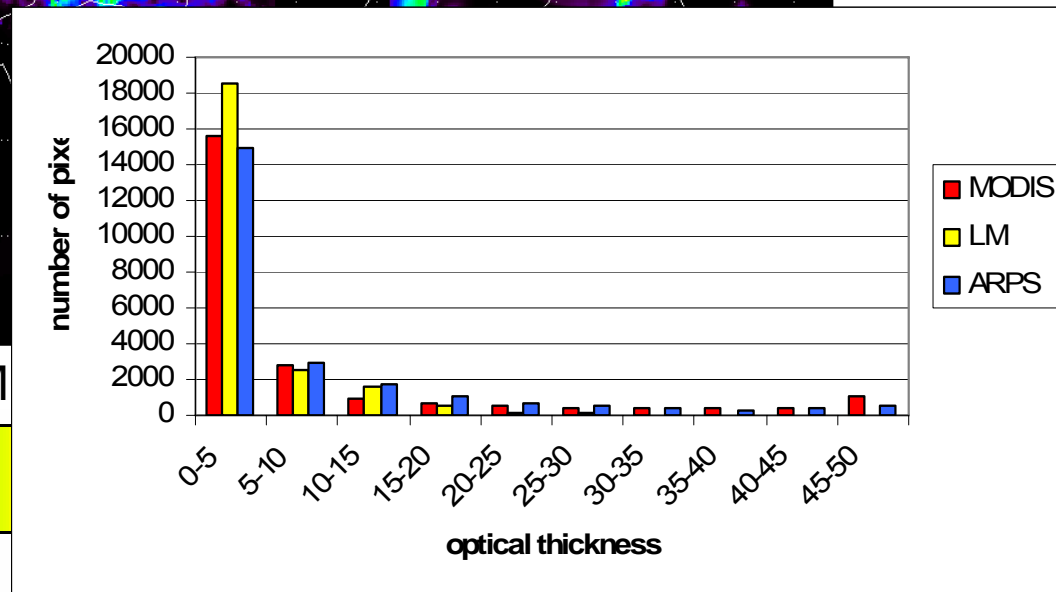
MODIS 12:15

COSMO



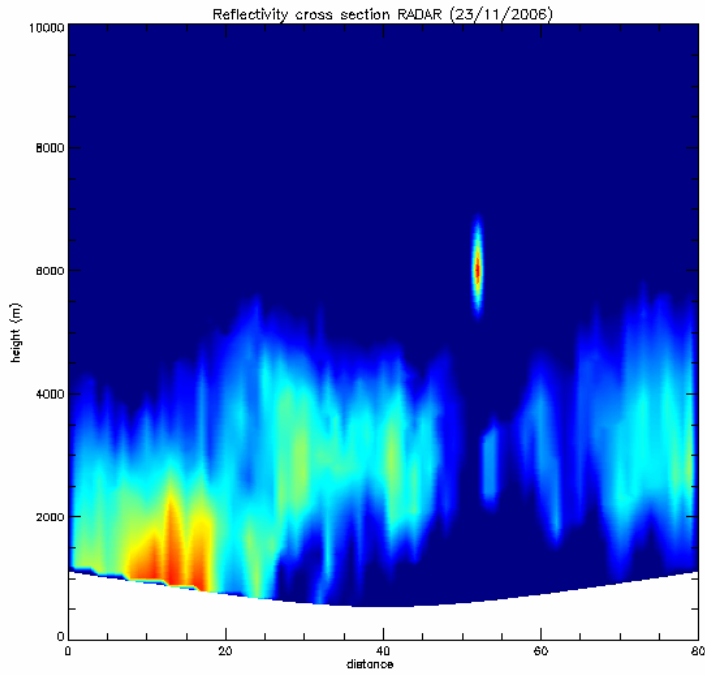
0

25

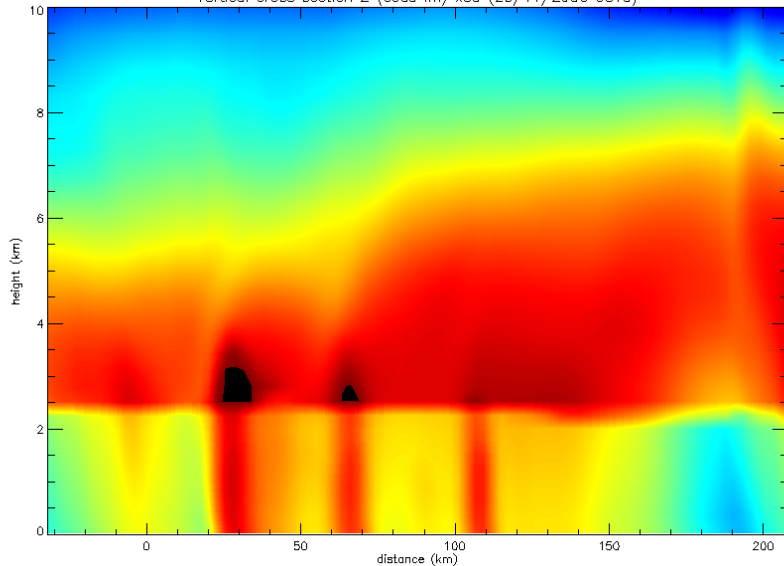


