

The formation of wider and deeper clouds through cold-pool dynamics

Linda Schlemmer, Cathy Hohenegger

Max-Planck-Institute for Meteorology, Hamburg

2013-05-14



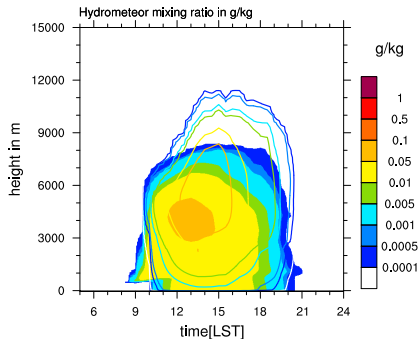
Max-Planck-Institut
für Meteorologie

Hans-Ertel-Zentrum für Wetterforschung
Deutscher Wetterdienst



Diurnal cycle of convection over land

How to get deep clouds??



Promotion of deeper and wider clouds that are less affected by entrainment: **cold pools**

(Khairoutdinov and Randall, 2006, Kuang and Bretherton, 2006, Böing et al., 2012)

Cold pools, driven by evaporation and melting and precipitation

- 1 Dynamical lifting of convective cells along the edges of cold pools (e.g. Lima et al., 2008)
- 2 Occurrence of a band with high equivalent potential temperature and water vapor along the edge of cold pools (Tompkins, 2001)

How do cold pools promote wider clouds? Different in different environments? Shallow <-> deep convection

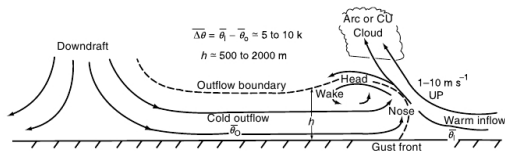


FIGURE 8.23 Schematic structure of a gust front. (Adapted from Goff (1976), Fankhauser (1982), and Wakimoto (1982))

Figure from Cotton (2011)

Large Eddy Simulations (LES)

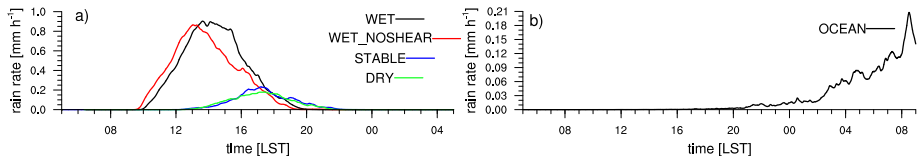
UCLA-LES

- Anelastic core
- Smagorinsky-type subgrid-scale diffusion
- $\Delta x = \Delta y = 250$ m, Δz stretched grid with a minimum of 70 m
- Domain size 250×250 km, model top at 21020 m
- Two-moment ice microphysics scheme (Seifert and Beheng, 2006)
- Interactive radiation scheme (Pincus and Stevens, 2009)
- Surface heat fluxes prescribed

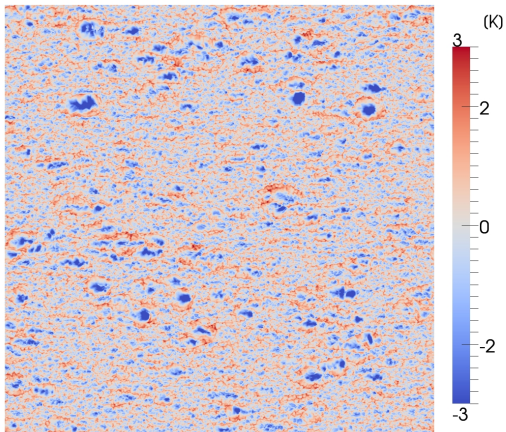
Different environments: cases

- Mid-latitude continental case (cf. Schlemmer et al., 2011)
 - WET
 - WET_noshear
 - STABLE
 - DRY

