Model for Prediction Across Scales (MPAS)







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Future Weather/Climate Atmospheric Dynamic Core

Problems with lat-lon coordinate for global models

- Pole singularities require special filtering
- Polar filters do not scale well on massively parallel computers
- Highly anisotropic grid cells at high latitudes

Consideration of alternative spatial discretizations:



Priority Requirements: • Efficient and scales well on massively parallel computers

- Well suited for cloud (nonhydrostatic) to global scales
- Capability for local grid refinement and regional domains
- Conserves at least mass and scalar quantities



Hexagonal C-Grid

Why a hexagonal grid?

- Removes the pole singularity
- Most isotropic (wave propagation).
- Provides good conservation properties in a finitevolume formulation
- Hexagonal grid permits larger explicit time steps.
- Readily generalized to arbitrarily structured grids

Why a C-grid staggering?

- Provides the highest accuracy for the fast (gravitywave) modes
- Provides twice the resolution of the A grid, and avoids the parasitic mode





Hexagonal C-Grid Problem: Non-Stationary Geostrophic Mode





Cell center is cell center-of-mass

Edges of dual grid intersect edges of primary grid at right angles.

Conjecture: smooth refinement on conformal meshes should mitigate many refinement problems.



(Michael Duda, MMM)

 A_{Aa}

A_{Ca}

Ac

 u_{2}

A_{Ba}

 A_{B}

Generalization for Arbitrarily Structured Grids

In discrete analogue of vorticity equation (ξ_t =-f δ_a), the divergence δ_a on the Delaunay triangulation must be identical to the divergence δ_A on the Voronoi hexagons used in the height equation (h_t =-H δ_A):

$$A_a\delta_a = A_{Aa}\delta_A + A_{Ba}\delta_B + A_{Ca}\delta_C$$

Construct tangential velocities from weighted sum of normal velocities on edges of adjacent grid cells.

$$d_e \underbrace{u_e^\perp}_j = \sum_j w_e^j l_j \underbrace{u_j}_j$$

Energy conservation achieved by requiring $w_e^j = -w_j^e$.



 u_7

 U_6

 u_{5}

Ac

A₄



Analytic Results for the Nonlinear Shallow-Water Equations

- 1. Stationary geostrophic mode is recovered.
- 2. Potential vorticity is conserved to round-off. PV is compatible with an underlying thickness evolution equation.
- 3. Total energy is conserved to within time truncation.
 - a. Coriolis force is energetically-neutral
 - b. Transport of KE is conservative
 - c. KE/PE exchange is equal and opposite.
- 4. It appears that potential enstrophy can be dissipated.

Results hold for a wide class of meshes: Lat/Lon, Voronoi Tessellations, Delaunay Triangulation and Conformally-mapped cubed sphere meshes.



Reconstructing the Nonlinear Coriolis Force

(The nonlinear Coriolis force Q_e^{\perp} is the PV flux perpendicular to the velocity)





The nonlinear Coriolis force will be energy neutral for any \bar{q}_{ej}

Reconstructing the Nonlinear Coriolis Force

(The nonlinear Coriolis force Q_e^{\perp} is the PV flux perpendicular to the velocity)





Upstream bias leads to dissipation of potential enstropy.

Anticipated Potential Vorticity Method (Sadourny & Basdevant, JAS 1985)

Shallow-Water Test Case 5: Flow over an Isolated Mountain

40962 cells, dx ~ 120 km 2nd order differencing **RK4** time differencing 10^{-2} Linf $\mathsf{L}_{\mathsf{inf}}$ and L_2 Norms of Thickness Tomita et al. (2001) 10⁻³ Linf 12 10⁻⁴ Lipscomb and Ringler (2005) Results with quasi-uniform meshes: 10⁻⁵ 50 250 100 150 200 300 350 400 0 New scheme is competitive Simulation Time (hours) with existing models.



SWTC#5: Convergence Rate as Measured by L₂ and L_{inf} Norms





SWTC#5: 100 Day Integration, Variable-Res Mesh

Computing potential vorticity with an upstream bias allows potential enstrophy dissipation while conserving energy

- d ~ 40 km over terrain
 - ~ 120 km elsewhere





SWTC#5 Relative Vorticity at 50 Days: Impact of Variable Mesh

Energy conserving and potential enstrophy dissipating simulation



Constant mesh, d ~ 60 km Variable Mesh, d ~ 40 - 120 km



SWTC#5: Globally-Averaged Potential Enstrophy Evolution





SWTC#5 Relative Vorticity at 50 days: Impact of Dissipating Potential Enstrophy



Limited-Area Cloud Model on Hexagonal C-Grid



Research Progress:

- Constructed a 3-D limited-area hexagonal-grid cloud model based on WRF/ARW numerics to evaluate performance.
- Documented that hexagonal-grid cloud simulations are at least as accurate and computationally more efficient than those on a conventional rectangular grid.





3-D Supercell Simulation ~500 m Horizontal Grid

Rectangular Grid

Hexagonal Grid



Vertical velocity contours at 1, 5, and 10 km (c.i. = 3 m/s)
30 m/s vertical velocity surface shaded in red
Rainwater surfaces shaded as transparent shells
Perturbation surface temperature shaded on baseplane



3-D Supercell Simulation - Maximum Vertical Velocity





Summary

New SW solver for SVCT C-grid

- Recovers stationary geostrophic mode.
- Conserves PV to round off.
- Conserves energy to time truncation.
- Solutions comparable to existing SW solvers, and no dissipation needed for standard SW test cases.
- Variable-resolution grid results are encouraging.
- Nonhydrostatic solver as accurate and efficient as existing cloud models
- Parallel SWL solver implemented for an unstructured grid.

Next Steps: MPAS development

- Hydrostatic 3D SVCT solver (based on SW solver).
- Nonhydrostatic solver (based on hydrostatic solver).
- Higher-order transport schemes (Bill Skamarock next talk).
- Continued variable-resolution grid testing.