# Results: November case

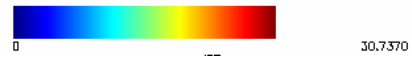
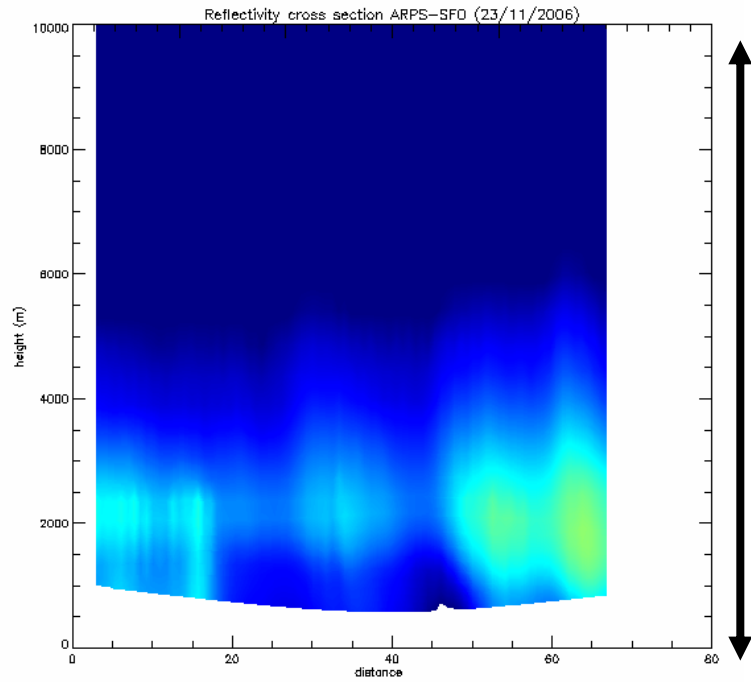
## Radar 6:00



Vertical cross section Z (code 1m) x50 (23/11/2006 6UTC)