- Maritime case (cf. Waite and Khoudier, 2010)
 - OCEAN

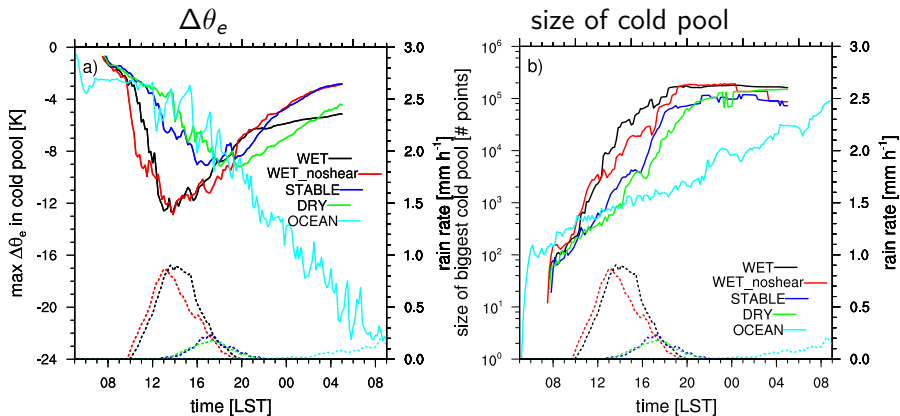


Cold pools

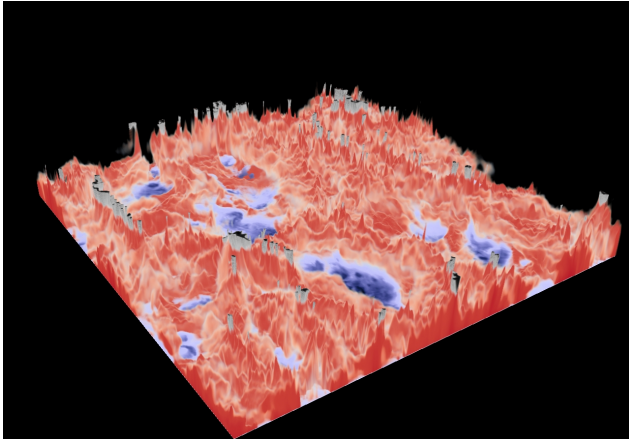


$\Delta\theta_e$ at 35 m height at
12 LST

Cold pools under different conditions

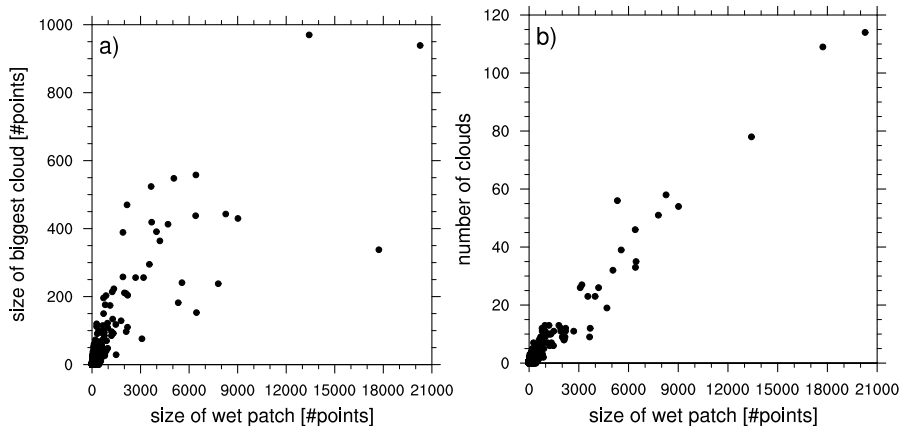


Signature in moisture field

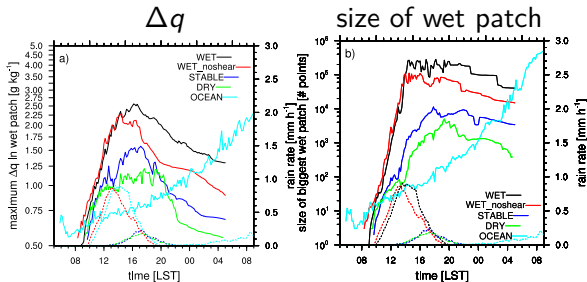


Dry areas in the cold
pools and moist
damns or patches
around them
Clouds are located on
these “wet patches.”

