

# An Inherently Mass-Conserving SISL Discretisation of the Nonhydrostatic Vertical Slice Equations

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# Overview

- Background
- Coupling conservation to momentum equations
- Governing equations
- Discretisation and iterative solver
- Results
- Conclusions



# Background

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- Mass conservation is important for climate simulations but is difficult for SL schemes.
- Coupling of a SL conservative scheme to the momentum equations has been shown for the SWE's, (Zerroukat *et al.* 2009).
- This is extended to a vertical slice model with orography.
- Constant reference height of the SWE becomes a vertically varying reference profile.

# Coupling the continuity equation to the momentum equations

- Standard (Lagrangian) continuity equation:

$$\frac{D\rho}{Dt} + \rho \nabla \cdot \mathbf{u} = 0$$

- Rewrite (more fundamentally) in conservative form:

$$\frac{D}{Dt} \left( \int_{\delta\mathcal{A}} \rho d\mathcal{A} \right) = 0$$

- $\mathcal{A}$  = vertical area element (control volume in 3D).
- Now apply SL scheme but key to stability is to couple directly to momentum equations

# Coupling the continuity equation to the momentum equations

- Define a reference density  $\rho^{\text{ref}}$  and rewrite:

$$\frac{D}{Dt} \left[ \int_{\delta\mathcal{A}} (\rho - \rho^{\text{ref}}) d\mathcal{A} \right] + \int_{\delta\mathcal{A}} \nabla \cdot (\rho^{\text{ref}} \mathbf{u}) d\mathcal{A} = 0$$

- Introduction of the reference profile allows a **divergence** term to be pulled off, coupling the momentum and continuity equations

# Coupling the continuity equation to the momentum equations

- Discretising along a trajectory and taking arrival volume  $\equiv$  control volume surrounding  $\rho \equiv \Delta\mathcal{A}$ :

$$\left[ (\rho - \rho^{\text{ref}}) + \alpha \Delta t \nabla \cdot (\rho^{\text{ref}} \mathbf{u}) \right]_A^{n+1} \equiv$$

$$\frac{1}{\Delta\mathcal{A}} \left\{ \int_{\delta\mathcal{A}} \left[ (\rho - \rho^{\text{ref}}) - \beta \Delta t \nabla \cdot (\rho^{\text{ref}} \mathbf{u}) \right] d\mathcal{A} \right\}_D^n ,$$

- Conservation is assured if RHS integration is conservative, (here the **SLICE** algorithm is used).

## Equations in height coordinates $z$

$$\frac{Du}{Dt} + c_p \theta \frac{\partial \pi}{\partial x} - fv + Fw = 0, \quad (1)$$

$$\frac{Dv}{Dt} + fu = 0, \quad (2)$$

$$\delta_V \frac{Dw}{Dt} + c_p \theta \frac{\partial \pi}{\partial z} + g - Fu = 0, \quad (3)$$

$$\frac{D}{Dt} \left( \int_{\delta \mathcal{A}} \rho d\mathcal{A} \right) = 0, \quad (4)$$

$$\frac{D\theta}{Dt} = 0, \quad (5)$$

$$\pi^{\frac{1-\kappa}{\kappa}} = \frac{R}{p_0} \rho \theta, \quad (6)$$

... in terrain following coordinates  $\eta(x, z)$

$$\frac{Du}{Dt} = -\frac{c_p\theta}{\partial z/\partial\eta} \left[ \frac{\partial}{\partial x} \left( \pi \frac{\partial z}{\partial\eta} \right) - \frac{\partial}{\partial\eta} \left( \pi \frac{\partial z}{\partial x} \right) \right] + fv - Fw, \quad (7)$$

$$\frac{Dv}{Dt} = -fu, \quad (8)$$

$$\delta_V \frac{Dw}{Dt} = -\frac{c_p\theta}{\partial z/\partial\eta} \frac{\partial\pi}{\partial\eta} - g + Fu, \quad (9)$$

$$\frac{D}{Dt} \left[ \int_{\delta\mathcal{A}} (\rho - \rho^{\text{ref}}) d\mathcal{A} \right] + \int_{\delta\mathcal{A}} \nabla \cdot (\rho^{\text{ref}} \mathbf{u}) d\mathcal{A} = 0, \quad (10)$$