## ARPS 6:00



0 31 dBZ

80 gridpoints  $\approx$  240km

## COSMO 6:00

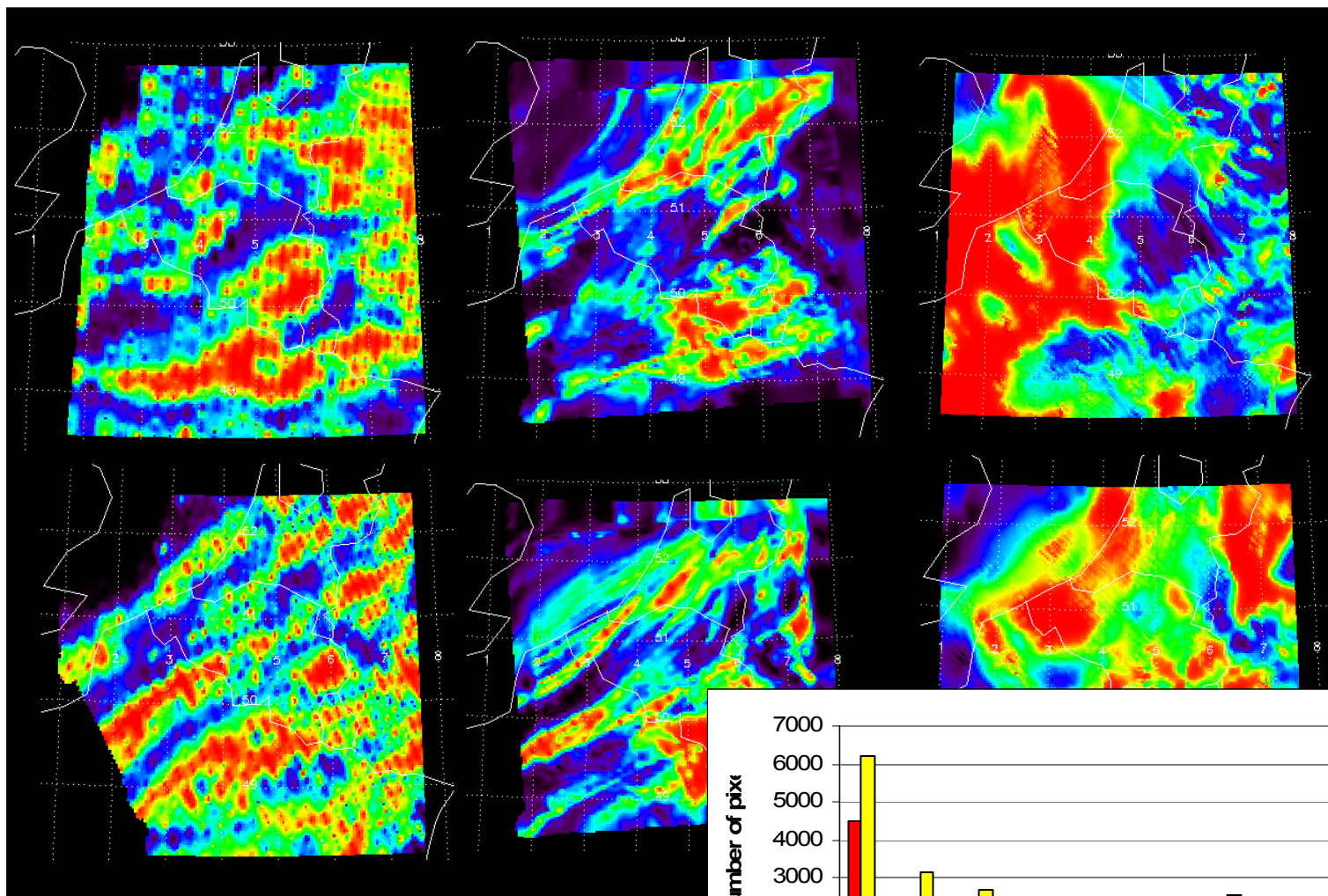


# Results: November case

MODIS 9:55

COSMO 9:00

ARPS 9:00



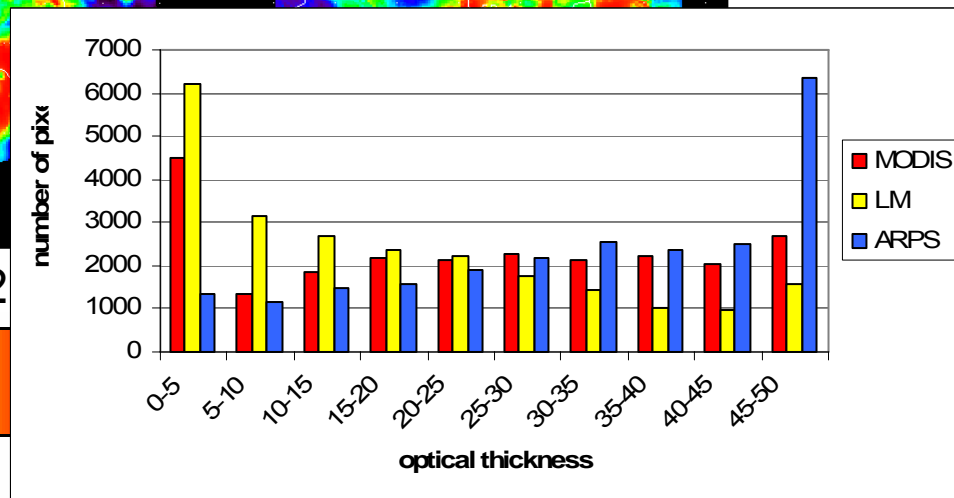
MODIS 11:40

COSMO 12



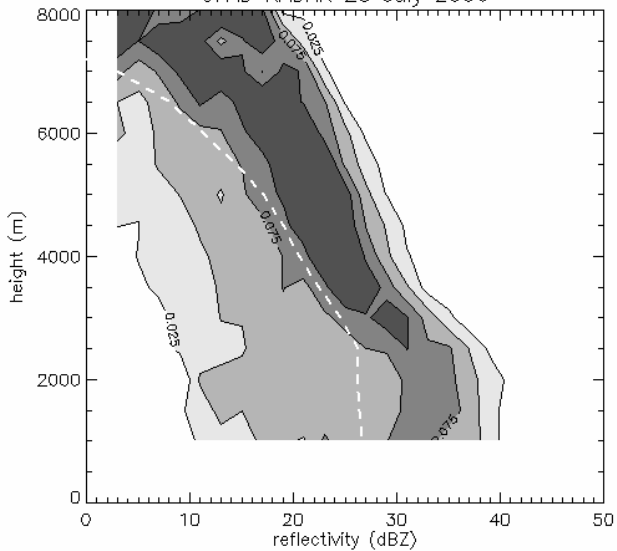
0

25

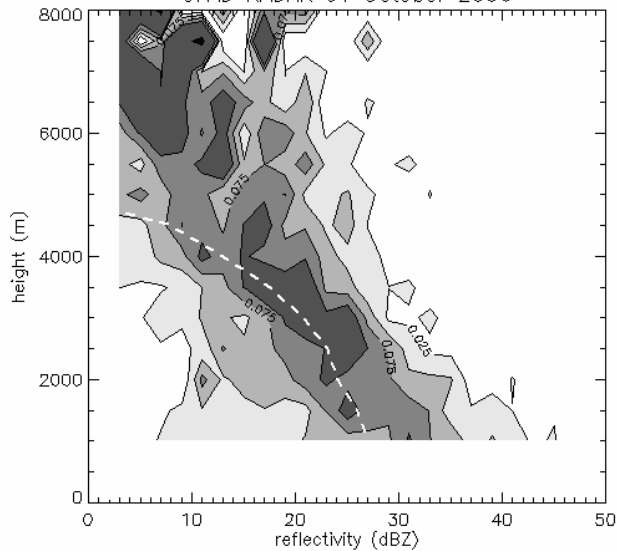


# Results

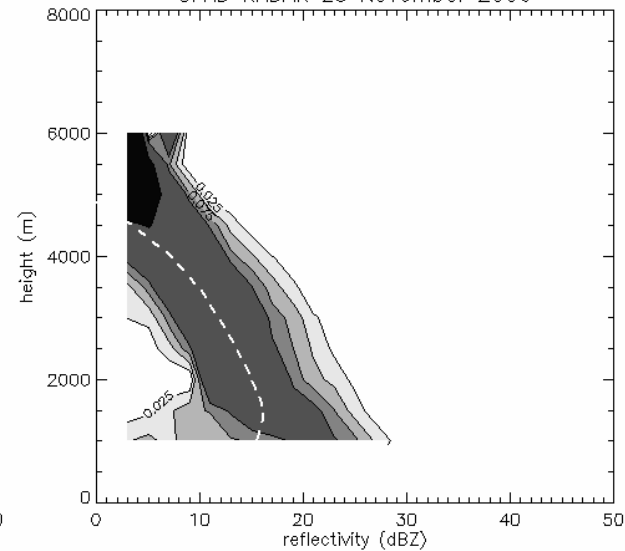
CFAD RADAR 28 July 2006