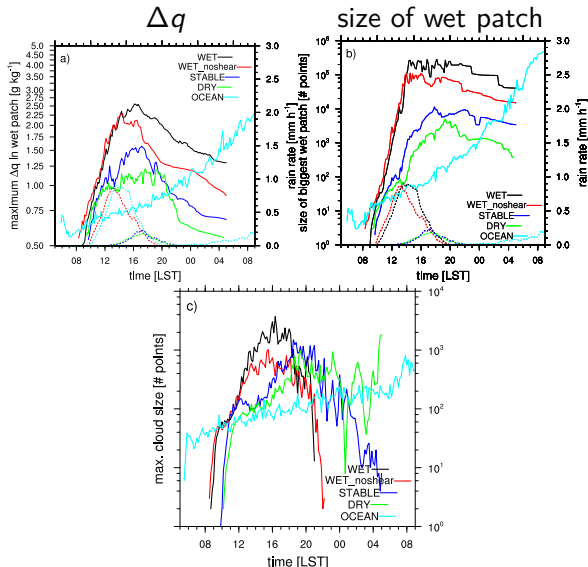
Wet patches and clouds



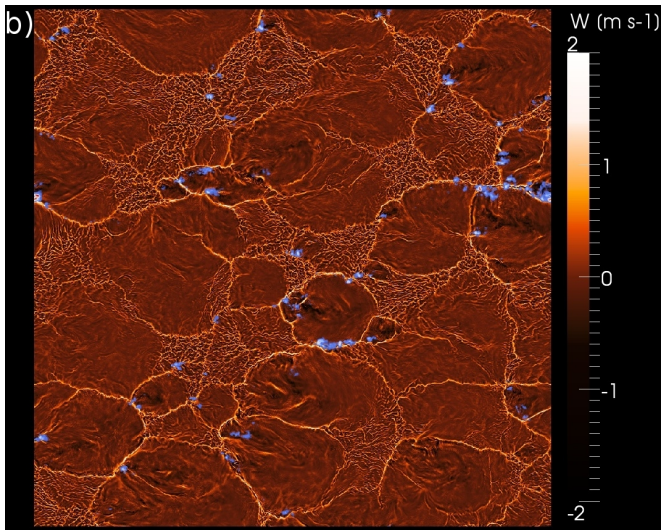
Wet patches and cloud sizes



Wet patches and cloud sizes



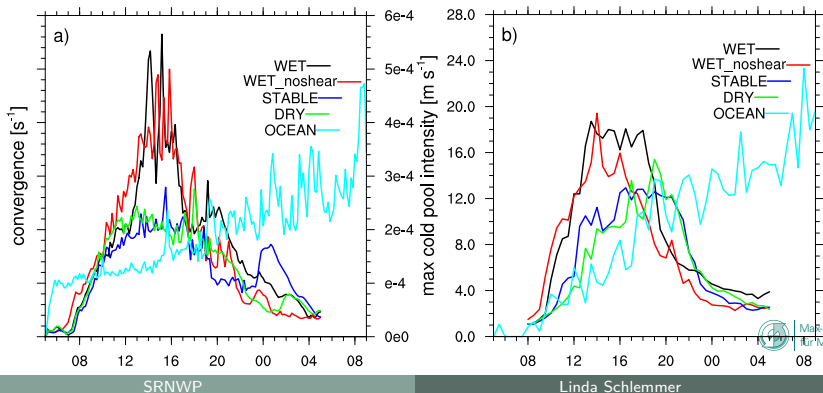
Dynamical lifting



Convergence & propagation speed

cold pool intensity c_i (Rotunno, 1988):

$$c_i^2 = 2 \cdot \int_0^{H_{cp}} (-B) dz$$



Cold pools - cloud width

- Both a dynamical triggering and an accumulation of moisture by cold pools seems to be important for the generation of wider convective cells, mechanisms hard to disentangle.
- Large wet patches yield wider clouds
- Cold pools act in a similar manner in different environments but are stronger and larger in a more humid environment

Pathway to a parameterization

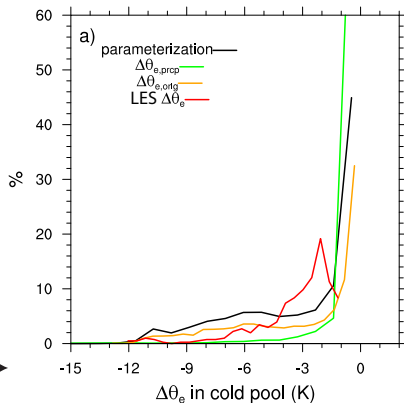
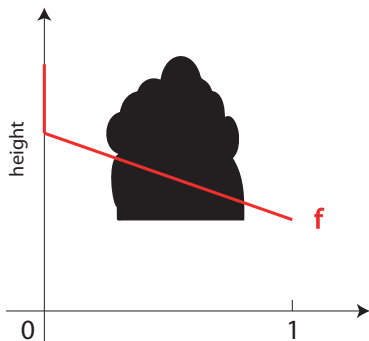
Successful attempts to include the effects of cold-pool wakes and associated dynamical triggering into parametrization schemes (e.g. Qian,1998, Mapes,2000, Rio,2009).