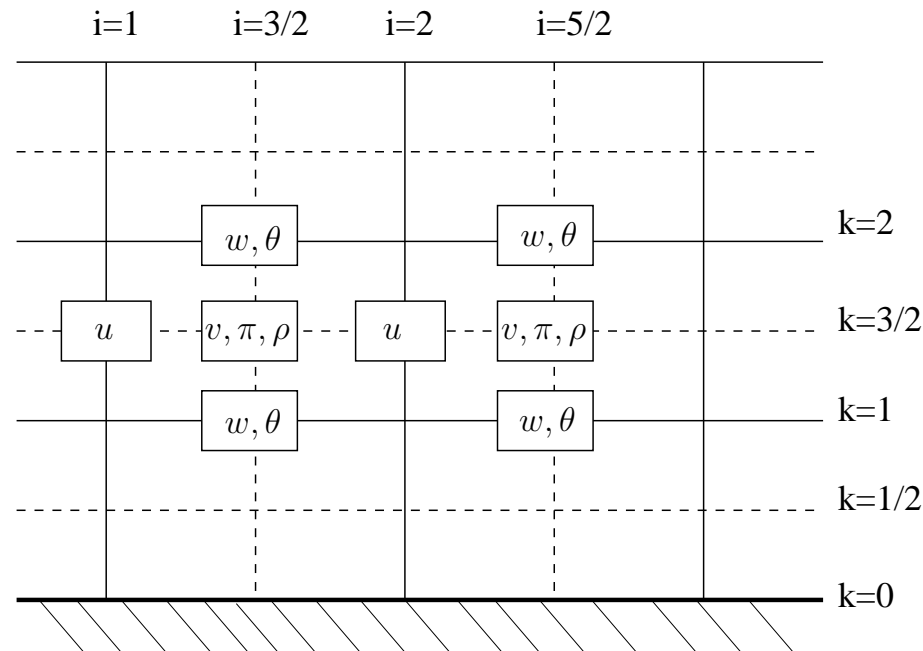
$$\frac{D}{Dt} (\theta - \theta^{\text{ref}}) + u \frac{\partial\theta^{\text{ref}}}{\partial x} + \dot{\eta} \frac{\partial\theta^{\text{ref}}}{\partial\eta} = 0, \quad (11)$$

$$\pi^{\frac{1-\kappa}{\kappa}} = \frac{R}{p_0} \rho\theta, \quad (12)$$



# Discretisation

- Equations are discretised in time using a two-time-level SISL scheme.
- Staggered Arakawa-C grid is used in the horizontal.
- Charney-Phillips grid is used in the vertical.
- Cubic-Lagrange polynomials used to interpolate fields to departure points.



# Model Features

- Coriolis terms are discretised following Thuburn and Staniforth (2004).
- Near boundary correction to departure point computations (Wood *et al.* 2009) is applied.
- A semi-Lagrangian discretisation of the relationship between  $\dot{\eta}$  and  $w$  is used (following Girard *et al.* 2005).
  - Eulerian:  $\dot{\eta} = \left(\frac{\partial z}{\partial \eta}\right)^{-1} \left(w - u \frac{\partial z}{\partial x}\right)$ .
  - Lagrangian:  $\dot{\eta} = \left(\frac{\partial z}{\partial \eta}\right)^{-1} \left(w - \frac{Dz}{Dt}|_{\eta}\right)$ .
- A variant of the absorbing layer (Klemp *et al.* 2008) is used to avoid spurious reflections from the domain lid.
- A multiply nested iterative approach (Zerroukat *et al.* 2009) is used to avoid any three-time-level extrapolations.