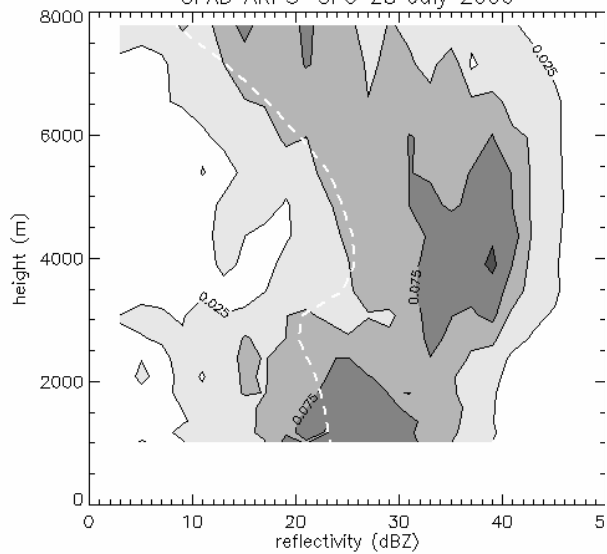
CFAD RADAR 01 October 2006



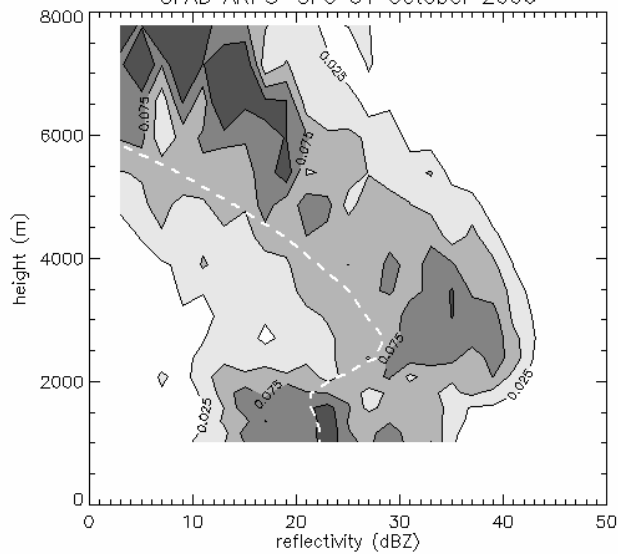
CFAD RADAR 23 November 2006



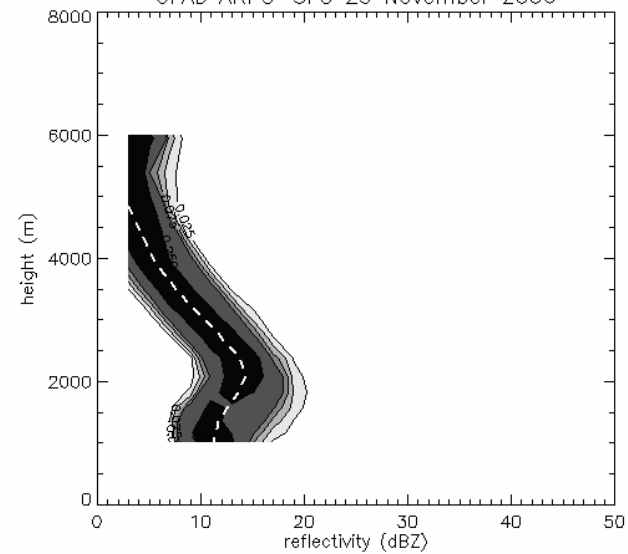
CFAD ARPS-SFO 28 July 2006



CFAD ARPS-SFO 01 October 2006



CFAD ARPS-SFO 23 November 2006



## Guidelines for model evaluation

- 3D radar volume data in combination with satellite data give useful complementary information for model evaluation
- By looking at volume data, additional information is available on:
  - the size of the convective cells
  - changes in the variability with height
- **ARPS** reflectivity very sensitive to amount of hail; this explains most discrepancies between radar and ARPS; too high hail amount just above the freezing level
- **COSMO2.8** version 3.21:
  - precipitating hydrometeors are separated
  - cloud optical thickness is too small for all these cases