⇒ include the role of a moisture aggregation.

Pathway to a parameterization

- 1 Predict θ_e depression in cold pool:

$$\Delta\theta_e = f \cdot \Delta\theta_{e,prcp} + (1 - f) \cdot \Delta\theta_{e,orig} \quad (1)$$

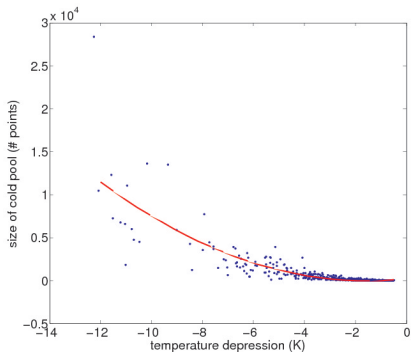


Pathway to a parameterization

- ① Predict θ_e depression in cold pool:

$$\Delta\theta_e = f \cdot \Delta\theta_{e,prcp} + (1 - f) \cdot \Delta\theta_{e,orig} \quad (1)$$

- ② Relate size of cold pool to temperature depression: $A_{cp} = c_{fit} \cdot \Delta\theta_e^2$



Pathway to a parameterization

- 1 Predict θ_e depression in cold pool:

$$\Delta\theta_e = f \cdot \Delta\theta_{e,prcp} + (1 - f) \cdot \Delta\theta_{e,orig} \quad (1)$$

- 2 Relate size of cold pool to temperature depression: $A_{cp} = c_{fit} \cdot \Delta\theta_e^2$
- 3 Assume: size of wet patch = size of dry area = size of cold pool
 $A_{cp} = A_{wp}$

Pathway to a parameterization

- 1 Predict θ_e depression in cold pool:

$$\Delta\theta_e = f \cdot \Delta\theta_{e,prcp} + (1 - f) \cdot \Delta\theta_{e,orig} \quad (1)$$

- 2 Relate size of cold pool to temperature depression: $A_{cp} = c_{fit} \cdot \Delta\theta_e^2$
- 3 Assume: size of wet patch = size of dry area = size of cold pool
 $A_{cp} = A_{wp}$
- 4 Size of biggest cloud related to size of wet patch

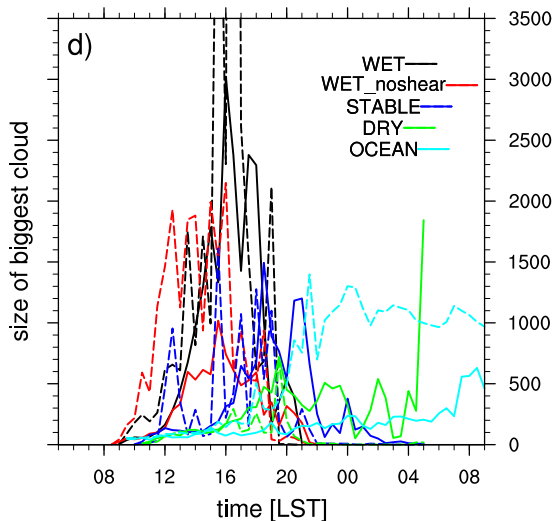
Pathway to a parameterization

- 1 Predict θ_e depression in cold pool:

$$\Delta\theta_e = f \cdot \Delta\theta_{e,prcp} + (1 - f) \cdot \Delta\theta_{e,orig} \quad (1)$$

- 2 Relate size of cold pool to temperature depression: $A_{cp} = c_{fit} \cdot \Delta\theta_e^2$
- 3 Assume: size of wet patch = size of dry area = size of cold pool
 $A_{cp} = A_{wp}$
- 4 Size of biggest cloud related to size of wet patch
- 5 Size of cloud \Rightarrow modified entrainment rate