# Iterative algorithm

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## Do time-step loop:

- given  $(\dot{\eta}, \theta, w, u, v, \rho, \pi)^n$  at level  $n$

## Do outer-loop iteration:

- compute  $(x_D^n, z_D^n)$  using  $(u, w)^n$  and latest estimate for  $(u, w)^{n+1}$
- evaluate  $R_D^n$  for required fields  $R$

## Do inner-loop iteration:

- evaluate Coriolis and nonlinear terms
- solve Helmholtz problem for  $\pi^{n+1}$
- update  $(\dot{\eta}, \theta, w, u, v, \rho)^{n+1}$

Enddo

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# Computational examples

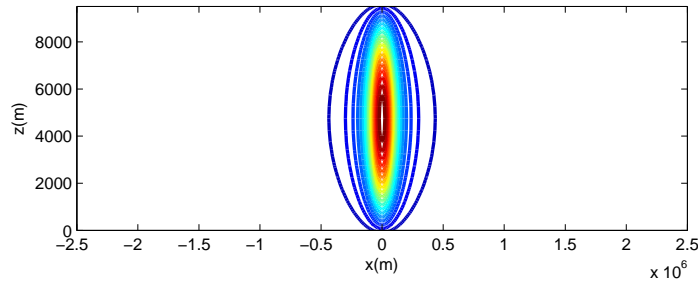
- Results tested on the standard test set:
  - Gravity waves.
  - Density Current.
  - Hydrostatic and nonhydrostatic mountain waves. [not shown]
  - Schär hill test.
- Linear  $\eta - z$  transformation is used:

$$\eta = \frac{z - z_S(x)}{z_T - z_S(x)}$$

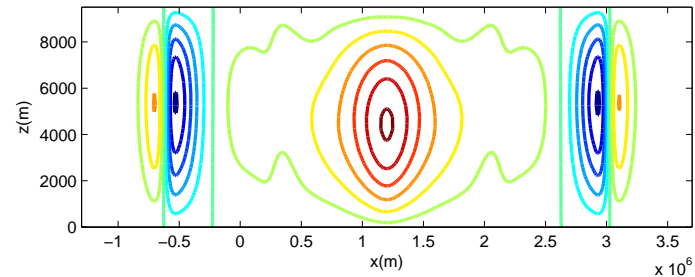
# Gravity Waves

- Hydrostatic inertia gravity waves

$t = 0$

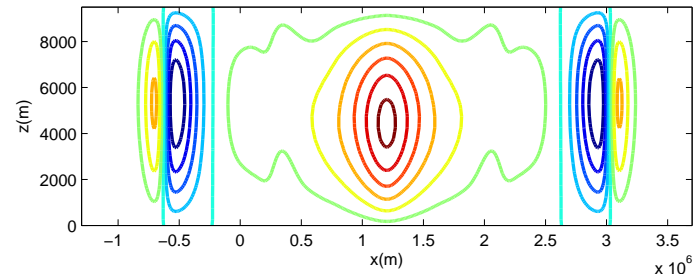
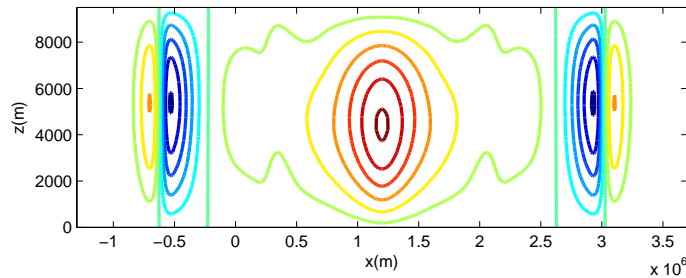


Mass-conserving  $t = 1000$  mins



SISL  $t = 1000$  mins,

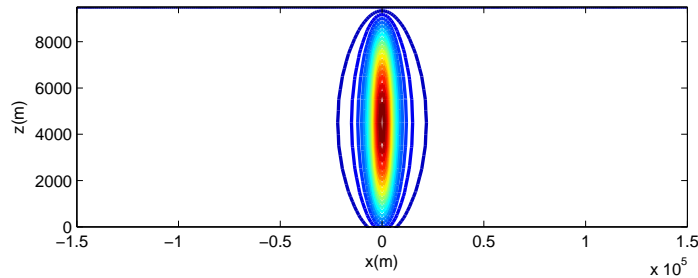
Mass-conserving,  $\delta_V = 0$ ,  $t = 1000$  mins



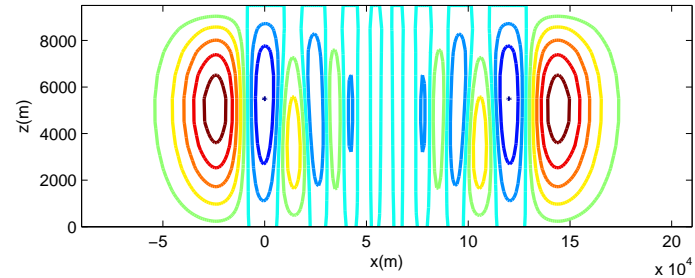
# Gravity Waves

- Nonhydrostatic gravity waves

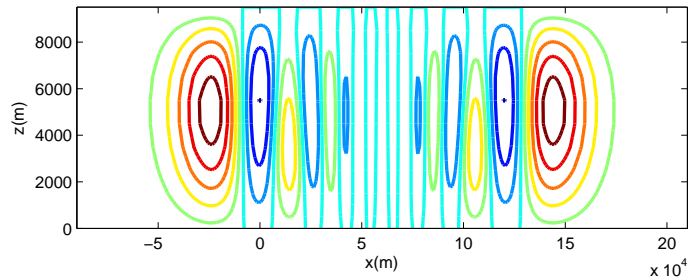
$t = 0$



Mass-conserving  $t = 1000$  mins



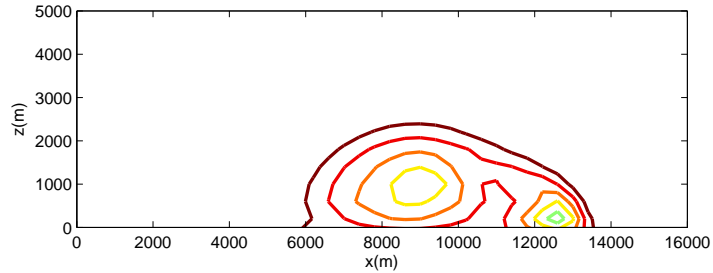
SISL  $t = 1000$  mins



# Density Current

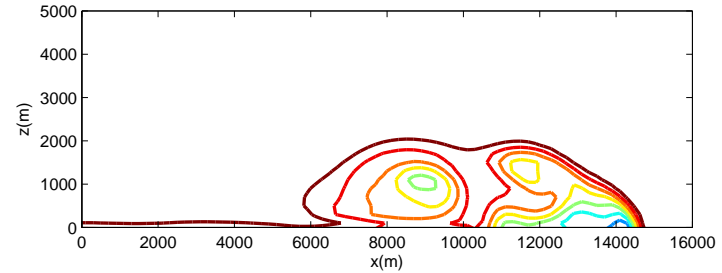
$\Delta x = 400m,$

$\Delta x = \Delta z = 400m, \Delta t = 4s$



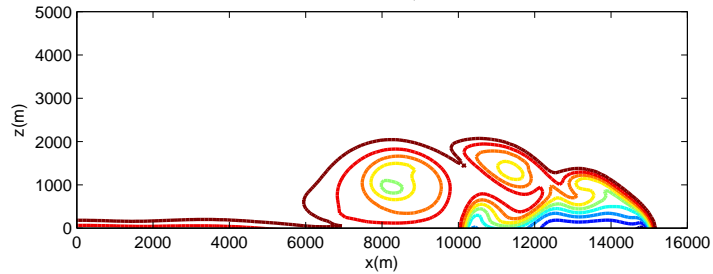
$\Delta x = 200m$