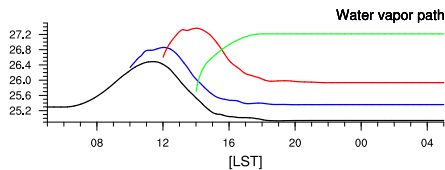
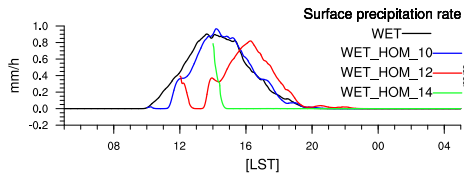
Pathway to a parameterization



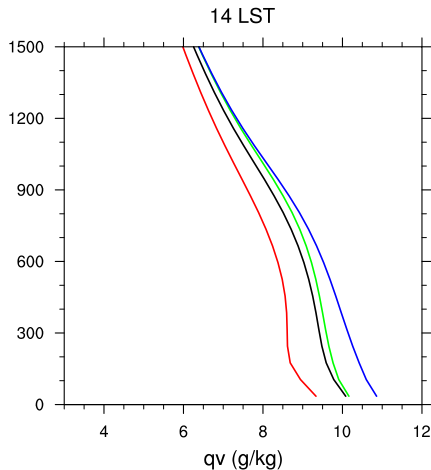
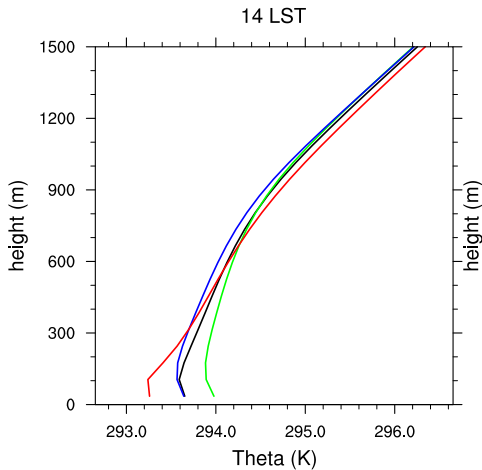
Conclusions

- Both a dynamical triggering and an accumulation of moisture by cold pools seems to be important for the generation of wider convective cells, mechanisms hard to disentangle.
- Cold pools act in a similar manner in different environments but are stronger and larger in a more humid environment
- Promising attempts towards a parameterization scheme

Homogenization



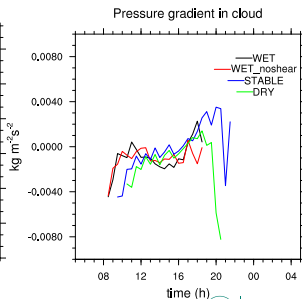
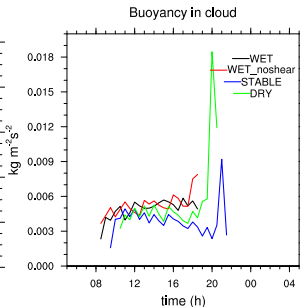
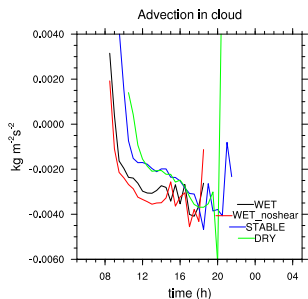
Vertical Profiles



Momentum equation

(neglect viscous terms and Coriolis terms)

$$\rho_0 \cdot \frac{\partial w}{\partial t} = - \left(\frac{\partial}{\partial x} (\rho_0 u w) + \frac{\partial}{\partial y} (\rho_0 v w) + \frac{\partial}{\partial z} (\rho_0 w w) \right) - \frac{\partial p'}{\partial z} + \rho_0 \cdot B$$



Momentum equation

(neglect viscous terms and Coriolis terms)

$$\rho_0 \cdot \frac{\partial w}{\partial t} = - \left(\frac{\partial}{\partial x} (\rho_0 u w) + \frac{\partial}{\partial y} (\rho_0 v w) + \frac{\partial}{\partial z} (\rho_0 w w) \right) - \frac{\partial p'}{\partial z} + \rho_0 \cdot B$$