$\Delta x = \Delta z = 200m, \Delta t = 2s$



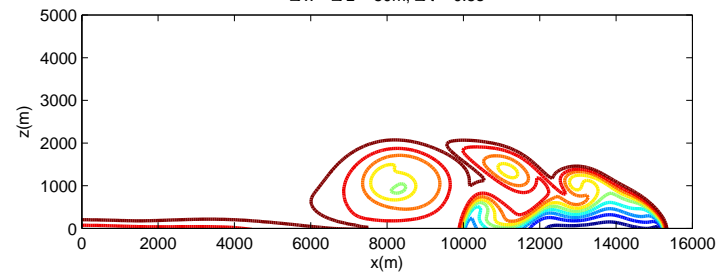
$\Delta x = 100m,$

$\Delta x = \Delta z = 100m, \Delta t = 1s$



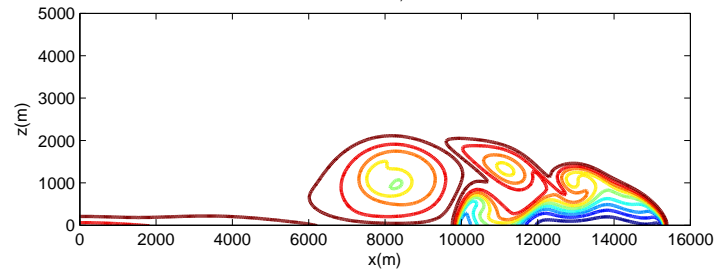
$\Delta x = 50m$

$\Delta x = \Delta z = 50m, \Delta t = 0.5s$



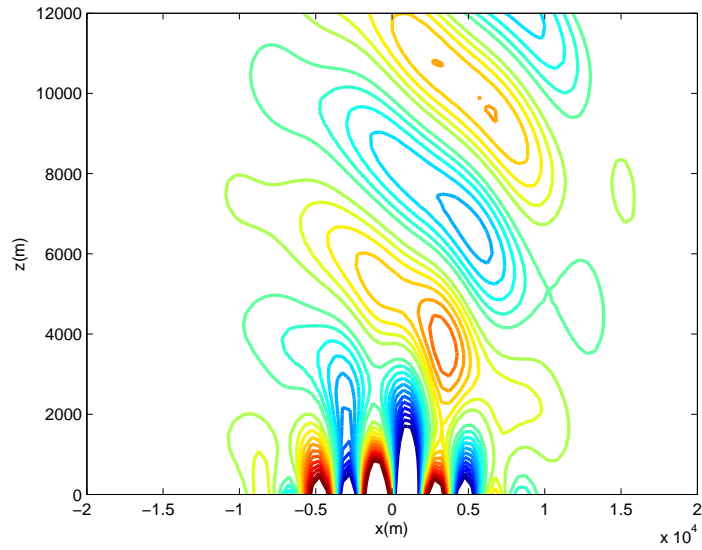
$\Delta x = 25m$

$\Delta x = \Delta z = 25m, \Delta t = 0.25s$



# Schär hill test

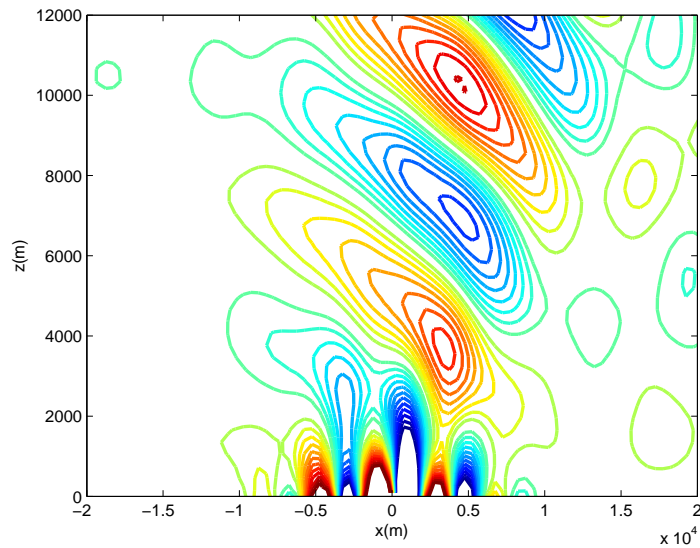
- Linear Analytic solution:



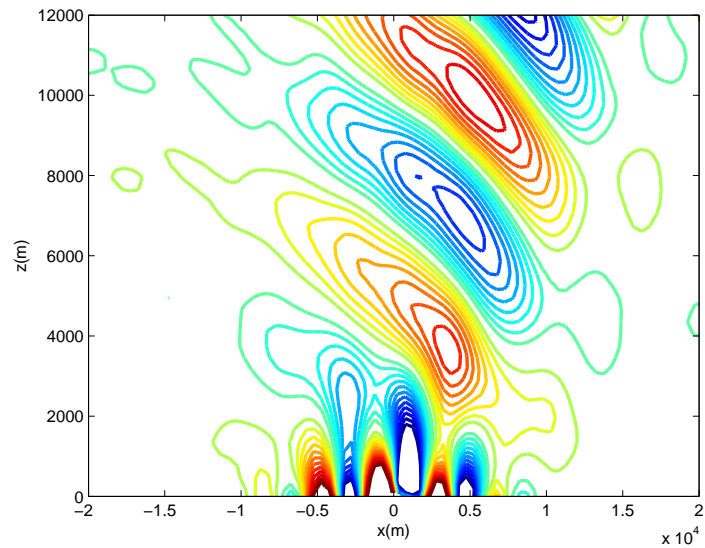


# Schär hill test

- Mass-Conserving mode:



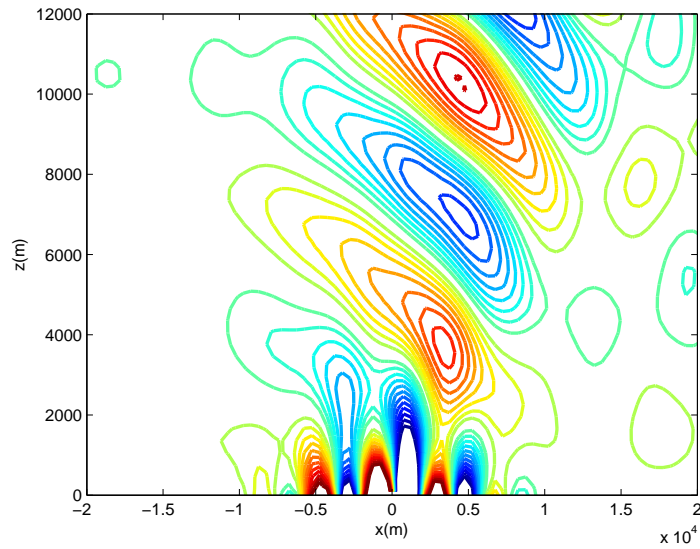
$$\Delta t = 8s,$$



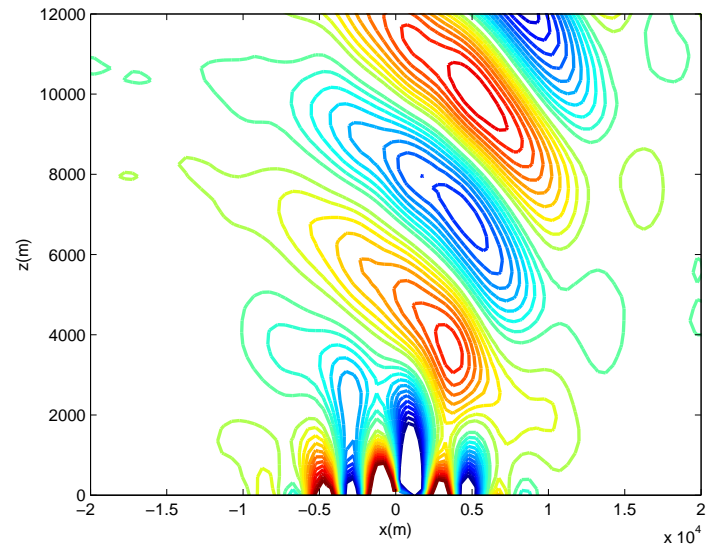
$$\Delta t = 40s$$

# Schär hill test

- Standard SISL mode:



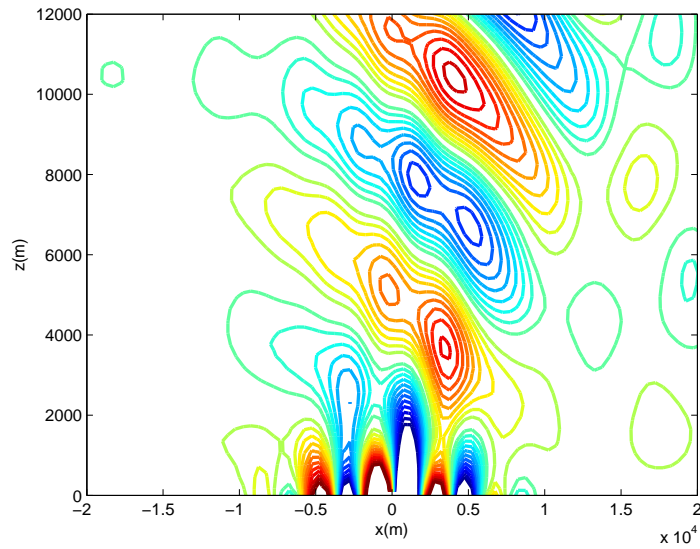
$$\Delta t = 8s,$$



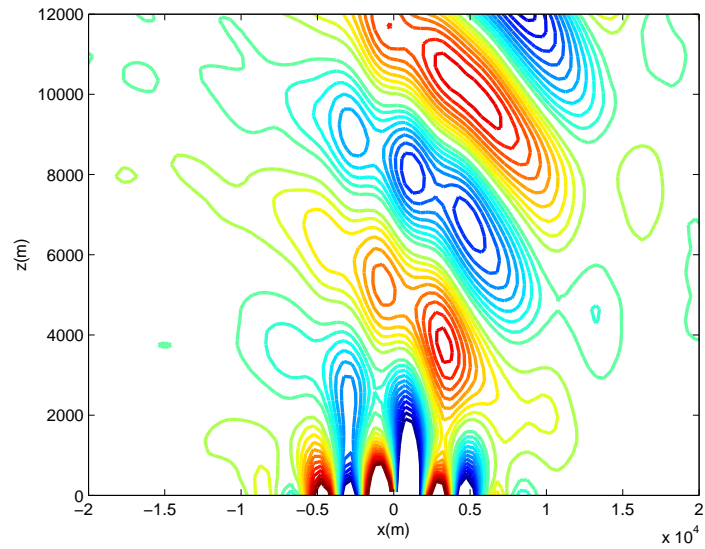
$$\Delta t = 40s$$

# Schär hill test

- Mass-Conserving mode with Eulerian  $\eta$  formulation:



$$\Delta t = 8s,$$



$$\Delta t = 40s$$



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# Conclusions

- Mass conserving algorithm (SLICE) has been coupled to a SL scheme
- Results for mass conserving mode are as good as standard SISL scheme and compare well to published results
- Lagrangian formulation of  $\dot{\eta}$  is important for numerical consistency.
- Remains to test coupling of the full 3D nonhydrostatic equations.



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## References

1. Melvin *et al.* "An inherently mass-conserving semi-implicit semi-Lagrangian discretisation of the nonhydrostatic vertical slice equations" - Submitted to QJRMS.
2. Zerroukat *et al.* "An inherently mass-conserving semi-implicit semi-Lagrangian discretisation of the shallow water equations on the sphere" - QJRMS **135**: 1104-1116.
3. Zerroukat *et al.* "SLICE: A Semi-Lagrangian Inherently Conserving and Efficient scheme for transport problems" - QJRMS **128**: 2801-2820.
4. Standard test set for 2D vertical slice dynamical cores:  
[http://box.mmm.ucar.edu/projects/srnwp\\_tests/index.html](http://box.mmm.ucar.edu/projects/srnwp_tests/index.html).